

Mineralogical and chemical features of gangue phases in relation to hydrothermal mineralization and their host rocks

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Abstract. Gangue minerals from hydrothermal deposits (apatite and carbonates) and their host rocks from three different volcanic areas (SE Spain, La Serena and Melipilla in Coastal Range of Chile) have been studied for broad elemental compositions. Carbonate units at the Au-Cu Palai-Islica epithermal deposit are Fe-Mn-bearing, with a slightly higher concentration of these elements in the orebody than in adjoining hydrothermal alteration zones. Apatite has a composition that correlates with its origin and with hydrothermal processes. Thus, volcanic apatite is Cl-rich, whereas apatite from the associated orebody is almost pure fluorapatite. Furthermore, apatite from hydrothermally altered volcanic rocks has a transitional composition between volcanic and ore-related apatite. Samples of carbonate from Mn, Cu(-Ag) and Ba-Ag deposits in the La Serena area are all Mn-bearing calcite. In addition, Mn(Fe)-poor and Mg-rich calcite is common in low-grade Mn areas. Similar features have been found at the Melipilla Cu (Cu-Ag) deposit where epithermal calcite is also enriched in Mn(+Fe) with respect to non-mineralized veins, carbonate host rock, and metamorphic mineralization. In summary, F in apatite and Mn have been introduced in to carbonate from the La Serena area by hydrothermal fluids and could provide an index of hydrothermal ore-forming activity.

Keywords. hydrothermal, gangue, carbonate, apatite, Spain, Chile

1 Introduction

Gangue minerals constitute an important proportion of hydrothermal deposits. For example, fluid inclusions and stable isotope compositions in these minerals can provide much valuable information on the genesis of ore deposits for researchers of economic geology. This contribution focuses attention on the chemical composition of apatite and carbonates from different hydrothermal deposits. This information sheds light on the genetic interpretation of the deposits studied, and provides pointers for mineral exploration.

2 Palai-islica: epithermal Au-Cu volcanic-hosted deposit, SE Spain

Palai-Islica is an Au-Cu epithermal deposit hosted by calc-alkaline andesitic to dacitic rocks of the Cabo de Gata-Cartagena volcanic belt, which is a magmatic expression of the Betic Orogeny formed during extension of a thickened lithosphere (e.g. Turner et al. 1999). The deposit consists of quartz veins with sulfides, sometimes stockwork-like, and massive replacement-type mineralization in zones of intense silicification. The volcanic host-rocks are commonly hydrothermal altered, but are not considerably mineralized. Details of the ore mineralogy and geochemistry of the deposit can be found in Morales Ruano et al. (2000) and Carrillo Rosúa et al. (2002 and 2003a). This work focuses on the composition of carbonate and phosphate as an aid to discriminate between hydrothermal alteration and regional low-temperature alteration.

Furthermore, dolomite and minor siderite are commonly formed during the later stages of mineralization. In associated hydrothermally altered volcanic rocks, dolomite very commonly replaces hornblende and plagioclase phenocrysts. Chemical analyses of both dolomites show slightly higher Fe and Mn in the orebody (averaging 1 at.% Fe and 0.5 at.% Mn) than in the adjacent alteration zones (averaging 0.51 at.% Fe and 0.44 at.% Mn). Idiomorphic crystals from the orebody show a chemical trend towards higher Fe and Mn content (Fig. 1). The isotopic differences between dolomite samples from the orebody ($\delta^{13}\text{C}$: -5.3 - -2.8‰) and hydrothermal alteration zones ($\delta^{13}\text{C}$: -8.5 - -4.5‰) are also important. This isotopic signature suggests that the carbon in the orebody has less of a magmatic input than that in the hydrothermal alteration zones.

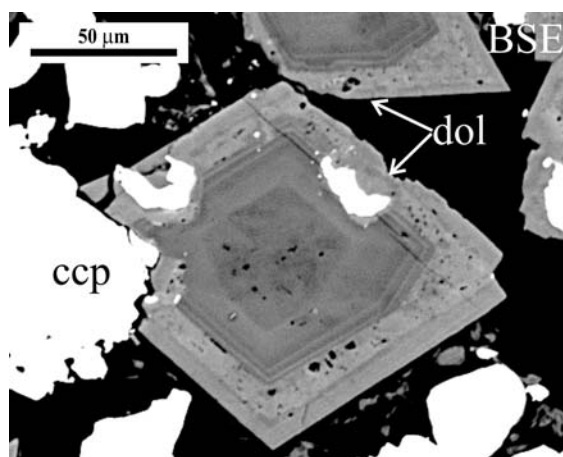


Figure 1: Back-scattered electron image of a dolomite crystal (dol) from the Palai-Islica orebody (SE Spain) with a marked zonation, the outer zones being richer (paler) in Mn-Fe than the inner (ccp: chalcopyrite).

It is noteworthy that the volcanic rocks of Cabo de Gata have undergone a regional alteration, unrelated to the mineral deposits. Mn- and Fe-poor calcite dominates this form of alteration. The calcite commonly replaces Ca-bearing phenocrysts or forms barren veins, and is characterized by $\delta^{13}\text{C}$ values of around -12‰. This isotopic signature suggests that the source region had an input of organic carbon. In the areas of propylitic alteration close to the Palai-Islica deposit, calcite is replaced by dolomite (Fig. 2), marking that the influence zone of hydrothermal ore-bearing fluids overprint the regional alteration.

At the Palai-Islica deposit, apatite is a minor phase that is disseminated in quartz veins containing sulfides. In hydrothermal alteration

zones, apatite is found in a mixture of chlorite \pm white mica \pm quartz \pm dolomite replacing volcanic hornblende phenocrysts. Apatite is also found in hornblende phenocrysts in andesitic to dacitic host-rocks that are not altered. The apatite in these rocks and sometimes in the orebody appears “clean” and is not porous. The apatite in the volcanic rocks may have crystallized directly from the magma, and in the orebody from hydrothermal fluids. In contrast, apatite found in hydrothermal alteration zones and sometimes in the orebody host xenotime inclusions, is always porous and is corroded. These crystals thought to be inherited from the volcanic rocks. These different styles of apatite also have distinctive compositions. Apatite in the volcanic rocks consists of 50–57% chlorapatite 24–36% hydroxylapatite, and 12–21% fluorapatite. Apatite in the orebody is mostly fluorapatite (between 95 and 100%) whereas apatite from the hydrothermally altered host-rock consists of 19–97% fluorapatite, 2–50% chloroapatite, and 0–44% hydroxylapatite. These observations indicate that the composition of apatite in hydrothermally altered zones is intermediate between those in volcanic rocks and the orebody. The textural and chemical features of the apatite from Palai-Islica suggest that apatite is hydrothermal altered from high Cl, OH-rich apatites in volcanic rocks to F-rich apatites (in which evolution constituted a “trap” of REE elements for the hydrothermal fluids). Enrichment in F is interpreted to relate to the intensity of hydrothermal processes and could thus provide a useful indicator of alteration and mineralization (Carrillo Rosúa and Morales Ruano, submitted).

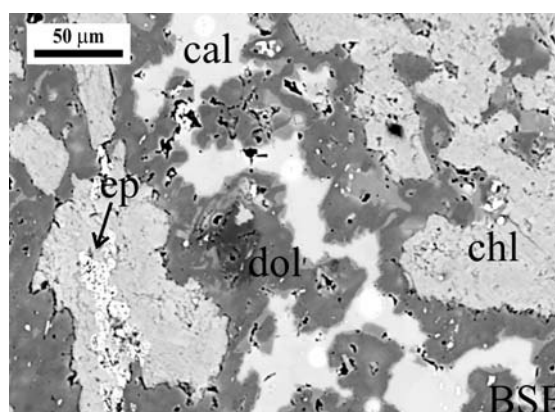


Figure 2: Back-scattered electron image of an intergrowth of dolomite crystal (dol) with chlorite (chl) and epidote (ep) in which dolomite is replacing calcite (cal) from a propylitically altered volcanic rock in the vicinity of Palai-Islica deposit.

3 La Serena: Mn, Cu(-Ag) stratabound, and Ba-Ag vein deposits, central Chile

A series of Mn, Cu(-Ag) stratabound and Ba-Ag vein deposits are hosted by late Cretaceous volcanic rocks located in the La Serena region of Chile (29°30' to 30°00' S). The deposits are in thick, volcanic, plagioclase-rich, porphyritic, high-K calc-alkaline to shoshonitic basaltic andesite and andesite rock, which have undergone low to very low-grade regional metamorphism (e.g. Aguirre and Egert 1965; Oyarzum et al. 1998; Morata and Aguirre 2003).

Manganese deposits in the La Serena region have ore mineralogy of braunite, piemontite, pyrolusite and hematite with calcite, barite, albite and white mica as gangue minerals (Carrillo Rosúa et al. 2003b). Calcite is the most abundant gangue mineral, which has concentrations of ≤ 0.94 at.% Mn (average of 0.24 at.%), ≤ 0.34 at.% Mg (average of 0.07 at.%) and ≤ 0.12 at.% Fe (average of 0.02 at.%). It is noteworthy that high-grade Mn areas contain calcite with low Mn (and Fe) content and high Mg content, whereas low-grade Mn areas are sulfide bearing and contain calcite with the opposite composition (Fig. 4).

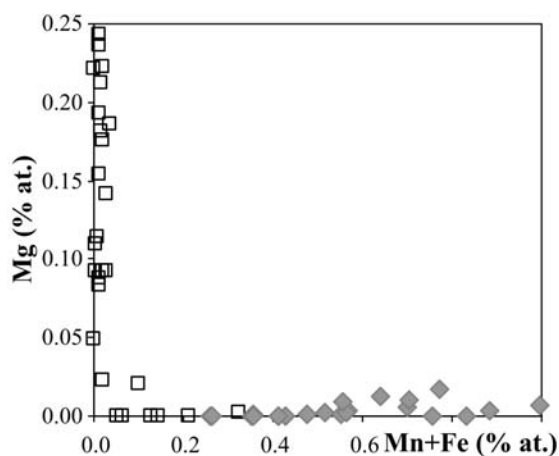


Figure 4: Binary diagram of Mn+Fe vs. Mg of calcites from Mn stratabound deposit (La Serena region, Central Chile) in \square Mn-rich areas non sulfide-bearing and in \blacklozenge Mn-poor areas sulfide-bearing.

Detailed studies of the Cu(-Ag) stratabound and Ba-Ag vein deposits show that the ore contains abundant and zoned calcite (Carrillo Rosúa et al. 2003b). Samples of calcite from Cu(-Ag) deposits have a high Mn content (0.04–1.63 at.%, average of 0.65 at.%), a significant Fe content (0.00 – 0.45 at.%, average of 0.12 at.%), and very low Mg

content Mg (0.00 – 0.05 at.%, average of 0.01 at.%), whereas calcite from Ba-Ag deposits only contains significant Mn (0.03 - 0.27 at.%, average of 0.12 at.%). It can be observed that Mn and Fe are proportional. Both types of deposits contain calcite with a complex zonation, suggesting that Mn(\pm Fe) calcite contain decreases towards the final stage of calcite precipitation (Fig. 5).

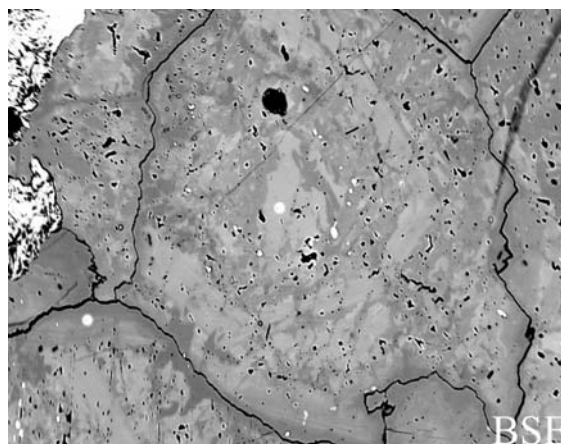


Figure 5: Back-scattered electron image of calcite from Ba-Ag veins of the La Serena region (Central Chile) with Mn-rich calcite (paler) replaced by Mn-poor calcite.

4 Melipilla: Cu (and Cu-Ag) hydrothermal vein and stratabound deposits, central Chile

The Cu (and Cu-Ag) vein and stratabound mineralization of the Melipilla area in the Coastal Range of Chile (33°45'S) is hosted by limestone and Early Cretaceous volcanic rocks similar to those in the La Serena area. Two kinds of mineralization are present in the area. The early type has a metamorphic paragenesis, and the later is epithermal containing significant quantities of calcite. The hydrothermally altered volcanic rocks also host non-mineralized or very low ore-grade calcite veins (Carrillo Rosúa et al. 2004). Calcite from the epithermal orebody usually has a high content of 0.01–1.92 at.% Mn (average of 0.44 at.%), and a low content of 0.03–0.63 at.% Mg (average of 0.14 at.%) and 0.00–0.16 at.% Fe (average 0.07 at.%). This is in contrast to calcite from carbonate clasts in the host-rock to the epithermal mineralization (Mn: 0.03 – 0.28 at.%; Mg: 0.06 – 0.20 at.%; Fe: 0.01 – 0.11 at.%), from the metamorphic mineralization (Mn: 0.00 – 0.01 at.%; Mg: 0.14 – 0.16 at.%; Fe: 0.00 – 0.01 at.%) and from the non-mineralized veins in hydrothermal alteration zones (Mn: 0.04 – 0.07 at.%; Mg: 0.00

