

1 **A new empirical model to estimate hourly diffuse photosynthetic photon flux density**

2

3

4 **I. Foyo-Moreno, (1,2), I. Alados (3,2) and L. Alados-Arboledas (1,2)**

5 **(2) Andalusian Institute for Earth System Research (IISTA-CEAMA), 18006**

6 **Granada, Spain**

7 **(3) Dpto de Física Aplicada II, Universidad de Málaga, Málaga, Spain.**

8

9 **Corresponding author:**

10 **I. Foyo-Moreno**

11 **Departamento de Física Aplicada**

12 **Facultad de Ciencias**

13 **Universidad de Granada**

14 **18071, Granada**

15 **Spain.**

16 **Phone: 34 58 240022**

17 **FAX: 34 58 243214**

18 **E-mail: ifoyo@ugr.es**

19

20

21

22

23

24

25 **ABSTRACT**

26

27 Knowledge of the photosynthetic photon flux density ( $Q_p$ ) is critical in different applications  
28 dealing with climate change, plant physiology, biomass production, and natural illumination  
29 in greenhouses. This is particularly true regarding its diffuse component ( $Q_{pd}$ ), which can  
30 enhance canopy light-use efficiency and thereby boost carbon uptake. Therefore, diffuse 31  
31 photosynthetic photon flux density is a key driving factor of ecosystem-productivity models.  
32 In this work, we propose a model to estimate this component, using a previous model to  
33 calculate  $Q_p$  and furthermore divide it into its components. We have used measurements in  
34 urban Granada (southern Spain), of global solar radiation ( $R_s$ ) to study relationships between  
35 the ratio  $Q_{pd}/R_s$  with different parameters accounting for solar position, water-vapour  
36 absorption and sky conditions. The model performance has been validated with experimental  
37 measurements from sites having varied climatic conditions. The model provides acceptable  
38 results, with the mean bias error and root mean square error varying between -0.3 and -8.8%  
39 and between 9.6 and 20.4%, respectively.

40

41

42 **1. INTRODUCTION**

43

44 The amount of photosynthetically active radiation (PAR), defined as the visible  
45 portion of global solar radiation ( $R_s$ ) used in photosynthesis to convert light energy into  
46 biomass (Udo and Aro, 1999; Jacovides et al., 2004; Tang et al., 2013), determines the  
47 exchanges of energy and water between the land surface and the atmosphere. Among the  
48 standard environmental statistics needed to evaluate plant photosynthesis are PAR data

49 (Akitsu et al., 2017). This radiation covering both photon and energy terms lies between 400  
50 and 700 nm. Thus, the photosynthetic photon flux density,  $Q_p$ , is defined as the photon flux  
51 density ( $1 \mu\text{mol photons m}^{-2} \text{s}^{-1} = 6.022 \times 10^{17} \text{ photons m}^{-2} \text{s}^{-1}$ ). This is the number of photons  
52 in the 400-700 nm waveband incident per unit time on a unit surface.

54 The amount and spectral quality of the radiation reaching the surface is altered by aerosols  
55 and clouds by absorption and scattering processes, varying the proportion of the components  
56 of radiation (direct and diffuse). Plants use both the direct and diffuse photosynthetic photon  
57 flux density ( $Q_{pb}$  and  $Q_{pd}$ ), but it is known that these components differ in the way they  
58 transfer energy through plant canopies and affect the summation of nonlinear processes such  
59 as photosynthesis differently than would occur at the leaf scale (Gu et al., 2002; Misson et  
60 al., 2005; Min, 2005; Jacovides et al., 2007; 2010). In fact, the efficiency with which  
61 incoming PAR is intercepted by a canopy, depends on its efficiency in intercepting direct and  
62 diffuse incoming radiation and on the proportions of  $Q_{pd}$  and  $Q_{pb}$  63 (Alados et al., 2002).

64 The analysis of the effects of  $Q_{pd}$  on ecosystem productivity has become one of the main  
65 goals in terrestrial carbon-cycle research (Gu *et al.*, 1999, 2002, 2003; Mercado *et al.*,  
66 2009; Sun and Zhou, 2010; Zhang *et al.*, 2010; He *et al.*, 2011; Zhang *et al.*, 2011), and thus  
67 the treatment of  $Q_{pd}$  in ecosystem models is needed (Gu *et al.*, 2003; Mercado *et al.*, 2009;  
68 Kanniah *et al.*, 2012). Hence, quantifying  $Q_{pd}$  and understanding its spatiotemporal variations  
69 are critical for estimating its impact on the carbon cycle of terrestrial ecosystems. Currently,  
70 many sites measure  $Q_p$ , although a worldwide network for routinely measuring  $Q_p$  has not yet  
71 been established, despite the biological importance of this radiometric quantity. Furthermore,  
72 few observation sites measure  $Q_{pd}$ , raising the need for models to estimate these data.  $Q_{pd}$  is  
73 usually estimated by multiplying  $Q_p$  by the diffuse fraction of global radiation ( $k_d = R_d/R_s$ ),  
74 where  $R_d$  is diffuse irradiance, but this is only a rough estimate because  $k_d$  is not equivalent to

75 the diffuse fraction in the visible range ( $k_{dQp}$ ), which is significantly greater than  $k_d$  under  
76 clear skies, although almost equivalent under cloudy skies (Spitters *et al.*, 1986).  
77 In the present work, we have evaluated a model to estimate  $Q_{pd}$  using a previous  
78 model to estimate  $Q_p$  (Foyo-Moreno *et al.*, 2017), considering the separation of  $R_s$  into its  
79 components: direct and diffuse. Prior to evaluating model performance, we have analysed  
80 dependences of the  $Q_{pd}/R_d$  ratio and  $Q_{pd}$  on different parameters. These parameters describe  
81 the solar position using the cosine of the solar zenith angle ( $\theta$ ) and indices such as sky  
82 clearness ( $\varepsilon = (R_d + R_b)/R_d$ ), sky brightness ( $\Delta = R_d/R_{so}\cos\theta$ ), and the clearness index ( $k_t =$   
83  $R_s/R_{so}$ ), where  $R_d$  is diffuse irradiance,  $R_b$  is direct normal irradiance and  $R_{so}$  is the  
84 extraterrestrial global solar irradiance (Pérez *et al.*, 1990; Alados *et al.*, 1996). For example,  
85 Jacovides *et al.* (2007; 2010) proposed an empirical model relating  $k_{dQp}$  to the clearness index  
86  $Q_p$  ( $k_{tQp} = Q_p/Q_{po}$ ), where  $Q_{po}$  is the extraterrestrial photosynthetic photon flux density. To  
87 consider the effects of water absorption in this spectral range, we also analysed dependence  
88 on meteorological parameters such as water-vapour pressure ( $e$ ) or dew-point temperature  
89 ( $T_d$ ). The model proposed in this work uses as input data the solar position as the cosine  
90 function of the solar zenith angle ( $\cos\theta$ ) and diffuse irradiance ( $R_d$ ). The model was evaluated  
91 with a set of independent data at various sites, two in Spain and another in Italy, with  
92 different climatological characteristics.

93

## 94 **2. EXPERIMENTAL SITE AND MEASUREMENTS**

95

96 In this work, we used data from three sites (Granada and Almería in Spain, and Renon  
97 in Italy). The data from Granada were used to study the main dependence between the  
98 variables of interest for this study, and all three stations were used to validate the proposed

99 model, since this proposal is based on a previous model to estimate  $Q_p$  (Foyo-Moreno et al.,  
100 2017), with data compiled for Granada during other years vs. those used in the present work.  
101 Ground-based data for two years at 1-min intervals were acquired at a station on the  
102 outskirts of Granada (37.18° N, 3.58° W, 660 m a.s.l.). From this data base, hourly values  
103 were generated for the entire two-year period to include a wide range of seasonal conditions  
104 and solar zenith angles. The photosynthetic active photon-flux density ( $Q_p$ ) was measured  
105 using LICOR model 190 SA quantum sensors (Lincoln, Nebraska, USA). Another quantum  
106 sensor has been equipped with a polar axis shadowband to measure the diffuse  
107 photosynthetic active photon flux density ( $Q_{pd}$ ). Global solar irradiance,  $R_s$ , was measured  
108 using a Kipp and Zonen model CM-11 radiometer (Delft, Netherlands), while another CM-11  
109 with a polar axis shadowband was used to measure diffuse solar irradiance ( $R_d$ ). The diffuse-  
110 irradiance measurements were corrected following the method proposed by Batlles et al.  
111 (1995). The quantum sensor has a relative error of less than 5% relative to the values  
112 measured, and global solar-irradiance measurements have an estimated experimental error  
113 of about 2-3%. The calibration constants of the instruments were periodically checked  
114 (Alados and Alados114 Arboledas, 1999).

115 To avoid problems associated with instrument deviations from the ideal cosine law,  
116 we limited our study to solar zenith angles of less than 85° (Alados et al., 1996). We  
117 evaluated the proposed model using data from different stations that had not been  
118 used in building the model, two in Spain (Granada and Almería), and another  
119 in Italy (Renon). Detailed information on the instrumentation and site characteristics  
120 can be found for Almería in Alados and Alados-Arboledas (1999). The radiometric  
121 sensors used at Almería are similar to those used at Granada. The Renon/Ritten site  
122 (Italy), is operated by the Forest Service and the Agency of the Environment

123 of the Autonomous Province of Bolzano (APB). In Italy,  $Q_p$  was measured  
124 by a BF2 sunshine sensor (Delta-T Devices, Burwell, Cambridge, United Kingdom).  
125 This device uses an array of silicon photodiodes and a shading pattern on the  
121 radiometer dome to determine  $Q_{pd}$ . The accuracy of these sensors is 15% (BF2, Delta-T  
122 Devices, 2005).

127 Table 1 presents some climatic data for the stations used. Granada is an inland location  
128 in south-eastern Spain, a non-industrialized medium-sized city situated in a natural basin  
129 surrounded by mountains with elevations between 1000 and 3500 m a.s.l.. Near continental  
130 conditions prevailing at this site are responsible for large seasonal temperature differences,  
131 providing cold winters and hot summers. Most rainfall occurs during winter and spring.  
132 Almería, located on the Mediterranean coast in south-eastern Spain, has frequent cloudless  
133 days and high humidity. Renon is situated at 1735m a.s.l. in the Italian Alps in Bolzano (Alto  
134 Adige, Italy). This site is influenced by an windy and humid alpine climate. Thus, the three  
135 sites present contrasting climatic characteristics and altitudes. The yearly precipitation  
136 markedly fluctuates.

137

### 138 **3. RESULTS AND DISCUSSION**

139

140

#### 141 **3.1. ANALYSIS OF $Q_{pd}/R_d$ RATIO**

142

143 As in our earlier work (Foyo-Moreno et al., 2017), before the direct study of the  
144 variable of interest ( $Q_{pd}$ ), we analysed the ratio between  $Q_{pd}$  and  $R_d$  ( $Q_{pd}/R_d$ ) with different  
145 parameters. Other works (Alados et al. 1996, Yu et al. 2015, Yu and Go, 2016) showed  
146 seasonal and daily variations of  $Q_p/R_s$  and also variations of  $Q_{pd}/R_d$  (Alados and Alados-

147 Arboledas, 1999). Since the most important factor influencing the solar-radiation levels  
148 reaching the Earth's surface is solar position, Figure 1 shows this dependence of this ratio on  
149  $\cos\theta$ . We noted that this ratio varied between 1.86 and 2.48  $\mu\text{mol J}^{-1}$  with a mean value of  
150  $2.19 \pm 0.13$  ( $\mu\text{mol J}^{-1}$ ). Jacovides et al. (2007) found a mean value in Athens (Greece) of  $2.43$   
151  $\pm 0.26$  ( $\mu\text{mol J}^{-1}$ ). For all hourly values together, we found no dependence. However, for the  
152 mean values of this ratio in different categories of  $\cos\theta$ , we detected a slight dependence with  
153 high dispersion. The correlation coefficient ( $R^2$ ) was 0.54 while the intercept and slope were  
154  $2.14 \pm 0.02$  ( $\mu\text{mol J}^{-1}$ ) and  $0.08 \pm 0.02$  ( $\mu\text{mol J}^{-1}$ ), respectively. The positive dependence with  
155  $\cos\theta$  implies a reduction in the relative contribution of photosynthetically active photon flux  
156 density over the whole solar spectrum when the optical air mass increases. In Almería (Spain)  
157 a similar dependence has been reported (Alados et al., 1999). In an initial approximation, we  
158 could assume that this ratio was constant and we calculated  $Q_{pd}$  from  $R_d$ .

159 The second dependence analysed was on  $R_s$  (Figure 2) and the third on  $R_d$  (Figure 3).  
160 There was no dependence on  $R_s$  for all hourly values but a positive correlation was found for  
161 the mean values considering the categories with the  $R_s$  values. However, a clearer  
162 dependence for all hourly values was detected, although with great dispersion, as a function  
163 of  $R_d$ , with larger values for low values of  $R_d$ . Thus, with decreasing  $R_d$  (clear skies),  
164 the diffuse component of  $Q_p$  was larger than the diffuse component for  $R_d$ .

165 In our search for a more direct relation with sky condition, Figure 4 shows the  
166 dependence of this ratio on the clearness index ( $k_t$ ). This parameter is defined as the ratio of  
167 the global irradiance ( $R_s$ ) to the extraterrestrial global irradiance ( $R_{so}$ ), both on a horizontal  
168 surface:

169

170

$$k_t = \frac{R_s}{R_{so}} \quad (1)$$

171

172

with  $R_{so} = E_o I_{sc} \cos\theta$ .  $E_o$  is the eccentricity correction factor and the value used of the solar

173

constant ( $I_{sc}$ ) is  $1367 \text{ Wm}^{-2}$  (Iqbal, 1983). Details of the calculation of  $\cos\theta$  can be found by

174

Iqbal (1983). The parameter  $k_t$  characterizes the sky condition including the attenuation

175

effects of the most dominant factors controlling solar radiation, such as clouds and aerosols.

176

Previous studies used similar parameters such as the sky clearness and the skylight brightness

177

(Pérez et al., 1990) to characterize the sky condition (Alados et al., 1996). Another parameter

178

to characterize sky condition is diffuse fraction ( $k_d$ ), defined as the ratio between diffuse

179

global irradiance ( $R_d$ ) and global irradiance ( $R_s$ ):

178

$$k_d = \frac{R_d}{R_s} \quad (2)$$

179

Knowledge of  $k_d$  can be useful to get an idea of atmospheric load indirectly, where a

180

low  $k_d$  indicates a clear sky and more pristine atmosphere, while high  $k_d$  values denote high

181

aerosol loads (Singh et al., 2013). In fact, the ratio between  $R_d$  and  $R_{bn}$  can be used to

182

estimate aerosol optical depth (Foyo-Moreno et al., 2014).

183

Figures 4 and 5 show the dependence of  $Q_{pd}/R_d$  on these parameters. There is a slight

184

dependence on  $k_t$  for all hourly data and for mean values in categories of  $k_t$  values, higher

185

values of the  $Q_{pd}/R_d$  ratio are detected for low and high  $k_t$  values. These high values of  $Q_{pd}/R_d$

186

for high values of  $k_t$  are consistent with Jacovides et al. (2007), who reported a weak

187

correlation for daily values, which is well fit by an exponential equation with a determination

188

coefficient of 0.276. They found hourly ratio values varying from  $2.56 \pm 0.27$  to  $2.42 \pm 0.24$

189

( $\mu\text{mol J}^{-1}$ ), and ratios under clear skies 5.5% higher than that for overcast skies. These results

190 agree with Alados and Alados-Arboledas (1999) and Min (2005). In any case, high  $k_t$  values  
191 imply low  $k_d$  values, so that this ratio had high values for low values of  $k_d$  (clear skies)  
192 although the dispersion of data was high (Figure 5), increasing the dispersion with  $k_d$ . Under  
193 clear skies, the regulating factors in the solar radiation modification are scattering processes,  
194 so that spectral investigations support the premise that, under these skies with aerosol loads,  
195 short wavelengths are preferentially scattered, thus increasing  $Q_{pd}$  more than  $R_d$  (Jacovides et  
196 al., 2000; Dye, 2004). However, the pattern observed for low values of  $k_t$  ( $<0.4$ ) is different  
197 (Figure 4), the ratio decreasing when  $k_t$  increases. This result could be associated with the  
198 mixed effects of aerosols and clouds, because these cases correspond to skies with high  
199 opacity.

200 From the above analysis, it is clear that this ratio varies with changing sky conditions.  
201 Clouds exhibit relatively constant extinction across the visible spectrum while aerosols  
202 commonly show extinction that significantly diminishes the greater the wavelength (Min,  
203 2005; Jacovides et al., 2007). This result can be attributed to the presence of clouds in skies  
204 with the high solar absorption in the infrared region (NIR). Clouds absorb NIR more strongly  
205 than  $Q_p$ , and hence the transmittance of  $Q_p$  through clouds is larger than NIR. In fact, clouds  
206 more markedly attenuate the total solar spectral range than they do shorter wavelengths such  
207 as the ultraviolet range (Foyo-Moreno et al., 2001, 2003).

208 As in an earlier work (Foyo-Moreno et al., 2017), here, to explicitly consider  
209 absorption by water vapour in the solar spectrum, we have shown in Figures 6 and 7 the  
210 dependence on meteorological parameters such as the dewpoint temperature ( $T_d$ ) – relevant  
211 due to its correlation with precipitable water (Reitan, 1963) – and the partial vapour pressure  
212 ( $e$ ).  $T_d$  and  $e$  have been calculated from the direct measurements of air temperature ( $T$ ) and  
213 the relative humidity ( $U$ ) at our station, with standard formulation. As before

214 (Foyo-Moreno et al., 2017), no dependence was detected.

215

### 216 **3.2. ANALYSIS OF PARAMETERS CHARACTERIZING SKY CONDITIONS**

217

218 Now, in order to find relationships between parameters considering sky conditions, it  
219 is possible to define indexes for the visible range equivalent to those for the shortwave  
220 broadband range. These parameters are  $k_{tQP} = Q_p/Q_{p0}$  (clearness index) and  $k_{dQP} = Q_{pd}/Q_p$   
221 (diffuse fraction), defined as in Eqs. 1 and 2, respectively but for the visible range.  $Q_{p0}$  can be  
222 derived from  $R_{so}$ , with a fraction of 0.5 suggested by most studies (Frouin and Pinker 1995;  
223 Hu et al. 2007, 2010; Wang et al. 2014), then  $Q_{p0}$  can be calculated by multiplying by the  
224 ratio 4.57 of the energy-flux density to photosynthetic photon flux density (Dye 2004).

225 Figure 8 shows the dependence between  $k_{dQP}$  and  $k_{tQP}$ . Jacovides et al. (2007) found a  
226 polynomial fitting, shown in Figure 8 as the dashed line for  $0.1 < k_{tQP} < 0.85$ . Our fitting is also  
227 a third-order polynomial but with different coefficients (solid line) with a determination  
228 coefficient of 0.88. For low values of  $k_{tQP}$  the diffuse component is similar to  $Q_p$  and for high  
229 values this proportion decreases to about 10% of  $Q_p$ . In the literature, few studies have  
230 focused on the empirical relationships between  $Q_p$  and  $Q_{pd}$  of the Liu and Jordan regression  
231 type (Jacovides et al., 2010). The applicability of several diffuse radiation empirical models  
232 has been analysed in Athens (Jacovides et al., 2010), showing mean bias error (MBE) values  
233 between -12.9 and 2.99 % and root mean square error (RMSE) values between 27.1 and  
234 35.6%, with a MBE of 0.11% and a RMSE of 26.8% for their fitting.

235 Figure 9 shows the dependence between the ratio of the diffuse fraction in the visible  
236 range ( $k_{dQP}$ ) and the diffuse fraction in the shortwave broadband range ( $k_d$ ) and clearness  
237 index ( $k_t$ ). We found a positive correlation, with the ratio increasing with clear skies. Spitters

238 et al. (1986) showed that the diffuse fraction for the visible range was 1.4 times the diffuse  
239 broadband fraction under clear sky conditions. It is evident that this ratio increases  
240 significantly when sky conditions change from cloudy to clear. In our case, the value found  
241 was  $1.17 \pm 0.07$  for clear skies. Jacovides et al. (2007) reported a value of  $1.38 \pm 0.32$ . The  
242 dispersion increases with higher  $k_t$  values. However, for low values of  $k_t$  ( $<0.4$ ; cloudy skies)  
243 the diffuse fraction in the visible range is of the same order as the diffuse fraction in the  
244 shortwave broadband range.

245

### 246 **3.3. ANALYSIS OF $Q_{pd}$ .**

247

248 In this section, we directly study the dependence of  $Q_{pd}$  on the same parameters  
249 considered in Section 3.1. First, we show that the solar zenith angle alone does not allow for  
250 a correct determination of  $Q_{pd}$  (Figure 10). While  $Q_{pd}$  increases with decreasing  $\theta$ , due to  
251 high dispersion it is possible to parameterise only the two *envelopes* of the data through a  
252 simple linear dependence on the cosine of  $\theta$ . This result is similar to earlier work with  $Q_p$   
253 (Foyo Moreno et al., 2017), and with the ultraviolet range (Foyo-Moreno et al., 1998, 2007).  
254 The upper envelope of the data corresponds to the maximum values that can be assumed to  
254 correspond to overcast skies, and the lower envelope can be attributed to values  
255 corresponding to clear skies for a given solar position.

257 We show the high positive correlation between  $Q_{pd}$  and  $R_d$  ( $R^2= 0.997$ ), with a slope of  
258  $2.125 \pm 0.002 \mu\text{mol J}^{-1}$  (Figure 11), close to the mean value found for the ratio  $Q_{pd}/R_d$  [ $2.19 \pm$   
259  $0.13 (\mu\text{mol J}^{-1})$ ]. The dependence on sky conditions using the parameters  $k_t$  and  $k_d$  is shown  
260 in Figures 12 and 13. The dispersion of data is high for all sky conditions.

261 Considering all of the above results, we found it more appropriate to model  $Q_{pd}$  rather

262 than  $Q_{pd}/R_d$ . Therefore, taking into account our previous model (Foyo-Moreno et al., 2017),  
 263 the expression of  $Q_p$  is:

$$264 \quad Q_p = a k_t \cos \theta \quad (3)$$

265 We can separate  $Q_p$  into its two components: direct and diffuse. Also, taking into account that  
 266 global irradiance ( $R_s$ ) is the sum of the two components (direct and diffuse), we arrive at this  
 267 expression for  $Q_{pd}$ :

$$268 \quad Q_{pd} = a \cos \theta \frac{R_d}{R_{so}} \quad (4)$$

268 with  $a = (2681 \pm 2) \mu\text{mol m}^{-2}\text{s}^{-1}$ .

270

### 271 **3.4. MODEL PERFORMANCE**

272

273 We have evaluated the model (Eq. 4) at the three stations which were not used in its  
 274 development. Table 2 shows the results found after including the correlation coefficient  $R^2$  –  
 275 (the fraction of experimental data variance explained by the model), slope  $b$ , and the intercept  
 276  $a$  of the linear regression of the  $Q_{pd}$  measured vs. the estimated values. Table 2 also shows the  
 277 mean bias error (MBE) and root mean square error (RMSE), both as percentages of the mean  
 278 experimental values. The MBE and RMSE, are given by the following expressions:

$$279 \quad MBE = \frac{100}{M_{ave}} \frac{1}{N} \sum_{i=1}^N (E_i - M_i) \quad (5)$$

$$280 \quad RMSE = \frac{100}{M_{ave}} \left( \frac{1}{N} \sum_{i=1}^N (E_i - M_i)^2 \right)^{0.5} \quad (6)$$

281 where  $E_i$  is the estimated value ( $i$ th number),  $M_i$  is the measured value,  $M_{ave}$  is the average of

282 the measured values and  $N$  is the number of observations. These two statistics enable the  
283 detection both of the differences between experimental data and of the model estimates and  
284 the existence of systematic over- or underestimation trends, respectively.

285 Figure 13 shows  $Q_{pd}$  estimated via Eq. (4) at all localities analysed in this work. The  
286 model evaluation is excellent for all radiometric stations, the variance explained for the  
287 models is better than 95%, and the values for MBE and RMSE vary between -0.3 and 8.6 %  
288 and 9.6 and 20.4%, respectively. The model underestimates the values for all localities. In any  
289 case, both the slope and correlation coefficient of the linear regression between the measured  
290 and estimated values reveals the goodness of the model estimations. Wang et al. (2017), using  
291 different soft computing techniques, found MBE values between -16 and -18% and RMSE  
292 ranging between 22 and 51% for six stations from the AmeriFlux network.

293

### 294 **3. CONCLUSIONS**

295

296 Two years of measurements made in Granada, an urban site in south-eastern Spain,  
297 were used to study the diffuse photosynthetic photon-flux density ( $Q_{pd}$ ), with the aim of  
298 proposing a model for all sky conditions. The model was evaluated at three stations, two in  
299 Spain (Granada and Almería) and another in Italy (Renon), with different climatic  
300 characteristics. Firstly, we analysed the ratio of  $Q_{pd}$  to diffuse irradiance ( $R_d$ ) to continue with  
301  $Q_{pd}$ . The dependence of both ( $Q_{pd}/R_d$  and  $Q_{pd}$ ) on different parameters, which characterize  
302 solar position and sky conditions, were considered together with other meteorological factors  
303 such as vapour pressure and dewpoint temperature ( $T_d$ ). The main findings can be  
304 summarized as follows:

305

- 306 1. A mean value of  $2.19 \pm 0.13 \mu\text{mol J}^{-1}$  was found for  $Q_{pd}/R_d$  with values varying between  
307 1.86 and  $2.48 \mu\text{mol J}^{-1}$ .
- 308 2.  $Q_{pd}/R_d$  showed dependence on parameters that characterize sky conditions such as  
309 clearness index ( $k_t$ ) and diffuse fraction ( $k_d$ ).
- 310 3. A model was proposed to estimate  $Q_{pd}$ , taking into account a previous model to estimate  
311 the photosynthetic active radiation-flux density ( $Q_p$ ). This model uses the solar zenith  
312 Angle ( $\theta$ ) and  $R_d$  as input data.
- 313 4. The model provided acceptable results, with a mean bias error varying between -0.3% and  
314 8.8%; and the root mean square error (RMSE) varying between 9.6 and 20.4%.
- 315 5. This work suggests that this empirical model can be widely applied from available  
316 measurements in most radiometric stations, with better accuracy than other empirical  
models.

318 **ACKNOWLEDGEMENTS:** This work was supported by the Andalusia Regional  
319 Government project P12-RNM-2409, by the Spanish Ministry of Economy and  
320 Competitiveness projects CGL2013-45410-R and CGL2016-81092-R, and by the European  
321 Union's Horizon 2020 research and innovation programme project ACTRIS-2 (grant  
322 agreement No 654109). The authors would like to thank the Forestry and Forest Products  
323 Research Institute FluxNet Database, FFPRI <http://www2.ffpri.affrc.go.jp/labs/flux/>.

324

## 325 REFERENCES

326 Akitsu, T., Nasahara, K.N., Hirose, Y., Ijima, O., Kume, A., 2017. Quantum sensors for  
327 accurate and stable long-term photosynthetically active radiation observations. Agric. For.  
328 Meteorol. 237-238, 171-183.

329 Alados, I., Foyo-Moreno, I., Alados-Arboledas, L., 1996. Photosynthetically active radiation:  
330 measurements and modeling. *Agric. For. Meteorol.* 78, 121-131.

331 Alados, I., Alados-Arboledas L., 1999. Direct and diffuse photosynthetically active radiation:  
332 measurements and modelling. *Agric. For. Meteorol.* 93, 27-38

333 Alados, I., Olmo, F.J., Foyo-Moreno, I., Alados-Arboledas, L., 2000. Estimation of  
334 photosynthetically active radiation under cloudy conditions. *Agric. For. Meteorol.* 102, 39-  
335 50.

336 Alados, I., Foyo-Moreno, I., Olmo, F.J., Alados-Arboledas, L., 2002. Improved estimation of  
337 diffuse photosynthetically active radiation using two spectral models. *Agric. For. Meteorol.*  
338 101, 1-12.

339 Batlles, F., Olmo, F., Alados-Arboledas, L., 1995. On shadowband correction methods for  
340 diffuse irradiance measurements. *Solar Energy* 54, 105-114

341 Dye, D., 2004. Spectral composition and quanta-to-energy ratio of diffuse photosynthetically  
342 active radiation under diverse cloud conditions. *J. Geophys. Res.* 109, D10203  
343 10.1029/2003JD004251

344 Foyo-Moreno, I., Vida, J., Alados-Arboledas, L., 1998. A simple all weather model to  
345 estimate ultraviolet solar radiation (290-385nm). *J. Apl. Metereol.*, 38, 1020–1026.

346 Foyo-Moreno, I., Alados, I., Olmo, F.J., Vida, J., Alados-Arboledas, L., 2001. On the  
use of a cloud modification factor for solar UV (290-385 nm) spectral range. *Theor. Appl.*  
348 *Climatol.*, 68, 41-50.

349 Foyo-Moreno, I., Alados, I., Olmo, F.J., Alados-Arboledas, L., 2003. The influence of  
350 cloudiness on UV global irradiance (295-385 nm). *Agric. For. Meteorol.* 120, 101-111.

351 Foyo-Moreno, I., Alados, I., Alados-Arboledas, L., 2007. Adaptation of an empirical model  
352 for erythemal ultraviolet irradiance. *Ann. Geophys.*, 25, 1–10.

353 Foyo-Moreno, I., Alados, I., Antón, M., Fernández-Gávez, J., Cazorla, A., Alados-Arboledas,  
354 L., 2014. Estimating aerosol characteristics from solar irradiance measurements at an urban  
355 location in southeastern Spain. *J. Geophys. Res.* doi: 10.1002/2013JD020599104

356 Foyo-Moreno, I., Alados, I., Alados-Arboledas, L., 2017. A new conventional regression  
357 model to estimate hourly photosynthetic photon flux density under all sky conditions. *Int.*  
358 *J. Climatol.* doi: 10.1002/joc.5063

359 Frouin, R., Pinker, R.T., 1995. Estimating photosynthetically active radiation (PAR) at the  
360 earth's surface from satellite observations. *Remote Sens Environ* 51:98–107

361 Gu, L., Fuentes, J.D., Shugart, H.H, Staebler, R. M., Black, T.A., 1999. Responses of net  
362 ecosystem exchanges of carbon dioxide to changes in cloudiness: Results from two North  
363 American deciduous forests. *J. Geophys. Res.*, 104, 31 421–31 434

364 Gu, L., Baldocchi, D., Verma, S.B., Black, T.A., Vesala, T., Falge, E.M., et al. 2002.  
365 Advantages of diffuse radiation for terrestrial ecosystem productivity. *J Geophys Res*,  
366 107(D6). doi:10.1029/2001JD001242

367 Gu, L.H., Baldocchi, D.D., Wofsy, S.C., et al., 2003. Response of a deciduous forest  
368 to the Mount Pinatubo eruption: Enhanced photosynthesis. *Science*, 299(5615): 2035–2038.

369 He, X.Z., Zhou, T., Jia, G.S. et al., 2011. Modeled effects of changes in the amount and  
370 diffuse fraction of PAR on forest GPP. *Journal of Natural Resources*, 26(4): 619–634

371 Hu, B., Wang, Y., Liu, G., 2007. Spatiotemporal characteristics of photosynthetically active  
372 radiation in China. *J. Geophys. Res.*, 112, doi: 10.1029/2006JD007965.

373 Hu, B., Wang, Y., Liu, G., 2010. Variation characteristics of ultraviolet radiation derived  
374 from measurement and reconstruction in Beijing, China, *Tellus B* 62, 100–108.

375 Iqbal, M. 1983. *An Introduction to Solar Radiation*. Academic Press: London.

376 Jacovides, C.P., Tymvios, F.S., Papaioannou, G., Asimakopoulos, D.N., Theofilou, K.M.,

377 2004. Ratio of PAR to broadband solar radiation measured in Cyprus, *Agric. For. Meteor.*,  
378 121, 135-140.

379 Jacovides, C.P., Tymvios, F.S., Asimakopoulos, V.D., Kaltsounides, N.A., 2007. The  
380 dependence of global and diffuse PAR radiation components on sky conditions at Athens,  
381 Greece, *Agric. For. Meteor.*, 143, 277-287.

382 Jacovides, C.P., Boland, J., Asimakopoulos, D.N., Kaltsounides, N.A., 2010. Comparing  
383 diffuse radiation models with one predictor for partitioning incident PAR radiation into its  
384 diffuse component in the eastern Mediterranean basin, *Renew. Energy*, 35, 1820-1827.

385 Kanniah, K.D., Beringer, J., North, P., Hutley, L., 2012. Control of atmospheric particles on  
386 diffuse radiation and terrestrial plant productivity A review. *Progr Phys Geogr* , 36(2), 209-  
387 237.

388 Mercado, L.M., Bellouin, N., Sitch, S. et al., 2009. Impact of changes in diffuse radiation on  
389 the global land carbon sink. *Nature*, 458(7241): 1014–1017.

390 Min, Q., 2005. Impacts of aerosols and clouds on forest-atmosphere carbon exchange. *J*  
391 *Geophys Res*110:D066203. doi:10.1029/2004JD004858.

392 Misson, L., Lunden, M., McKay, M., Goldstein, A.H., 2005. Atmospheric aerosol light  
393 scattering and surface wetness influences the diurnal pattern of net ecosystem exchange in  
394 a semi-arid ponderosa pine plantation. *Agric For Meteorol* 129:60–83.

395 Perez, R., Ineichen, P., Seals, R., Michalsky, J.J., Stewart, R., 1990. Modelling daylight  
396 availability and irradiance components from direct and global irradiance. *Solar Energy* 44,  
397 271-289.

398 Reitan, C.H., 1963. Surface dewpoint and water vapour aloft. *J. Appl. Meteorol.*, 2: 776-779.

399 Singh, J., Bhattacharya, B.K., Kumar, M., Mallick, K.. 2013. Modelling monthly diffuse  
400 solar radiation fraction and its validity over the Indian sub-tropics. *Int. J. Climatol.*, 33, 77-

401 86.

400 Spitters, C.J.T., Toussaint, H.A.J.M., Goudriaan, J. 1986. Separating the diffuse and direct  
401 component of global radiation and its implications for modelling canopy photosynthesis.  
402 *Agric For Meteorol*, 38:217–29

403 Sun, J.S., Zhou, G.S., 2010. Review of advances in measurements and effects of diffuse  
404 radiation on terrestrial ecosystem productivity. *Chinese Journal of Plant Ecology*, 34(4):  
405 452–461. (in Chinese).

407 Tang, W., Qin, J., Yang, K., Niu, X., Zhang, X., 2013. Reconstruction of daily  
408 photosynthetically active radiation and its trend over China. *J. Geophys. Res. Atm.*, 118,  
409 13292-13302.

410 Udo, S., Aro, T., 1999. Global PAR related to global solar radiation for central Nigeria.  
411 *Agric. For. Meteorol.* 97, 21–31, [http://dx.doi.org/10.1016/S0168-1923\(99\)55-6](http://dx.doi.org/10.1016/S0168-1923(99)55-6).

412 Wang, L., Gong, W., Ma, Y., Hu, B., Zhang, M., 2014. Photosynthetically active radiation  
413 and its relationship with global solar radiation in Central China. *Int. J. Biometeorol.*, 58,  
414 1265-1277.

415 Wang, L., Hu, B., Kisi, O., Zounemat-Kermani, M., Gong, W., 2017. Prediction of diffuse  
416 photosynthetically active radiation using different soft computing techniques. *Q. J. R.  
417 Meteorol. Soc.*, doi: 10.1002/qj.3081.

418 Yu, X., Wu, Z., Jiang, W., Guo, X., 2015. Predicting daily photosynthetically active radiation  
419 from global solar radiation in the Contiguous United States. *Energy Convers. Manage.*, 89,  
420 71-82.

421 Yu, X., Guo, X., 2016. Hourly photosynthetically active radiation estimation in Midwestern  
422 United States from artificial neural networks and conventional regression models. *Int. J.  
423 Biometeorol.*, 60, 1247-1259.

424 Zhang, M., Yu, G.R., Zhang, L.M. et al., 2010. Impact of cloudiness on net ecosystem  
425 exchange of carbon dioxide in different types of forest ecosystems in China.  
426 *Biogeosciences*, 7(2): 711–722.

427 Zhang, M., Yu, G.R., Zhuang, J. et al., 2011. Effects of cloudiness change on net ecosystem  
428 exchange, light use efficiency, and water use efficiency in typical ecosystems of China.  
429 *Agric.For.Meteorol.*151(7),803–816.

431

Sites	T (°C)	TM (°C)	Tm (°C)	R (mm)	U (%)
<b>Granada, Spain</b> (37.16° N, 3.6° W, 650 m a.s.l.)	15.4	23.0	7.8	365	58
<b>Almería, Spain</b> (36.83° N, 2.41° W, 21 m a.s.l.)	19.1	23.4	14.7	200	65
<b>Renon, Italy</b> (46.42° N, 11.28° E, 1735 m a.s.l.)	4.1	18.0	5.2	1010	

432 Table 1: Description and climatic data of study sites. T (yearly average temperature), TM  
433 (yearly average of the maximum temperature), Tm (yearly average of the minimum  
434 temperature), R (yearly average of precipitation), U (yearly average of relative humidity).

435

436

437

438

439

440

441

442

443

444

445

446

447

448

449

450

451

452

453

454

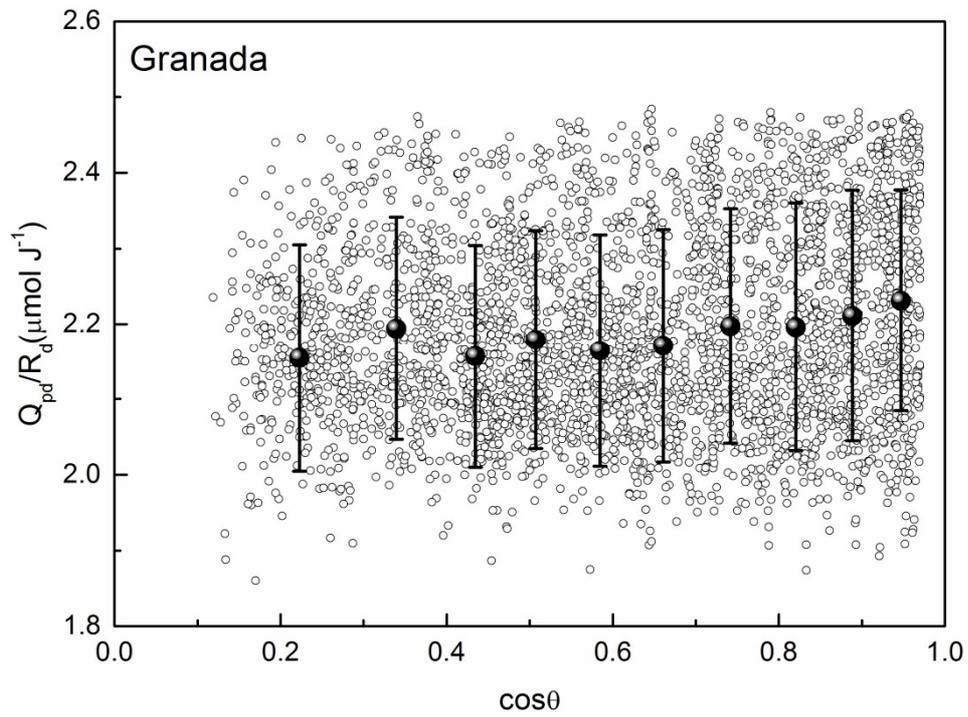


<b>Sites</b>	<b>Years</b>	<b>N</b>	<b>kt</b>	<b>k</b>	<b>Q<sub>pd,ave</sub> (μmol m<sup>-2</sup>s<sup>-1</sup>)</b>	<b>a (μmol m<sup>-2</sup>s<sup>-1</sup>)</b>	<b>b</b>	<b>R<sub>2</sub></b>	<b>MBE (%)</b>	<b>RMSE (%)</b>
<b>Granada</b>	94-95	3448	0.54±0.19	0.55±0.30	428	-30.8 ± 0.8	0.992 ± 0.002	0.99	-7.7	9.6
<b>Almeria</b>	93-95	5797	0.58±0.18	0.50±0.30	419	-19.5 ± 1.1	0.959 ± 0.002	0.96	-8.8	12.9
<b>Renon</b>	13-15	16980	0.45 ±0.27	0.59±0.35	296	-60.3 ±0.7	1.200±0.002	0.95	-0.3	20.4

457 Table 2: Statistical results of the new empirical model to estimate Q<sub>pd</sub> at each station. N (total number of observations). Q<sub>pd,ave</sub>

458 (average values of Q<sub>pd</sub>). Linear regression statistics: a (intercept), b (slope) and R<sup>2</sup> (correlation coefficient). MBE (Mean Bias Error).

459 RMSE (Root Mean Square Error).



460

461

462

463 Figure 1: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance ( $Q_{pd}/R_d$ ) vs.

464 the cosine of solar zenith angle ( $\cos \theta$ ). Small dots denote experimental data, black symbols

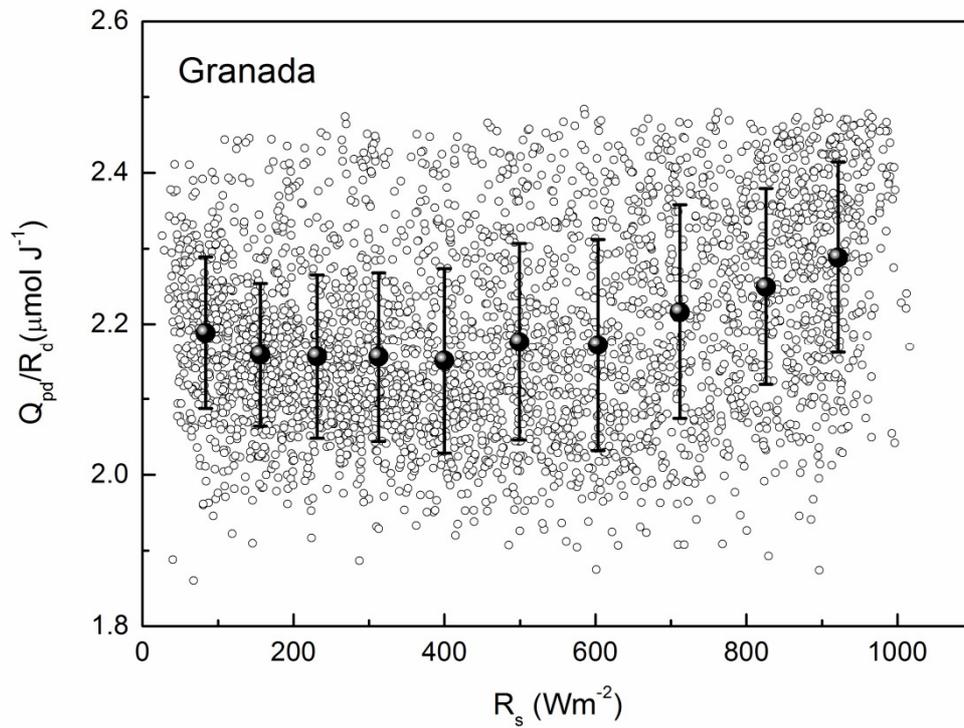
465 represent mean values, and bars denote standard deviations for each of the intervals in  $\cos \theta$ .

466

467

468

469



470

471

472

473 Figure 2: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance ( $Q_{pd}/R_d$ ) vs.  
 474 broadband solar irradiance ( $R_s$ ). Small dots denote experimental data, black symbols represent  
 475 mean values, and bars denote standard deviations for each of the intervals in  $R_s$ .

476

477

478

479

480

481

482

483

484

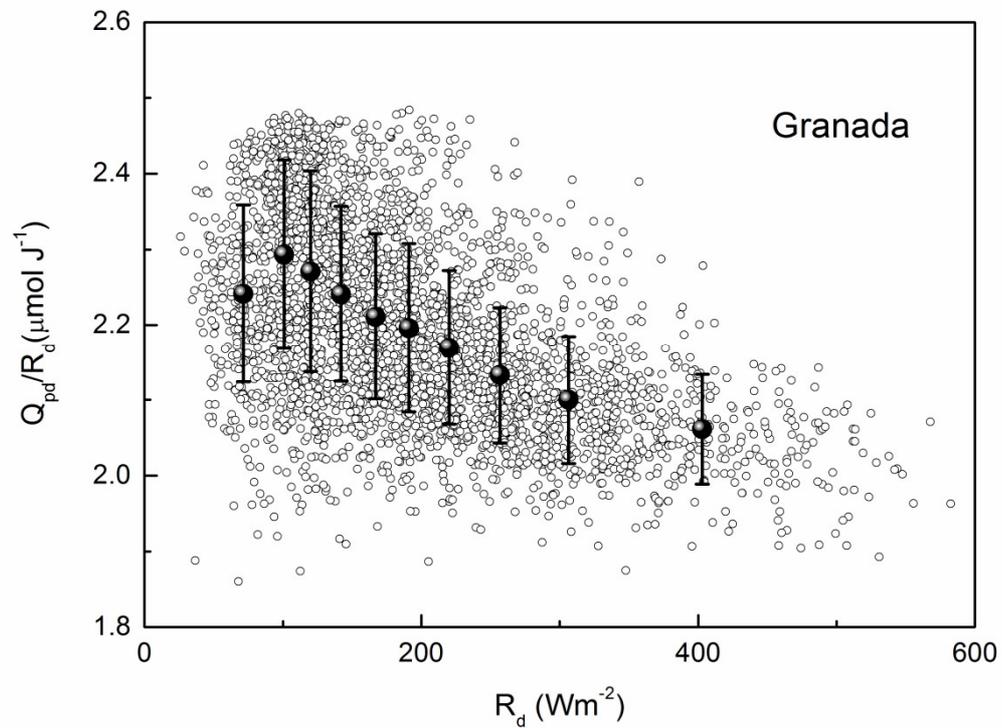
485

486

487

488

489



491

492 Figure 3: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance ( $Q_{pd}/R_d$ ) vs.493 diffuse irradiance ( $R_d$ ). Small dots denote experimental data, black symbols represent mean494 values, and bars denote standard deviations for each of the intervals in  $R_d$ .

495

496

497

498

499

500

501

502

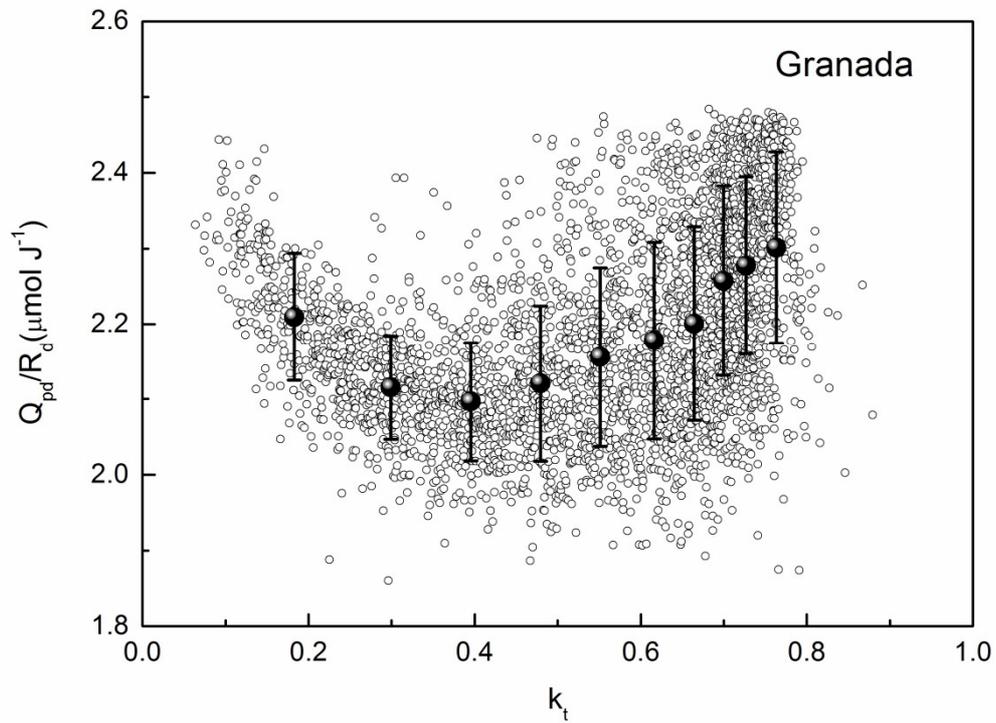
503

504

505

506

507



508

509

510

511 Figure 4: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance ( $Q_{pd}/R_d$ ) vs.

512 clearness index ( $k_t$ ). Small dots denote experimental data, black symbols represent mean

513 values, and bars denote standard deviations for each of the intervals in  $k_t$ .

514

515

516

517

518

519

520

521

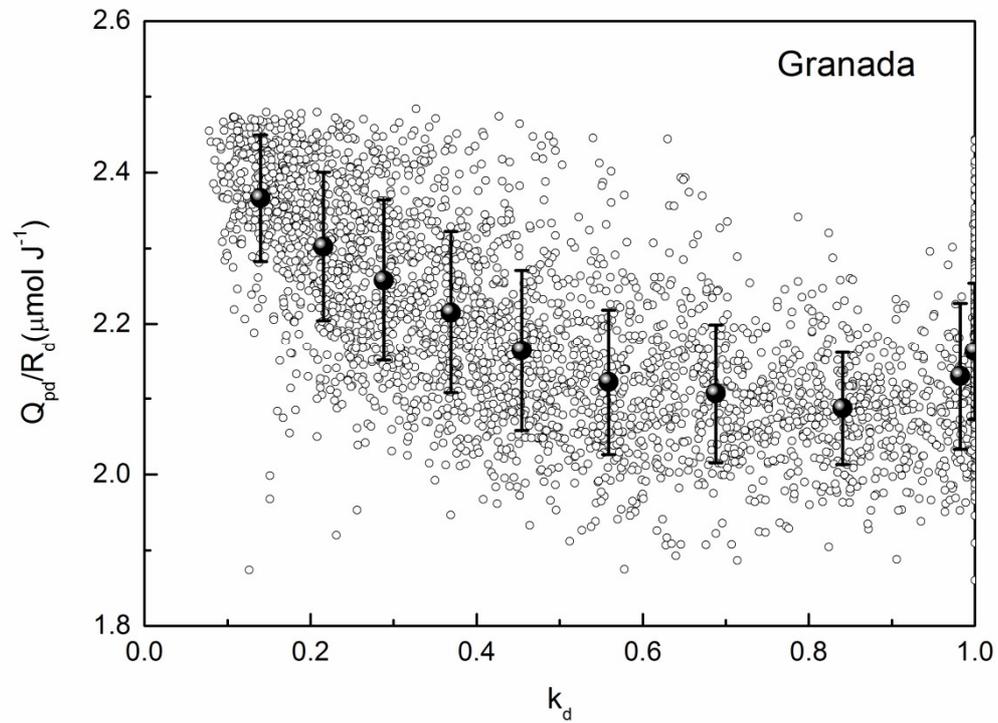
522

523

524

525

526



527

528

529 Figure 5: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance ( $Q_{pd}/R_d$ )

530 vs.  $k_d$ . Small dots denote experimental data, black symbols represent mean values and bars

531 denote standard deviations for each of the intervals in  $k_d$ .

532

533

534

535

536

537

538

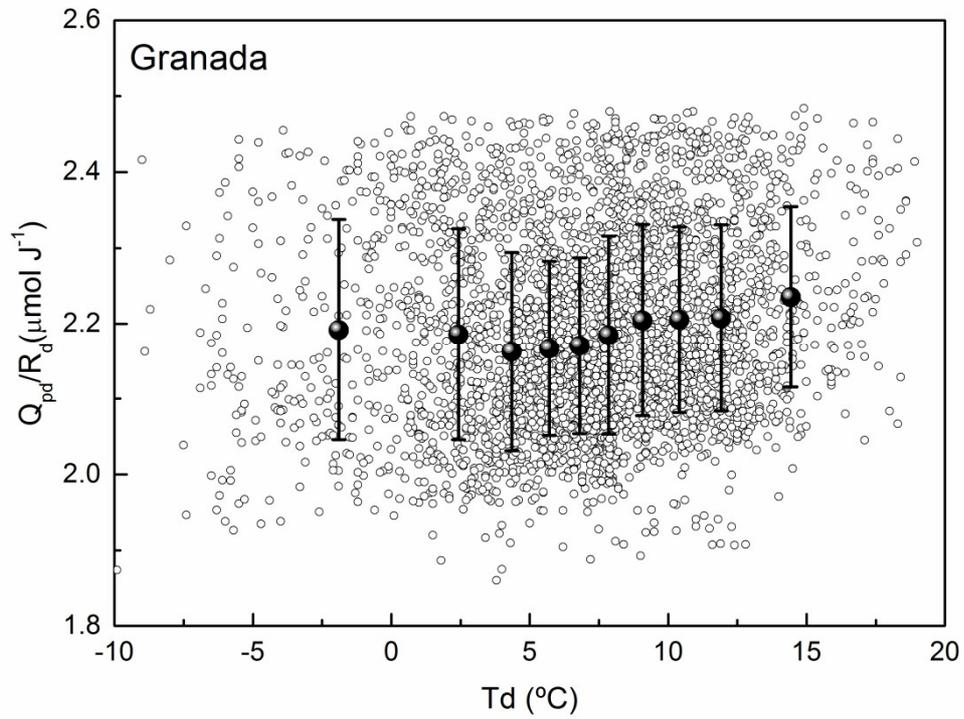
539

540

541

542

543



544

545

546 Figure 6: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance ( $Q_{pd}/R_d$ ) vs.  
547 the dewpoint temperature ( $T_d$ ). Small dots denote experimental data, black symbols represent  
548 mean values, and bars denote standard deviations for each of the intervals in  $T_d$ .

549

550

551

552

553

554

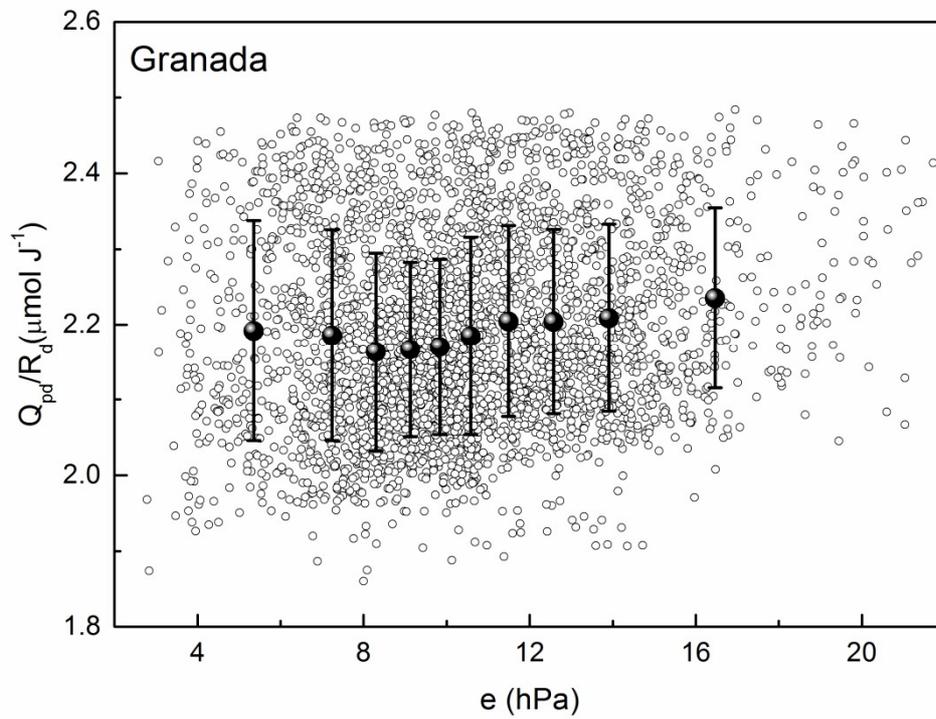
555

556

557

558

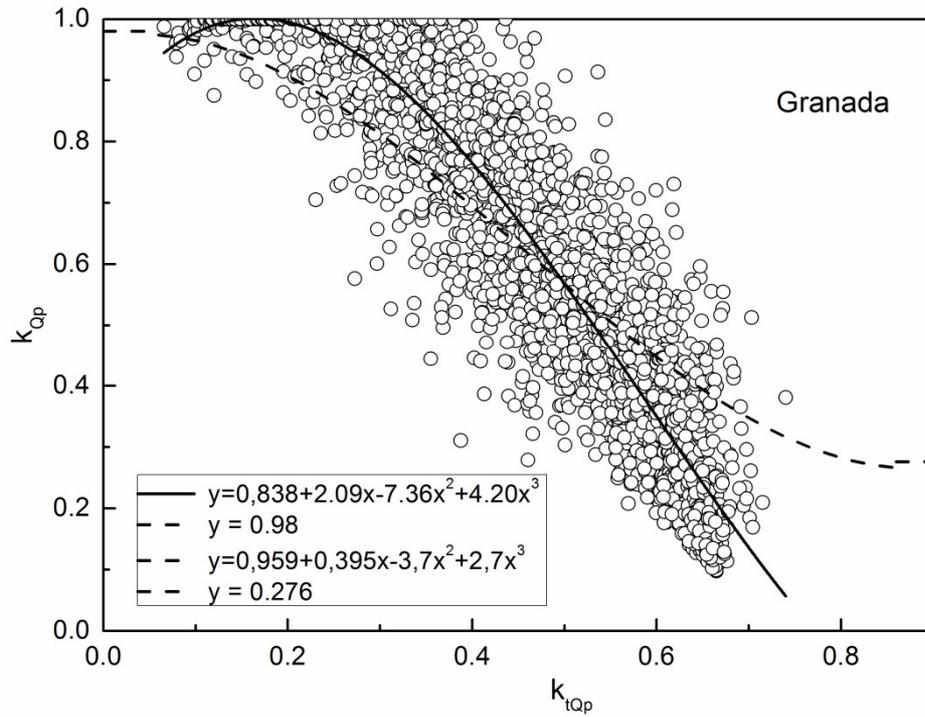
559  
560  
561



562  
563  
564  
565  
566  
567  
568  
569  
570  
571  
572  
573  
574  
575  
576  
577

Figure 7: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance ( $Q_{pd}/R_d$ ) vs. the partial vapour pressure ( $e$ ). Small dots denote experimental data, black symbols represent mean values, and bars denote standard deviations for each of the intervals in  $e$ .

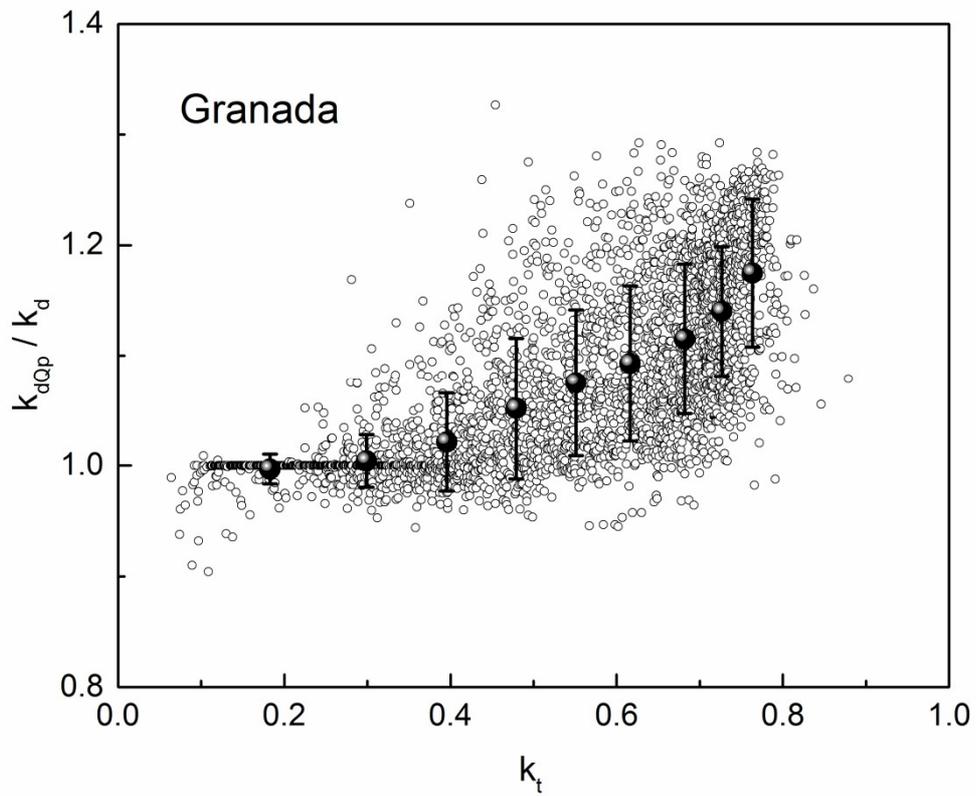
578  
579  
580  
581



582  
583 Figure 8: Diffuse fraction in the visible range ( $k_{dQp}$ ) vs. clearness index in the visible range  
584 ( $k_{tQp}$ ). Small dots denote experimental data. Dashed lines denote polynomial fitting (Jacovides  
585 et al. 2007). Solid lines denote our polynomial fitting.

586  
587  
588  
589  
590  
591  
592  
593  
594  
595

596  
597  
598  
599

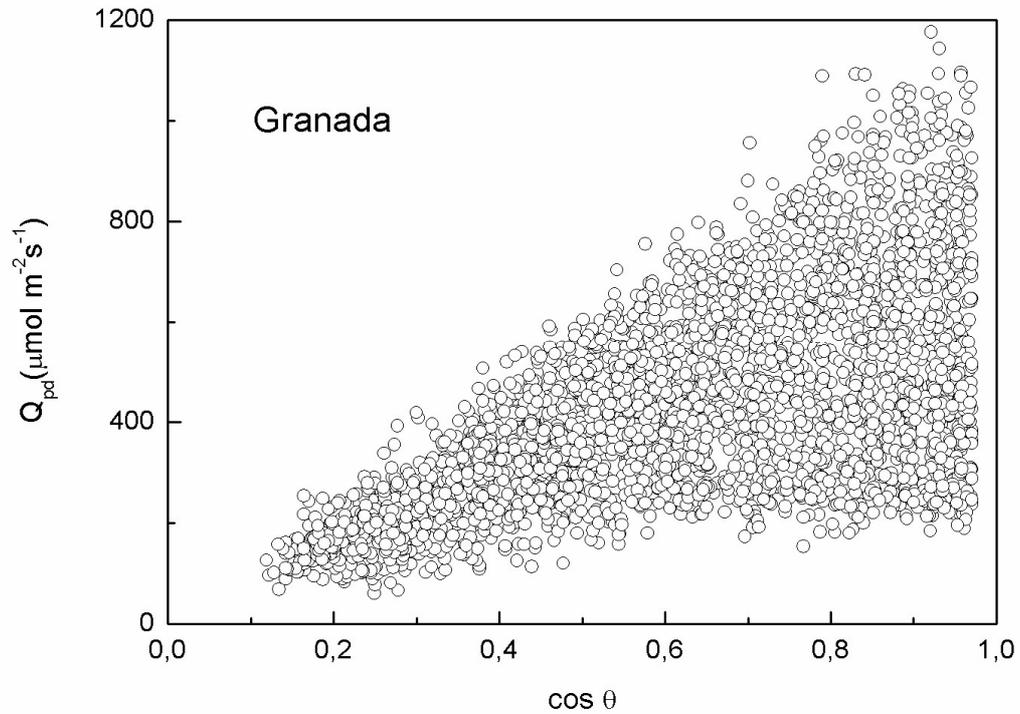


600  
601  
602  
603  
604  
605  
606  
607  
608  
609  
610  
611

Figure 9: Ratio of diffuse fraction in the visible range ( $k_{dQP}$ ) and diffuse fraction in the shortwave broadband range ( $k_d$ ) vs. clearness index ( $k_t$ ). Small dots denote experimental data, black symbols represent mean values, and bars denote standard deviations for each of the interval in  $k_t$ .

612

613



614

615

616

617 Figure 10: Diffuse photosynthetic photon flux density ( $Q_{pd}$ ) vs. cosine of the solar zenith

618 angle ( $\cos \theta$ ). Dots represent experimental data.

619

620

621

622

623

624

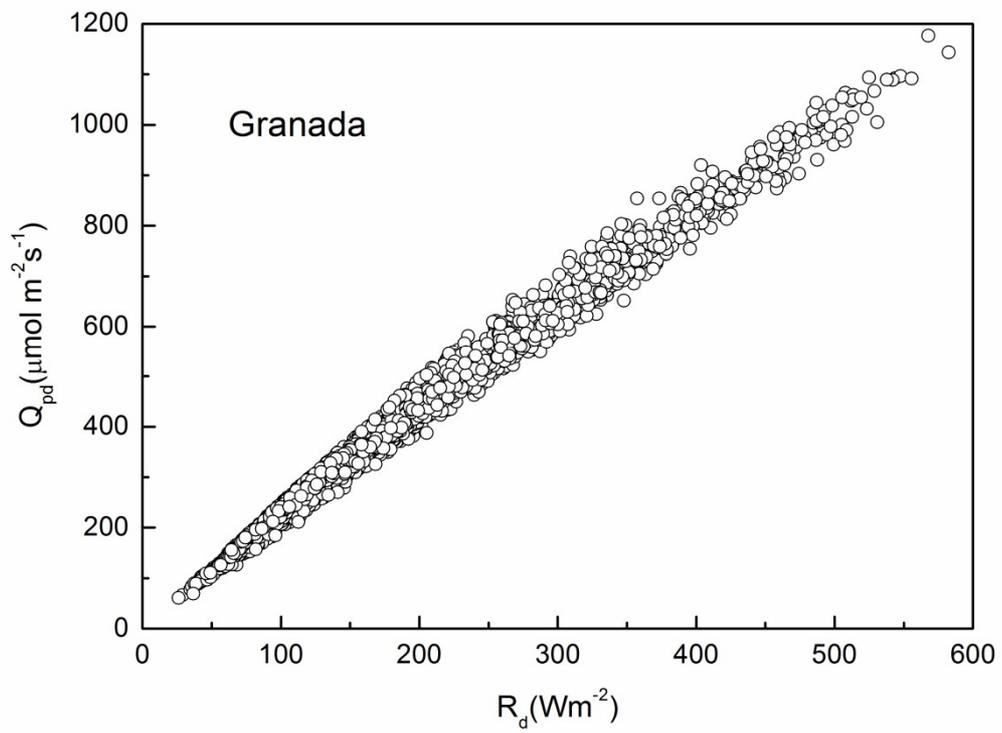
625

626

627

628

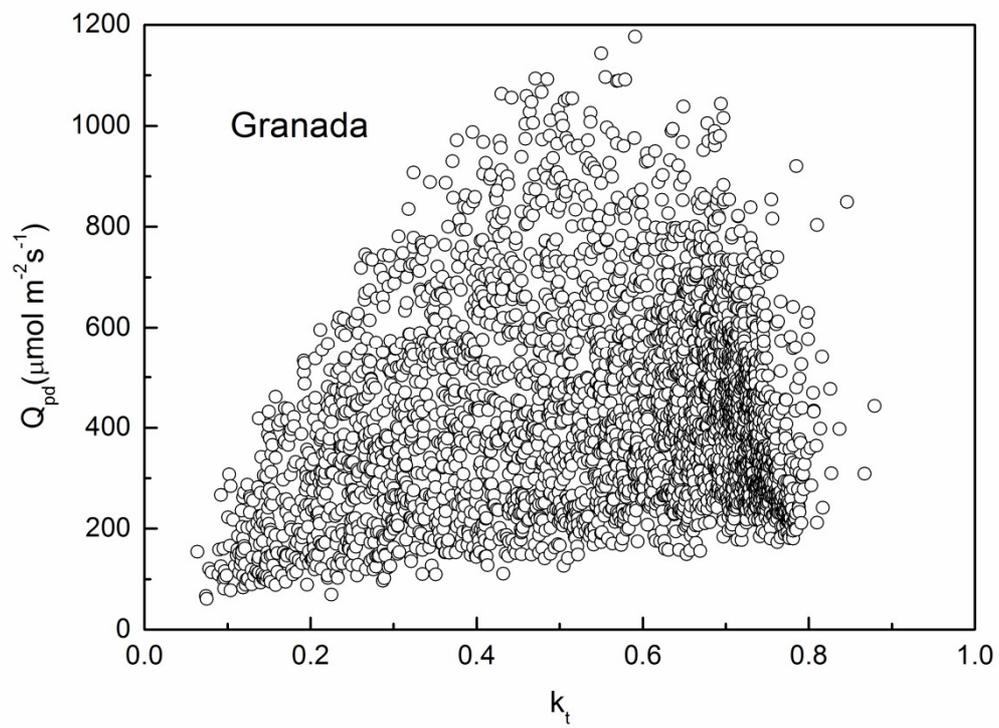
629  
630  
631  
632



633  
634  
635  
636  
637  
638  
639  
640  
641  
642  
643  
644  
645  
646

Figure 11: Diffuse photosynthetic photon flux density ( $Q_{pd}$ ) vs.  $R_d$ . Dots represent experimental data.

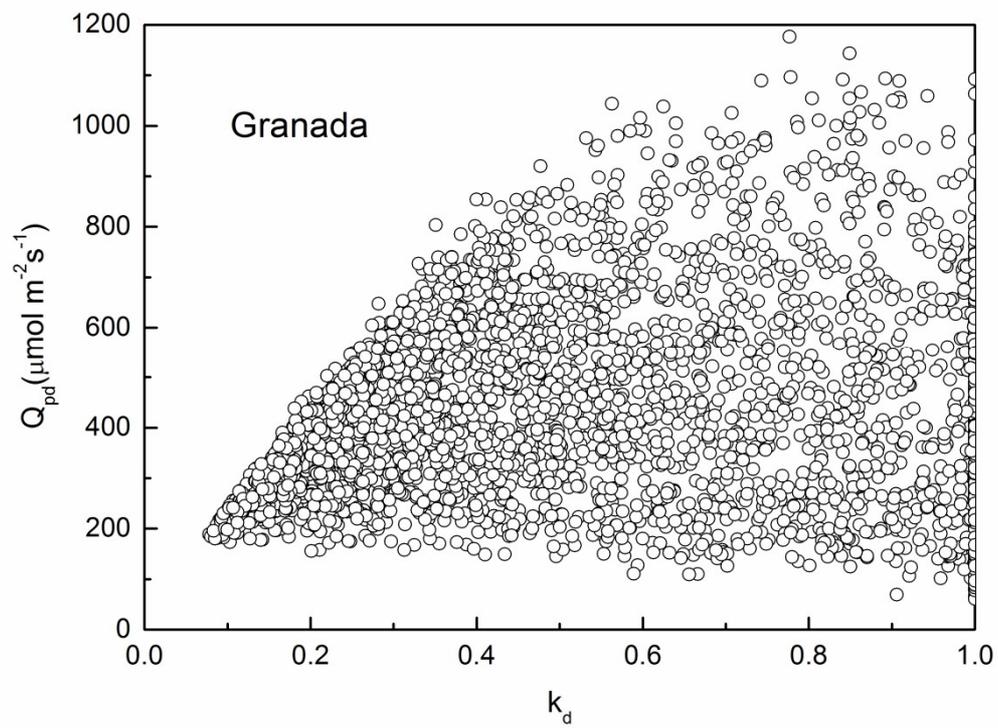
647  
648  
649  
650



651  
652  
653  
654  
655  
656  
657  
658  
659  
660  
661  
662  
663  
664

Figure 12: Diffuse photosynthetic photon flux density ( $Q_{pd}$ ) vs.  $k_t$ . Dots represent experimental data.

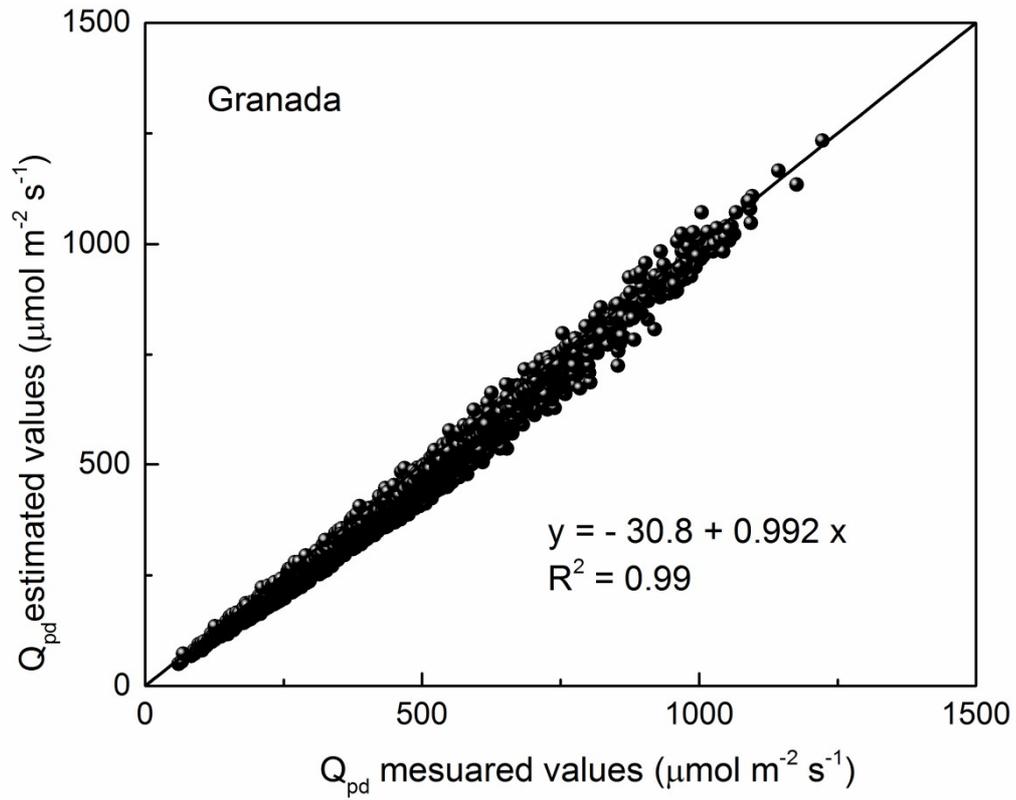
665  
666  
667  
668



669  
670  
671  
672  
673  
674  
675  
676  
677  
678  
679  
680  
681  
682  
683

Figure 13: Diffuse photosynthetic photon flux density ( $Q_{pd}$ ) vs.  $k_d$ . Dots represent experimental data.

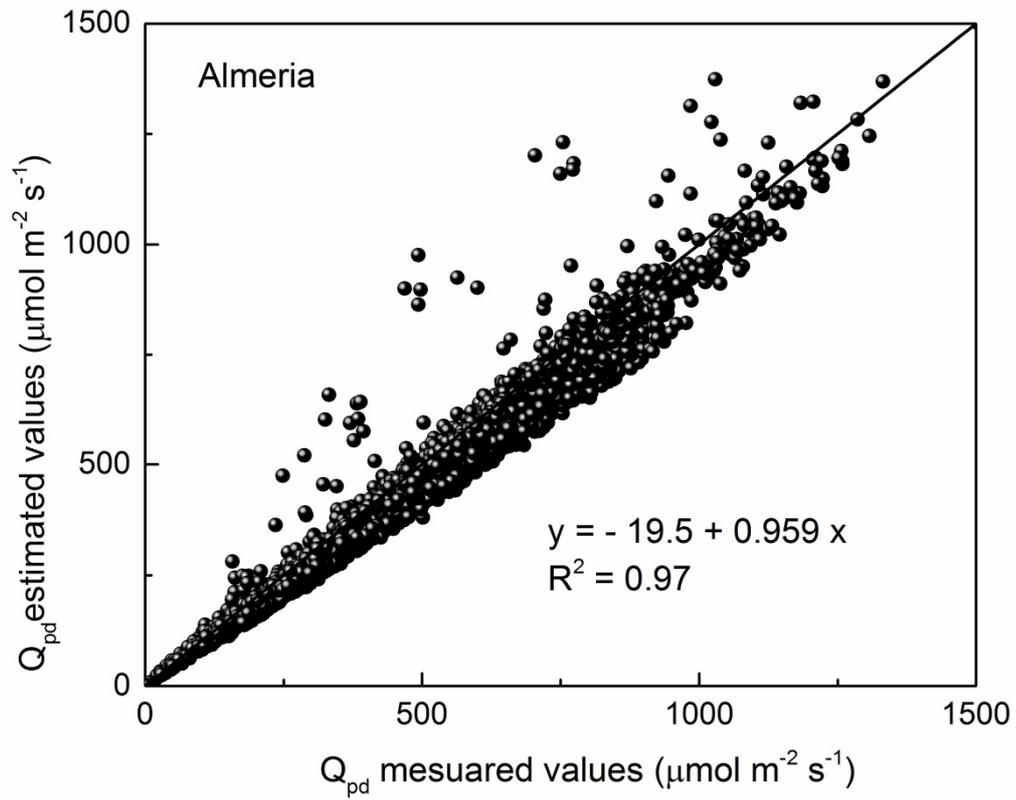
684  
685  
686  
687  
688  
689



690  
691 Figure 14 a: Scatter plot of estimated vs. measured values of photosynthetic photon flux  
692 density ( $Q_{pd}$ ) at Granada site.

693  
694  
695  
696  
697  
698  
699  
700

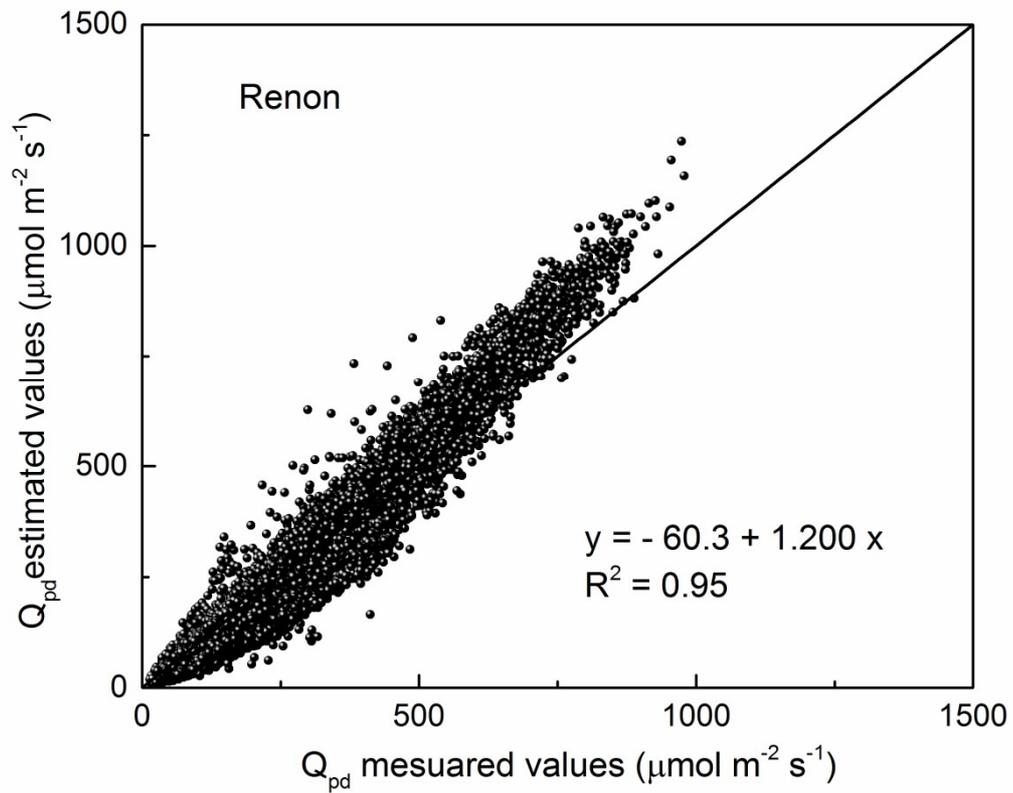
701  
702  
703  
704  
705



706  
707  
708  
709  
710  
711  
712  
713  
714  
715  
716  
717  
718

Figure 14 b: Scatter plot of estimated vs. measured values of photosynthetic photon flux density ( $Q_{pd}$ ) at Almeria site.

719  
720  
721  
722  
723  
724  
725  
726



727  
728 Figure 14c: Scatter plot of estimated vs. measured values of photosynthetic photon flux  
729 density ( $Q_{pd}$ ) at Renon site.  
730