

1 **A new empirical model to estimate hourly diffuse photosynthetic photon flux density**

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25 **ABSTRACT**

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27 Knowledge of the photosynthetic photon flux density (Q_p) is critical in different applications
28 dealing with climate change, plant physiology, biomass production, and natural illumination
29 in greenhouses. This is particularly true regarding its diffuse component (Q_{pd}), which can
30 enhance canopy light-use efficiency and thereby boost carbon uptake. Therefore, diffuse 31
31 photosynthetic photon flux density is a key driving factor of ecosystem-productivity models.
32 In this work, we propose a model to estimate this component, using a previous model to
33 calculate Q_p and furthermore divide it into its components. We have used measurements in
34 urban Granada (southern Spain), of global solar radiation (R_s) to study relationships between
35 the ratio Q_{pd}/R_s with different parameters accounting for solar position, water-vapour
36 absorption and sky conditions. The model performance has been validated with experimental
37 measurements from sites having varied climatic conditions. The model provides acceptable
38 results, with the mean bias error and root mean square error varying between -0.3 and -8.8%
39 and between 9.6 and 20.4%, respectively.

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42 **1. INTRODUCTION**

43

44 The amount of photosynthetically active radiation (PAR), defined as the visible
45 portion of global solar radiation (R_s) used in photosynthesis to convert light energy into
46 biomass (Udo and Aro, 1999; Jacovides et al., 2004; Tang et al., 2013), determines the
47 exchanges of energy and water between the land surface and the atmosphere. Among the
48 standard environmental statistics needed to evaluate plant photosynthesis are PAR data

49 (Akitsu et al., 2017). This radiation covering both photon and energy terms lies between 400
50 and 700 nm. Thus, the photosynthetic photon flux density, Q_p , is defined as the photon flux
51 density ($1 \mu\text{mol photons m}^{-2} \text{s}^{-1} = 6.022 \times 10^{17} \text{ photons m}^{-2} \text{s}^{-1}$). This is the number of photons
52 in the 400-700 nm waveband incident per unit time on a unit surface.

54 The amount and spectral quality of the radiation reaching the surface is altered by aerosols
55 and clouds by absorption and scattering processes, varying the proportion of the components
56 of radiation (direct and diffuse). Plants use both the direct and diffuse photosynthetic photon
57 flux density (Q_{pb} and Q_{pd}), but it is known that these components differ in the way they
58 transfer energy through plant canopies and affect the summation of nonlinear processes such
59 as photosynthesis differently than would occur at the leaf scale (Gu et al., 2002; Misson et
60 al., 2005; Min, 2005; Jacovides et al., 2007; 2010). In fact, the efficiency with which
61 incoming PAR is intercepted by a canopy, depends on its efficiency in intercepting direct and
62 diffuse incoming radiation and on the proportions of Q_{pd} and Q_{pb} 63 (Alados et al., 2002).

64 The analysis of the effects of Q_{pd} on ecosystem productivity has become one of the main
65 goals in terrestrial carbon-cycle research (Gu *et al.*, 1999, 2002, 2003; Mercado *et al.*,
66 2009; Sun and Zhou, 2010; Zhang *et al.*, 2010; He *et al.*, 2011; Zhang *et al.*, 2011), and thus
67 the treatment of Q_{pd} in ecosystem models is needed (Gu *et al.*, 2003; Mercado *et al.*, 2009;
68 Kanniah *et al.*, 2012). Hence, quantifying Q_{pd} and understanding its spatiotemporal variations
69 are critical for estimating its impact on the carbon cycle of terrestrial ecosystems. Currently,
70 many sites measure Q_p , although a worldwide network for routinely measuring Q_p has not yet
71 been established, despite the biological importance of this radiometric quantity. Furthermore,
72 few observation sites measure Q_{pd} , raising the need for models to estimate these data. Q_{pd} is
73 usually estimated by multiplying Q_p by the diffuse fraction of global radiation ($k_d = R_d/R_s$),
74 where R_d is diffuse irradiance, but this is only a rough estimate because k_d is not equivalent to

75 the diffuse fraction in the visible range (k_{dQp}), which is significantly greater than k_d under
76 clear skies, although almost equivalent under cloudy skies (Spitters *et al.*, 1986).
77 In the present work, we have evaluated a model to estimate Q_{pd} using a previous
78 model to estimate Q_p (Foyo-Moreno *et al.*, 2017), considering the separation of R_s into its
79 components: direct and diffuse. Prior to evaluating model performance, we have analysed
80 dependences of the Q_{pd}/R_d ratio and Q_{pd} on different parameters. These parameters describe
81 the solar position using the cosine of the solar zenith angle (θ) and indices such as sky
82 clearness ($\varepsilon = (R_d + R_b)/R_d$), sky brightness ($\Delta = R_d/R_{so}\cos\theta$), and the clearness index ($k_t =$
83 R_s/R_{so}), where R_d is diffuse irradiance, R_b is direct normal irradiance and R_{so} is the
84 extraterrestrial global solar irradiance (Pérez *et al.*, 1990; Alados *et al.*, 1996). For example,
85 Jacovides *et al.* (2007; 2010) proposed an empirical model relating k_{dQp} to the clearness index
86 Q_p ($k_{tQp} = Q_p/Q_{po}$), where Q_{po} is the extraterrestrial photosynthetic photon flux density. To
87 consider the effects of water absorption in this spectral range, we also analysed dependence
88 on meteorological parameters such as water-vapour pressure (e) or dew-point temperature
89 (T_d). The model proposed in this work uses as input data the solar position as the cosine
90 function of the solar zenith angle ($\cos\theta$) and diffuse irradiance (R_d). The model was evaluated
91 with a set of independent data at various sites, two in Spain and another in Italy, with
92 different climatological characteristics.

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94 **2. EXPERIMENTAL SITE AND MEASUREMENTS**

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96 In this work, we used data from three sites (Granada and Almería in Spain, and Renon
97 in Italy). The data from Granada were used to study the main dependence between the
98 variables of interest for this study, and all three stations were used to validate the proposed

99 model, since this proposal is based on a previous model to estimate Q_p (Foyo-Moreno et al.,
100 2017), with data compiled for Granada during other years vs. those used in the present work.
101 Ground-based data for two years at 1-min intervals were acquired at a station on the
102 outskirts of Granada (37.18° N, 3.58° W, 660 m a.s.l.). From this data base, hourly values
103 were generated for the entire two-year period to include a wide range of seasonal conditions
104 and solar zenith angles. The photosynthetic active photon-flux density (Q_p) was measured
105 using LICOR model 190 SA quantum sensors (Lincoln, Nebraska, USA). Another quantum
106 sensor has been equipped with a polar axis shadowband to measure the diffuse
107 photosynthetic active photon flux density (Q_{pd}). Global solar irradiance, R_s , was measured
108 using a Kipp and Zonen model CM-11 radiometer (Delft, Netherlands), while another CM-11
109 with a polar axis shadowband was used to measure diffuse solar irradiance (R_d). The diffuse-
110 irradiance measurements were corrected following the method proposed by Batlles et al.
111 (1995). The quantum sensor has a relative error of less than 5% relative to the values
112 measured, and global solar-irradiance measurements have an estimated experimental error
113 of about 2-3%. The calibration constants of the instruments were periodically checked
114 (Alados and Alados114 Arboledas, 1999).

115 To avoid problems associated with instrument deviations from the ideal cosine law,
116 we limited our study to solar zenith angles of less than 85° (Alados et al., 1996). We
117 evaluated the proposed model using data from different stations that had not been
118 used in building the model, two in Spain (Granada and Almería), and another
119 in Italy (Renon). Detailed information on the instrumentation and site characteristics
120 can be found for Almería in Alados and Alados-Arboledas (1999). The radiometric
121 sensors used at Almería are similar to those used at Granada. The Renon/Ritten site
122 (Italy), is operated by the Forest Service and the Agency of the Environment

123 of the Autonomous Province of Bolzano (APB). In Italy, Q_p was measured
124 by a BF2 sunshine sensor (Delta-T Devices, Burwell, Cambridge, United Kingdom).
125 This device uses an array of silicon photodiodes and a shading pattern on the
121 radiometer dome to determine Q_{pd} . The accuracy of these sensors is 15% (BF2, Delta-T
122 Devices, 2005).

127 Table 1 presents some climatic data for the stations used. Granada is an inland location
128 in south-eastern Spain, a non-industrialized medium-sized city situated in a natural basin
129 surrounded by mountains with elevations between 1000 and 3500 m a.s.l.. Near continental
130 conditions prevailing at this site are responsible for large seasonal temperature differences,
131 providing cold winters and hot summers. Most rainfall occurs during winter and spring.
132 Almería, located on the Mediterranean coast in south-eastern Spain, has frequent cloudless
133 days and high humidity. Renon is situated at 1735m a.s.l. in the Italian Alps in Bolzano (Alto
134 Adige, Italy). This site is influenced by an windy and humid alpine climate. Thus, the three
135 sites present contrasting climatic characteristics and altitudes. The yearly precipitation
136 markedly fluctuates.

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138 **3. RESULTS AND DISCUSSION**

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141 **3.1. ANALYSIS OF Q_{pd}/R_d RATIO**

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143 As in our earlier work (Foyo-Moreno et al., 2017), before the direct study of the
144 variable of interest (Q_{pd}), we analysed the ratio between Q_{pd} and R_d (Q_{pd}/R_d) with different
145 parameters. Other works (Alados et al. 1996, Yu et al. 2015, Yu and Go, 2016) showed
146 seasonal and daily variations of Q_p/R_s and also variations of Q_{pd}/R_d (Alados and Alados-

147 Arboledas, 1999). Since the most important factor influencing the solar-radiation levels
148 reaching the Earth's surface is solar position, Figure 1 shows this dependence of this ratio on
149 $\cos\theta$. We noted that this ratio varied between 1.86 and 2.48 $\mu\text{mol J}^{-1}$ with a mean value of
150 2.19 ± 0.13 ($\mu\text{mol J}^{-1}$). Jacovides et al. (2007) found a mean value in Athens (Greece) of 2.43
151 ± 0.26 ($\mu\text{mol J}^{-1}$). For all hourly values together, we found no dependence. However, for the
152 mean values of this ratio in different categories of $\cos\theta$, we detected a slight dependence with
153 high dispersion. The correlation coefficient (R^2) was 0.54 while the intercept and slope were
154 2.14 ± 0.02 ($\mu\text{mol J}^{-1}$) and 0.08 ± 0.02 ($\mu\text{mol J}^{-1}$), respectively. The positive dependence with
155 $\cos \theta$ implies a reduction in the relative contribution of photosynthetically active photon flux
156 density over the whole solar spectrum when the optical air mass increases. In Almería (Spain)
157 a similar dependence has been reported (Alados et al., 1999). In an initial approximation, we
158 could assume that this ratio was constant and we calculated Q_{pd} from R_d .

159 The second dependence analysed was on R_s (Figure 2) and the third on R_d (Figure 3).
160 There was no dependence on R_s for all hourly values but a positive correlation was found for
161 the mean values considering the categories with the R_s values. However, a clearer
162 dependence for all hourly values was detected, although with great dispersion, as a function
163 of R_d , with larger values for low values of R_d . Thus, with decreasing R_d (clear skies),
164 the diffuse component of Q_p was larger than the diffuse component for R_d .

165 In our search for a more direct relation with sky condition, Figure 4 shows the
166 dependence of this ratio on the clearness index (k_t). This parameter is defined as the ratio of
167 the global irradiance (R_s) to the extraterrestrial global irradiance (R_{so}), both on a horizontal
168 surface:

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$$k_t = \frac{R_s}{R_{so}} \quad (1)$$

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with $R_{so} = E_o I_{sc} \cos\theta$. E_o is the eccentricity correction factor and the value used of the solar

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constant (I_{sc}) is 1367 Wm^{-2} (Iqbal, 1983). Details of the calculation of $\cos\theta$ can be found by

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Iqbal (1983). The parameter k_t characterizes the sky condition including the attenuation

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effects of the most dominant factors controlling solar radiation, such as clouds and aerosols.

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Previous studies used similar parameters such as the sky clearness and the skylight brightness

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(Pérez et al., 1990) to characterize the sky condition (Alados et al., 1996). Another parameter

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to characterize sky condition is diffuse fraction (k_d), defined as the ratio between diffuse

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global irradiance (R_d) and global irradiance (R_s):

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$$k_d = \frac{R_d}{R_s} \quad (2)$$

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Knowledge of k_d can be useful to get an idea of atmospheric load indirectly, where a

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low k_d indicates a clear sky and more pristine atmosphere, while high k_d values denote high

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aerosol loads (Singh et al., 2013). In fact, the ratio between R_d and R_{bn} can be used to

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estimate aerosol optical depth (Foyo-Moreno et al., 2014).

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Figures 4 and 5 show the dependence of Q_{pd}/R_d on these parameters. There is a slight

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dependence on k_t for all hourly data and for mean values in categories of k_t values, higher

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values of the Q_{pd}/R_d ratio are detected for low and high k_t values. These high values of Q_{pd}/R_d

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for high values of k_t are consistent with Jacovides et al. (2007), who reported a weak

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correlation for daily values, which is well fit by an exponential equation with a determination

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coefficient of 0.276. They found hourly ratio values varying from 2.56 ± 0.27 to 2.42 ± 0.24

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($\mu\text{mol J}^{-1}$), and ratios under clear skies 5.5% higher than that for overcast skies. These results

190 agree with Alados and Alados-Arboledas (1999) and Min (2005). In any case, high k_t values
191 imply low k_d values, so that this ratio had high values for low values of k_d (clear skies)
192 although the dispersion of data was high (Figure 5), increasing the dispersion with k_d . Under
193 clear skies, the regulating factors in the solar radiation modification are scattering processes,
194 so that spectral investigations support the premise that, under these skies with aerosol loads,
195 short wavelengths are preferentially scattered, thus increasing Q_{pd} more than R_d (Jacovides et
196 al., 2000; Dye, 2004). However, the pattern observed for low values of k_t (<0.4) is different
197 (Figure 4), the ratio decreasing when k_t increases. This result could be associated with the
198 mixed effects of aerosols and clouds, because these cases correspond to skies with high
199 opacity.

200 From the above analysis, it is clear that this ratio varies with changing sky conditions.
201 Clouds exhibit relatively constant extinction across the visible spectrum while aerosols
202 commonly show extinction that significantly diminishes the greater the wavelength (Min,
203 2005; Jacovides et al., 2007). This result can be attributed to the presence of clouds in skies
204 with the high solar absorption in the infrared region (NIR). Clouds absorb NIR more strongly
205 than Q_p , and hence the transmittance of Q_p through clouds is larger than NIR. In fact, clouds
206 more markedly attenuate the total solar spectral range than they do shorter wavelengths such
207 as the ultraviolet range (Foyo-Moreno et al., 2001, 2003).

208 As in an earlier work (Foyo-Moreno et al., 2017), here, to explicitly consider
209 absorption by water vapour in the solar spectrum, we have shown in Figures 6 and 7 the
210 dependence on meteorological parameters such as the dewpoint temperature (T_d) – relevant
211 due to its correlation with precipitable water (Reitan, 1963) – and the partial vapour pressure
212 (e). T_d and e have been calculated from the direct measurements of air temperature (T) and
213 the relative humidity (U) at our station, with standard formulation. As before

214 (Foyo-Moreno et al., 2017), no dependence was detected.

215

216 **3.2. ANALYSIS OF PARAMETERS CHARACTERIZING SKY CONDITIONS**

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218 Now, in order to find relationships between parameters considering sky conditions, it
219 is possible to define indexes for the visible range equivalent to those for the shortwave
220 broadband range. These parameters are $k_{tQP} = Q_p/Q_{po}$ (clearness index) and $k_{dQP} = Q_{pd}/Q_p$
221 (diffuse fraction), defined as in Eqs. 1 and 2, respectively but for the visible range. Q_{po} can be
222 derived from R_{so} , with a fraction of 0.5 suggested by most studies (Frouin and Pinker 1995;
223 Hu et al. 2007, 2010; Wang et al. 2014), then Q_{po} can be calculated by multiplying by the
224 ratio 4.57 of the energy-flux density to photosynthetic photon flux density (Dye 2004).

225 Figure 8 shows the dependence between k_{dQP} and k_{tQP} . Jacovides et al. (2007) found a
226 polynomial fitting, shown in Figure 8 as the dashed line for $0.1 < k_{tQP} < 0.85$. Our fitting is also
227 a third-order polynomial but with different coefficients (solid line) with a determination
228 coefficient of 0.88. For low values of k_{tQP} the diffuse component is similar to Q_p and for high
229 values this proportion decreases to about 10% of Q_p . In the literature, few studies have
230 focused on the empirical relationships between Q_p and Q_{pd} of the Liu and Jordan regression
231 type (Jacovides et al., 2010). The applicability of several diffuse radiation empirical models
232 has been analysed in Athens (Jacovides et al., 2010), showing mean bias error (MBE) values
233 between -12.9 and 2.99 % and root mean square error (RMSE) values between 27.1 and
234 35.6%, with a MBE of 0.11% and a RMSE of 26.8% for their fitting.

235 Figure 9 shows the dependence between the ratio of the diffuse fraction in the visible
236 range (k_{dQP}) and the diffuse fraction in the shortwave broadband range (k_d) and clearness
237 index (k_t). We found a positive correlation, with the ratio increasing with clear skies. Spitters

238 et al. (1986) showed that the diffuse fraction for the visible range was 1.4 times the diffuse
239 broadband fraction under clear sky conditions. It is evident that this ratio increases
240 significantly when sky conditions change from cloudy to clear. In our case, the value found
241 was 1.17 ± 0.07 for clear skies. Jacovides et al. (2007) reported a value of 1.38 ± 0.32 . The
242 dispersion increases with higher k_t values. However, for low values of k_t (<0.4 ; cloudy skies)
243 the diffuse fraction in the visible range is of the same order as the diffuse fraction in the
244 shortwave broadband range.

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246 **3.3. ANALYSIS OF Q_{pd} .**

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248 In this section, we directly study the dependence of Q_{pd} on the same parameters
249 considered in Section 3.1. First, we show that the solar zenith angle alone does not allow for
250 a correct determination of Q_{pd} (Figure 10). While Q_{pd} increases with decreasing θ , due to
251 high dispersion it is possible to parameterise only the two *envelopes* of the data through a
252 simple linear dependence on the cosine of θ . This result is similar to earlier work with Q_p
253 (Foyo Moreno et al., 2017), and with the ultraviolet range (Foyo-Moreno et al., 1998, 2007).
254 The upper envelope of the data corresponds to the maximum values that can be assumed to
254 correspond to overcast skies, and the lower envelope can be attributed to values
255 corresponding to clear skies for a given solar position.

257 We show the high positive correlation between Q_{pd} and R_d ($R^2= 0.997$), with a slope of
258 $2.125 \pm 0.002 \mu\text{mol J}^{-1}$ (Figure 11), close to the mean value found for the ratio Q_{pd}/R_d [$2.19 \pm$
259 $0.13 (\mu\text{mol J}^{-1})$]. The dependence on sky conditions using the parameters k_t and k_d is shown
260 in Figures 12 and 13. The dispersion of data is high for all sky conditions.

261 Considering all of the above results, we found it more appropriate to model Q_{pd} rather

262 than Q_{pd}/R_d . Therefore, taking into account our previous model (Foyo-Moreno et al., 2017),
 263 the expression of Q_p is:

$$264 \quad Q_p = a k_t \cos \theta \quad (3)$$

265 We can separate Q_p into its two components: direct and diffuse. Also, taking into account that
 266 global irradiance (R_s) is the sum of the two components (direct and diffuse), we arrive at this
 267 expression for Q_{pd} :

$$268 \quad Q_{pd} = a \cos \theta \frac{R_d}{R_{so}} \quad (4)$$

268 with $a = (2681 \pm 2) \mu\text{mol m}^{-2}\text{s}^{-1}$.

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271 **3.4. MODEL PERFORMANCE**

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273 We have evaluated the model (Eq. 4) at the three stations which were not used in its
 274 development. Table 2 shows the results found after including the correlation coefficient R^2 –
 275 (the fraction of experimental data variance explained by the model), slope b , and the intercept
 276 a of the linear regression of the Q_{pd} measured vs. the estimated values. Table 2 also shows the
 277 mean bias error (MBE) and root mean square error (RMSE), both as percentages of the mean
 278 experimental values. The MBE and RMSE, are given by the following expressions:

$$279 \quad MBE = \frac{100}{M_{ave}} \frac{1}{N} \sum_{i=1}^N (E_i - M_i) \quad (5)$$

$$280 \quad RMSE = \frac{100}{M_{ave}} \left(\frac{1}{N} \sum_{i=1}^N (E_i - M_i)^2 \right)^{0.5} \quad (6)$$

281 where E_i is the estimated value (i th number), M_i is the measured value, M_{ave} is the average of

282 the measured values and N is the number of observations. These two statistics enable the
283 detection both of the differences between experimental data and of the model estimates and
284 the existence of systematic over- or underestimation trends, respectively.

285 Figure 13 shows Q_{pd} estimated via Eq. (4) at all localities analysed in this work. The
286 model evaluation is excellent for all radiometric stations, the variance explained for the
287 models is better than 95%, and the values for MBE and RMSE vary between -0.3 and 8.6 %
288 and 9.6 and 20.4%, respectively. The model underestimates the values for all localities. In any
289 case, both the slope and correlation coefficient of the linear regression between the measured
290 and estimated values reveals the goodness of the model estimations. Wang et al. (2017), using
291 different soft computing techniques, found MBE values between -16 and -18% and RMSE
292 ranging between 22 and 51% for six stations from the AmeriFlux network.

293

294 **3. CONCLUSIONS**

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296 Two years of measurements made in Granada, an urban site in south-eastern Spain,
297 were used to study the diffuse photosynthetic photon-flux density (Q_{pd}), with the aim of
298 proposing a model for all sky conditions. The model was evaluated at three stations, two in
299 Spain (Granada and Almería) and another in Italy (Renon), with different climatic
300 characteristics. Firstly, we analysed the ratio of Q_{pd} to diffuse irradiance (R_d) to continue with
301 Q_{pd} . The dependence of both (Q_{pd}/R_d and Q_{pd}) on different parameters, which characterize
302 solar position and sky conditions, were considered together with other meteorological factors
303 such as vapour pressure and dewpoint temperature (T_d). The main findings can be
304 summarized as follows:

305

- 306 1. A mean value of $2.19 \pm 0.13 \mu\text{mol J}^{-1}$ was found for Q_{pd}/R_d with values varying between
307 1.86 and $2.48 \mu\text{mol J}^{-1}$.
- 308 2. Q_{pd}/R_d showed dependence on parameters that characterize sky conditions such as
309 clearness index (k_t) and diffuse fraction (k_d).
- 310 3. A model was proposed to estimate Q_{pd} , taking into account a previous model to estimate
311 the photosynthetic active radiation-flux density (Q_p). This model uses the solar zenith
312 Angle (θ) and R_d as input data.
- 313 4. The model provided acceptable results, with a mean bias error varying between -0.3% and
314 8.8%; and the root mean square error (RMSE) varying between 9.6 and 20.4%.
- 315 5. This work suggests that this empirical model can be widely applied from available
316 measurements in most radiometric stations, with better accuracy than other empirical
models.

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324

325 REFERENCES

326 Akitsu, T., Nasahara, K.N., Hirose, Y., Ijima, O., Kume, A., 2017. Quantum sensors for
327 accurate and stable long-term photosynthetically active radiation observations. Agric. For.
328 Meteorol. 237-238, 171-183.

329 Alados, I., Foyo-Moreno, I., Alados-Arboledas, L., 1996. Photosynthetically active radiation:
330 measurements and modeling. *Agric. For. Meteorol.* 78, 121-131.

331 Alados, I., Alados-Arboledas L., 1999. Direct and diffuse photosynthetically active radiation:
332 measurements and modelling. *Agric. For. Meteorol.* 93, 27-38

333 Alados, I., Olmo, F.J., Foyo-Moreno, I., Alados-Arboledas, L., 2000. Estimation of
334 photosynthetically active radiation under cloudy conditions. *Agric. For. Meteorol.* 102, 39-
335 50.

336 Alados, I., Foyo-Moreno, I., Olmo, F.J., Alados-Arboledas, L., 2002. Improved estimation of
337 diffuse photosynthetically active radiation using two spectral models. *Agric. For. Meteorol.*
338 101, 1-12.

339 Batlles, F., Olmo, F., Alados-Arboledas, L., 1995. On shadowband correction methods for
340 diffuse irradiance measurements. *Solar Energy* 54, 105-114

341 Dye, D., 2004. Spectral composition and quanta-to-energy ratio of diffuse photosynthetically
342 active radiation under diverse cloud conditions. *J. Geophys. Res.* 109, D10203
343 10.1029/2003JD004251

344 Foyo-Moreno, I., Vida, J., Alados-Arboledas, L., 1998. A simple all weather model to
345 estimate ultraviolet solar radiation (290-385nm). *J. Appl. Meteorol.*, 38, 1020-1026.

346 Foyo-Moreno, I., Alados, I., Olmo, F.J., Vida, J., Alados-Arboledas, L., 2001. On the
use of a cloud modification factor for solar UV (290-385 nm) spectral range. *Theor. Appl.*
348 *Climatol.*, 68, 41-50.

349 Foyo-Moreno, I., Alados, I., Olmo, F.J., Alados-Arboledas, L., 2003. The influence of
350 cloudiness on UV global irradiance (295-385 nm). *Agric. For. Meteorol.* 120, 101-111.

351 Foyo-Moreno, I., Alados, I., Alados-Arboledas, L., 2007. Adaptation of an empirical model
352 for erythemal ultraviolet irradiance. *Ann. Geophys.*, 25, 1-10.

353 Foyo-Moreno, I., Alados, I., Antón, M., Fernández-Gávez, J., Cazorla, A., Alados-Arboledas,
354 L., 2014. Estimating aerosol characteristics from solar irradiance measurements at an urban
355 location in southeastern Spain. *J. Geophys. Res.* doi: 10.1002/2013JD020599104

356 Foyo-Moreno, I., Alados, I., Alados-Arboledas, L., 2017. A new conventional regression
357 model to estimate hourly photosynthetic photon flux density under all sky conditions. *Int.*
358 *J. Climatol.* doi: 10.1002/joc.5063

359 Frouin, R., Pinker, R.T., 1995. Estimating photosynthetically active radiation (PAR) at the
360 earth's surface from satellite observations. *Remote Sens Environ* 51:98–107

361 Gu, L., Fuentes, J.D., Shugart, H.H, Staebler, R. M., Black, T.A., 1999. Responses of net
362 ecosystem exchanges of carbon dioxide to changes in cloudiness: Results from two North
363 American deciduous forests. *J. Geophys. Res.*, 104, 31 421–31 434

364 Gu, L., Baldocchi, D., Verma, S.B., Black, T.A., Vesala, T., Falge, E.M., et al. 2002.
365 Advantages of diffuse radiation for terrestrial ecosystem productivity. *J Geophys Res*,
366 107(D6). doi:10.1029/2001JD001242

367 Gu, L.H., Baldocchi, D.D., Wofsy, S.C., et al., 2003. Response of a deciduous forest
368 to the Mount Pinatubo eruption: Enhanced photosynthesis. *Science*, 299(5615): 2035–2038.

369 He, X.Z., Zhou, T., Jia, G.S. et al., 2011. Modeled effects of changes in the amount and
370 diffuse fraction of PAR on forest GPP. *Journal of Natural Resources*, 26(4): 619–634

371 Hu, B., Wang, Y., Liu, G., 2007. Spatiotemporal characteristics of photosynthetically active
372 radiation in China. *J. Geophys. Res.*, 112, doi: 10.1029/2006JD007965.

373 Hu, B., Wang, Y., Liu, G., 2010. Variation characteristics of ultraviolet radiation derived
374 from measurement and reconstruction in Beijing, China, *Tellus B* 62, 100–108.

375 Iqbal, M. 1983. *An Introduction to Solar Radiation*. Academic Press: London.

376 Jacovides, C.P., Tymvios, F.S., Papaioannou, G., Asimakopoulos, D.N., Theofilou, K.M.,

377 2004. Ratio of PAR to broadband solar radiation measured in Cyprus, *Agric. For. Meteor.*,
378 121, 135-140.

379 Jacovides, C.P., Tymvios, F.S., Asimakopoulos, V.D., Kaltsounides, N.A., 2007. The
380 dependence of global and diffuse PAR radiation components on sky conditions at Athens,
381 Greece, *Agric. For. Meteor.*, 143, 277-287.

382 Jacovides, C.P., Boland, J., Asimakopoulos, D.N., Kaltsounides, N.A., 2010. Comparing
383 diffuse radiation models with one predictor for partitioning incident PAR radiation into its
384 diffuse component in the eastern Mediterranean basin, *Renew. Energy*, 35, 1820-1827.

385 Kanniah, K.D., Beringer, J., North, P., Hutley, L., 2012. Control of atmospheric particles on
386 diffuse radiation and terrestrial plant productivity A review. *Progr Phys Geogr* , 36(2), 209-
387 237.

388 Mercado, L.M., Bellouin, N., Sitch, S. et al., 2009. Impact of changes in diffuse radiation on
389 the global land carbon sink. *Nature*, 458(7241): 1014–1017.

390 Min, Q., 2005. Impacts of aerosols and clouds on forest-atmosphere carbon exchange. *J*
391 *Geophys Res*110:D066203. doi:10.1029/2004JD004858.

392 Misson, L., Lunden, M., McKay, M., Goldstein, A.H., 2005. Atmospheric aerosol light
393 scattering and surface wetness influences the diurnal pattern of net ecosystem exchange in
394 a semi-arid ponderosa pine plantation. *Agric For Meteorol* 129:60–83.

395 Perez, R., Ineichen, P., Seals, R., Michalsky, J.J., Stewart, R., 1990. Modelling daylight
396 availability and irradiance components from direct and global irradiance. *Solar Energy* 44,
397 271-289.

398 Reitan, C.H., 1963. Surface dewpoint and water vapour aloft. *J. Appl. Meteorol.*, 2: 776-779.

399 Singh, J., Bhattacharya, B.K., Kumar, M., Mallick, K.. 2013. Modelling monthly diffuse
400 solar radiation fraction and its validity over the Indian sub-tropics. *Int. J. Climatol.*, 33, 77-

401 86.

400 Spitters, C.J.T., Toussaint, H.A.J.M., Goudriaan, J. 1986. Separating the diffuse and direct
401 component of global radiation and its implications for modelling canopy photosynthesis.
402 *Agric For Meteorol*, 38:217–29

403 Sun, J.S., Zhou, G.S., 2010. Review of advances in measurements and effects of diffuse
404 radiation on terrestrial ecosystem productivity. *Chinese Journal of Plant Ecology*, 34(4):
405 452–461. (in Chinese).

407 Tang, W., Qin, J., Yang, K., Niu, X., Zhang, X., 2013. Reconstruction of daily
408 photosynthetically active radiation and its trend over China. *J. Geophys. Res. Atm.*, 118,
409 13292-13302.

410 Udo, S., Aro, T., 1999. Global PAR related to global solar radiation for central Nigeria.
411 *Agric. For. Meteorol.* 97, 21–31, [http://dx.doi.org/10.1016/S0168-1923\(99\)55-6](http://dx.doi.org/10.1016/S0168-1923(99)55-6).

412 Wang, L., Gong, W., Ma, Y., Hu, B., Zhang, M., 2014. Photosynthetically active radiation
413 and its relationship with global solar radiation in Central China. *Int. J. Biometeorol.*, 58,
414 1265-1277.

415 Wang, L., Hu, B., Kisi, O., Zounemat-Kermani, M., Gong, W., 2017. Prediction of diffuse
416 photosynthetically active radiation using different soft computing techniques. *Q. J. R.
417 Meteorol. Soc.*, doi: 10.1002/qj.3081.

418 Yu, X., Wu, Z., Jiang, W., Guo, X., 2015. Predicting daily photosynthetically active radiation
419 from global solar radiation in the Contiguous United States. *Energy Convers. Manage.*, 89,
420 71-82.

421 Yu, X., Guo, X., 2016. Hourly photosynthetically active radiation estimation in Midwestern
422 United States from artificial neural networks and conventional regression models. *Int. J.
423 Biometeorol.*, 60, 1247-1259.

424 Zhang, M., Yu, G.R., Zhang, L.M. et al., 2010. Impact of cloudiness on net ecosystem
425 exchange of carbon dioxide in different types of forest ecosystems in China.
426 *Biogeosciences*, 7(2): 711–722.

427 Zhang, M., Yu, G.R., Zhuang, J. et al., 2011. Effects of cloudiness change on net ecosystem
428 exchange, light use efficiency, and water use efficiency in typical ecosystems of China.
429 *Agric.For.Meteorol.*151(7),803–816.

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Sites	T (°C)	TM (°C)	Tm (°C)	R (mm)	U (%)
Granada, Spain (37.16° N, 3.6° W, 650 m a.s.l.)	15.4	23.0	7.8	365	58
Almería, Spain (36.83° N, 2.41° W, 21 m a.s.l.)	19.1	23.4	14.7	200	65
Renon, Italy (46.42° N, 11.28° E, 1735 m a.s.l.)	4.1	18.0	5.2	1010	

432 Table 1: Description and climatic data of study sites. T (yearly average temperature), TM
433 (yearly average of the maximum temperature), Tm (yearly average of the minimum
434 temperature), R (yearly average of precipitation), U (yearly average of relative humidity).

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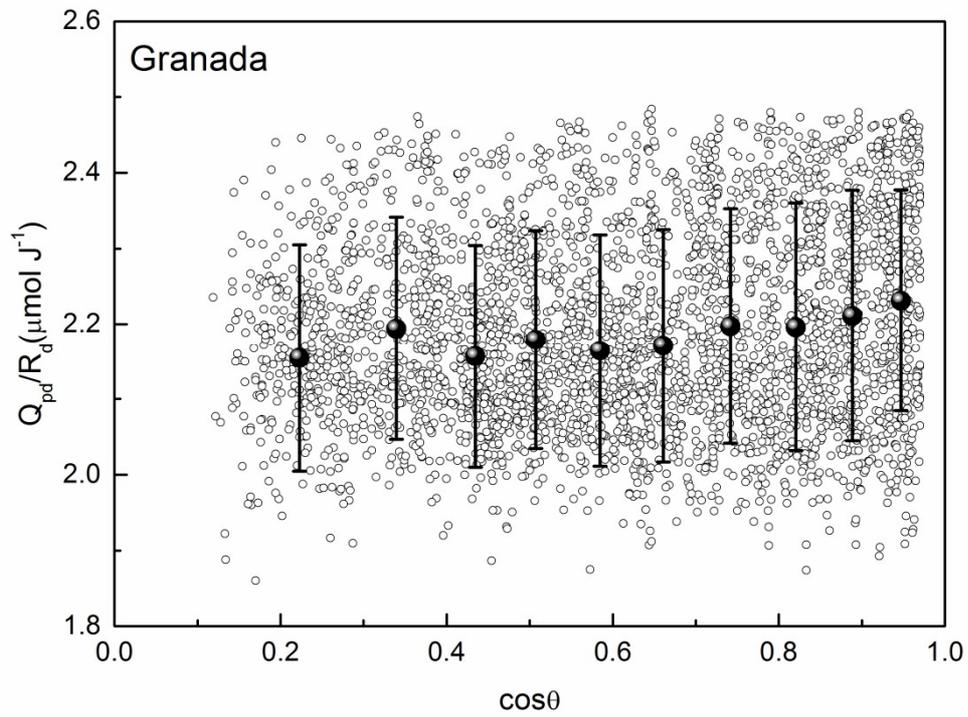
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Sites	Years	N	kt	k	Q_{pd,ave} (μmol m⁻²s⁻¹)	a (μmol m⁻²s⁻¹)	b	R₂	MBE (%)	RMSE (%)
Granada	94-95	3448	0.54±0.19	0.55±0.30	428	-30.8 ± 0.8	0.992 ± 0.002	0.99	-7.7	9.6
Almeria	93-95	5797	0.58±0.18	0.50±0.30	419	-19.5 ± 1.1	0.959 ± 0.002	0.96	-8.8	12.9
Renon	13-15	16980	0.45 ±0.27	0.59±0.35	296	-60.3 ±0.7	1.200±0.002	0.95	-0.3	20.4

457 Table 2: Statistical results of the new empirical model to estimate Q_{pd} at each station. N (total number of observations). Q_{pd,ave}

458 (average values of Q_{pd}). Linear regression statistics: a (intercept), b (slope) and R² (correlation coefficient). MBE (Mean Bias Error).

459 RMSE (Root Mean Square Error).



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463 Figure 1: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance (Q_{pd}/R_d) vs.

464 the cosine of solar zenith angle ($\cos \theta$). Small dots denote experimental data, black symbols

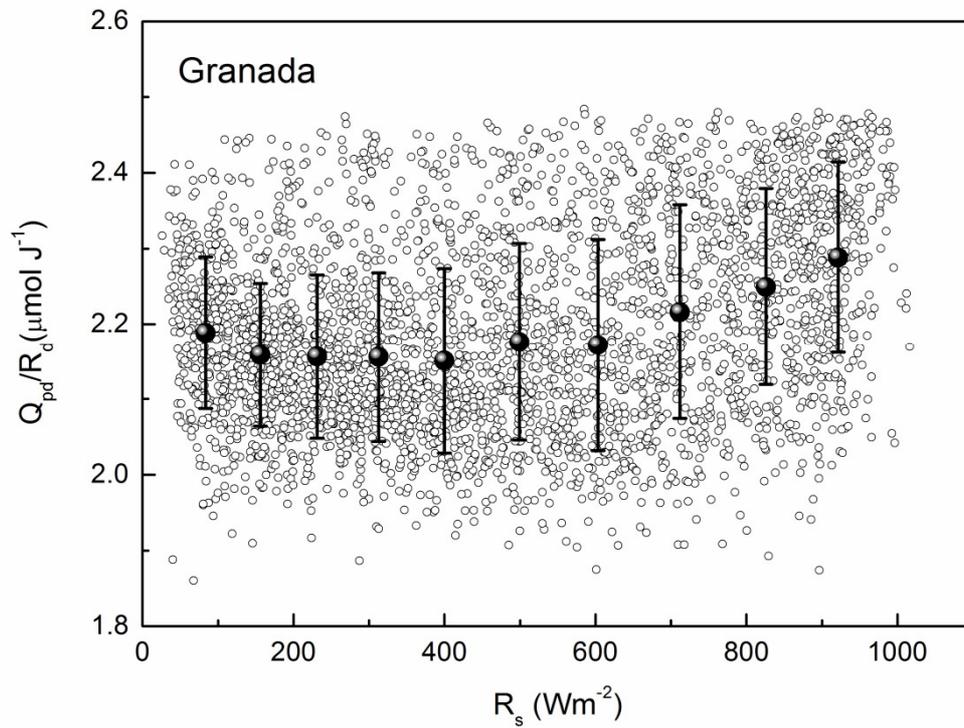
465 represent mean values, and bars denote standard deviations for each of the intervals in $\cos \theta$.

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473 Figure 2: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance (Q_{pd}/R_d) vs.
 474 broadband solar irradiance (R_s). Small dots denote experimental data, black symbols represent
 475 mean values, and bars denote standard deviations for each of the intervals in R_s .

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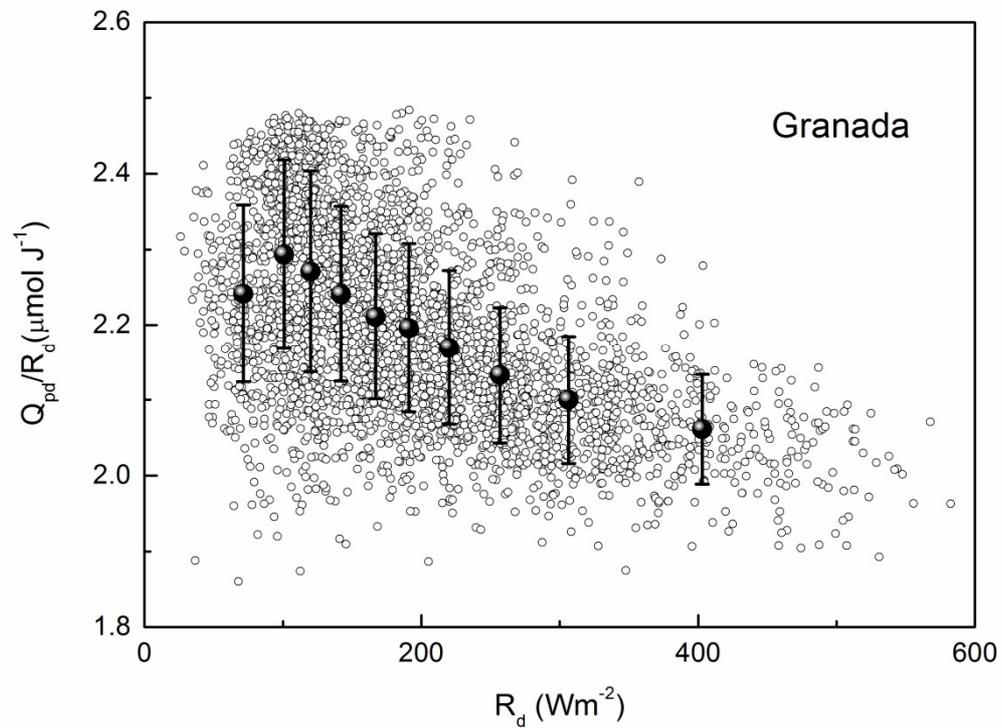
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492 Figure 3: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance (Q_{pd}/R_d) vs.493 diffuse irradiance (R_d). Small dots denote experimental data, black symbols represent mean494 values, and bars denote standard deviations for each of the intervals in R_d .

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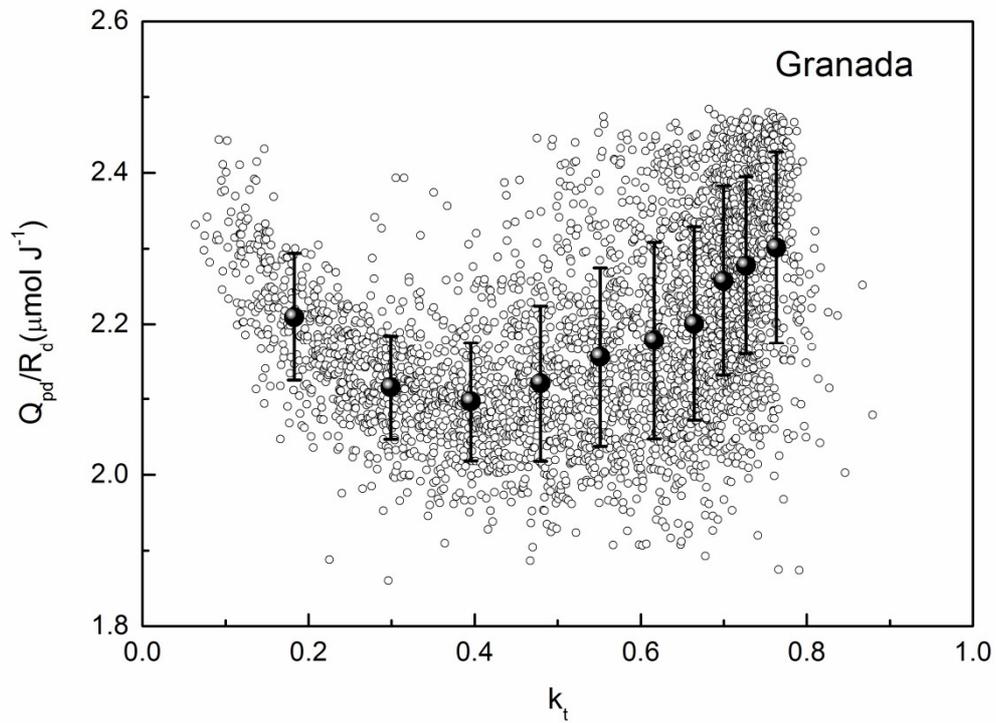
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511 Figure 4: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance (Q_{pd}/R_d) vs.

512 clearness index (k_t). Small dots denote experimental data, black symbols represent mean

513 values, and bars denote standard deviations for each of the intervals in k_t .

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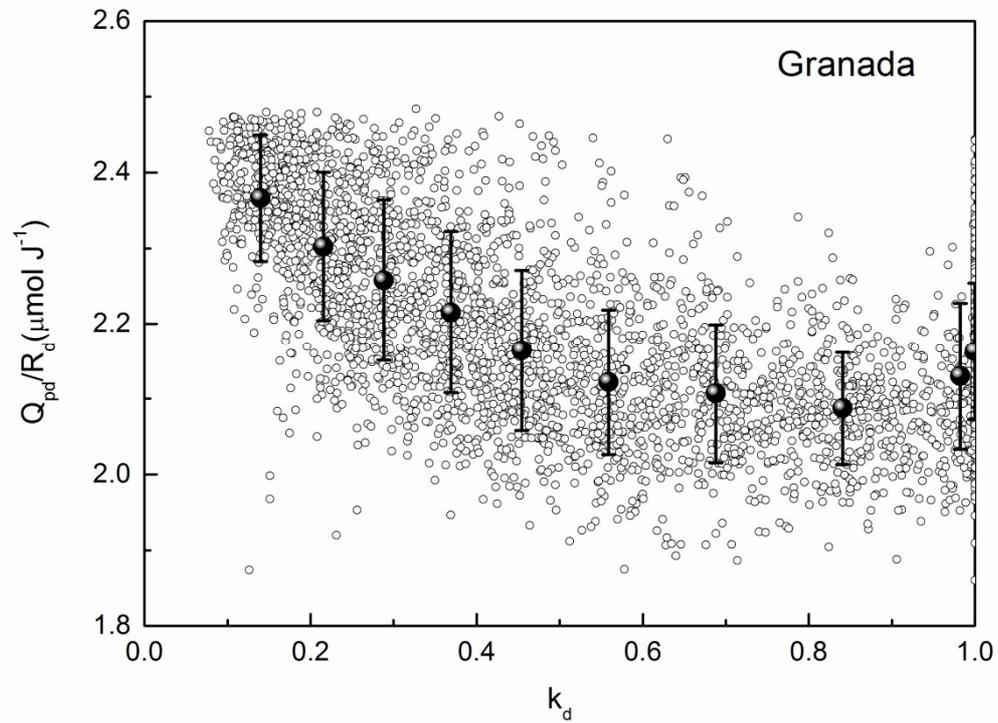
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529 Figure 5: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance (Q_{pd}/R_d)

530 vs. k_d . Small dots denote experimental data, black symbols represent mean values and bars

531 denote standard deviations for each of the intervals in k_d .

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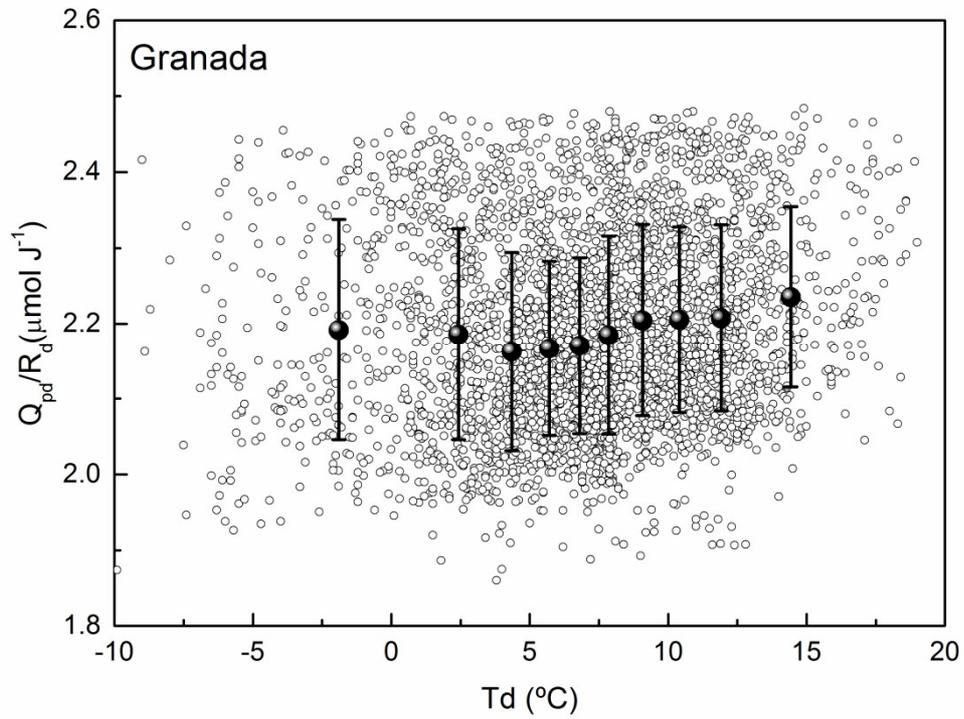
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546 Figure 6: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance (Q_{pd}/R_d) vs.
547 the dewpoint temperature (T_d). Small dots denote experimental data, black symbols represent
548 mean values, and bars denote standard deviations for each of the intervals in T_d .

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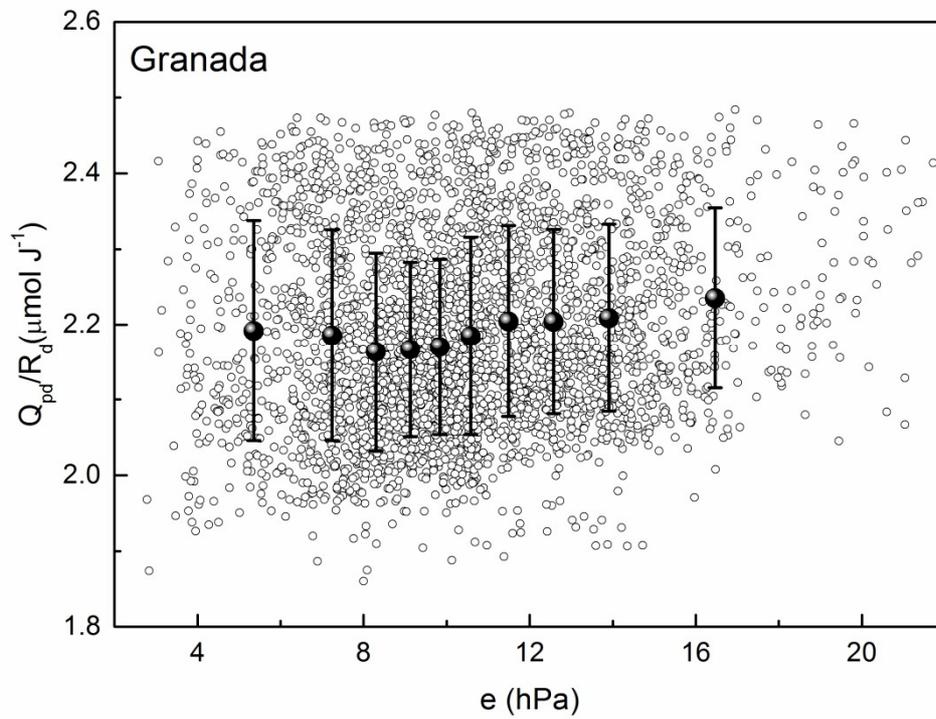
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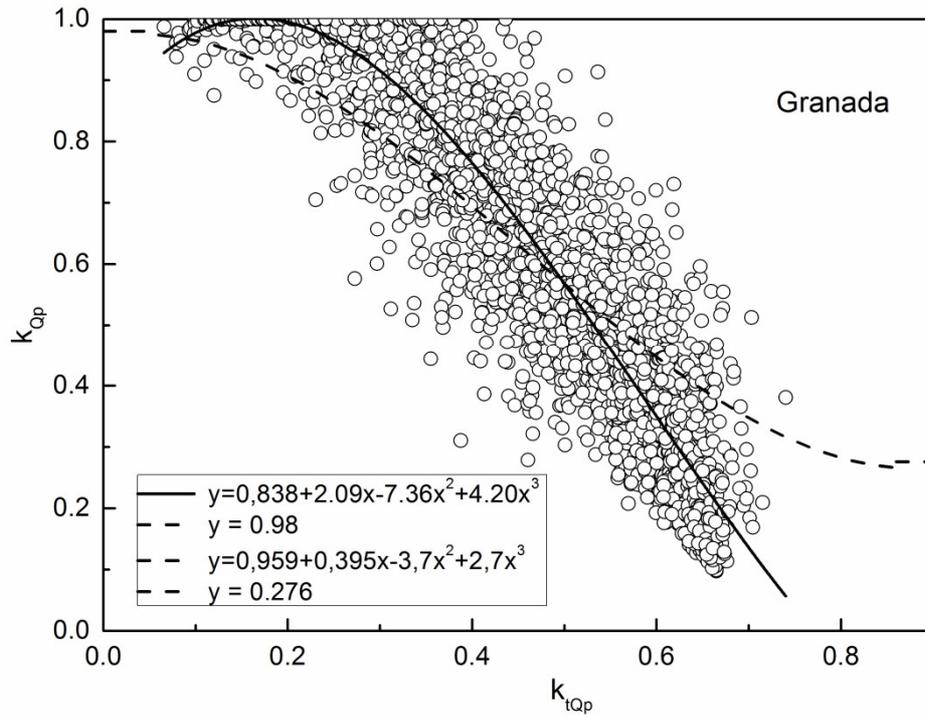
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Figure 7: Ratio of diffuse photosynthetic photon flux density to diffuse irradiance (Q_{pd}/R_d) vs. the partial vapour pressure (e). Small dots denote experimental data, black symbols represent mean values, and bars denote standard deviations for each of the intervals in e .

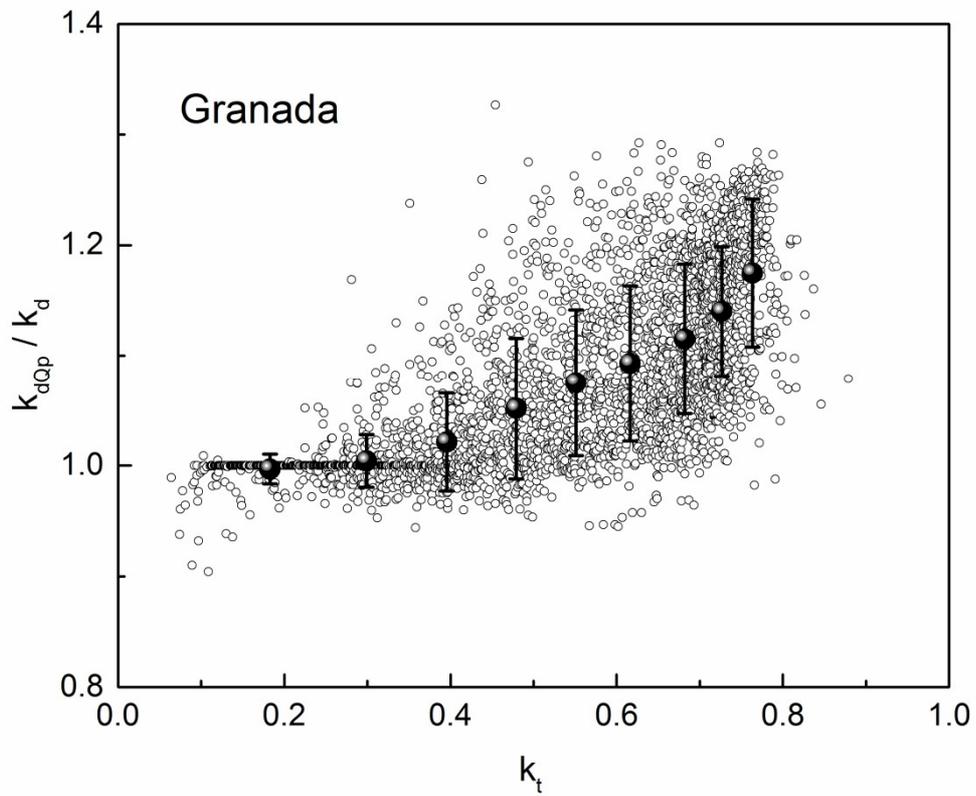
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583 Figure 8: Diffuse fraction in the visible range (k_{dQp}) vs. clearness index in the visible range
584 (k_{tQp}). Small dots denote experimental data. Dashed lines denote polynomial fitting (Jacovides
585 et al. 2007). Solid lines denote our polynomial fitting.

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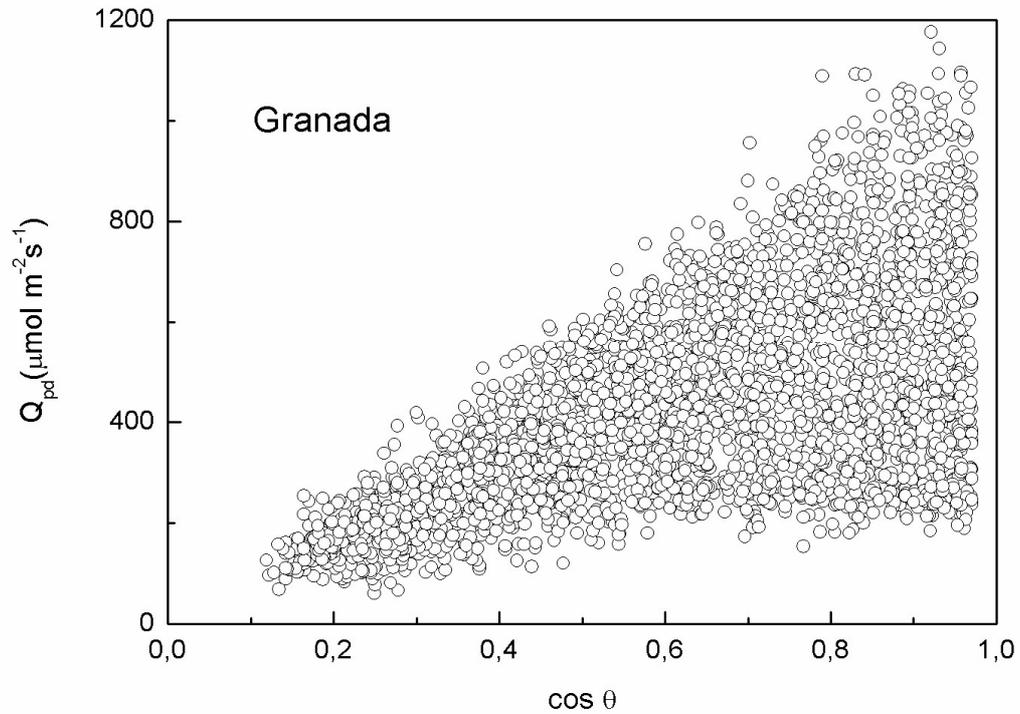


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Figure 9: Ratio of diffuse fraction in the visible range (k_{dQP}) and diffuse fraction in the shortwave broadband range (k_d) vs. clearness index (k_t). Small dots denote experimental data, black symbols represent mean values, and bars denote standard deviations for each of the interval in k_t .

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617 Figure 10: Diffuse photosynthetic photon flux density (Q_{pd}) vs. cosine of the solar zenith

618 angle ($\cos \theta$). Dots represent experimental data.

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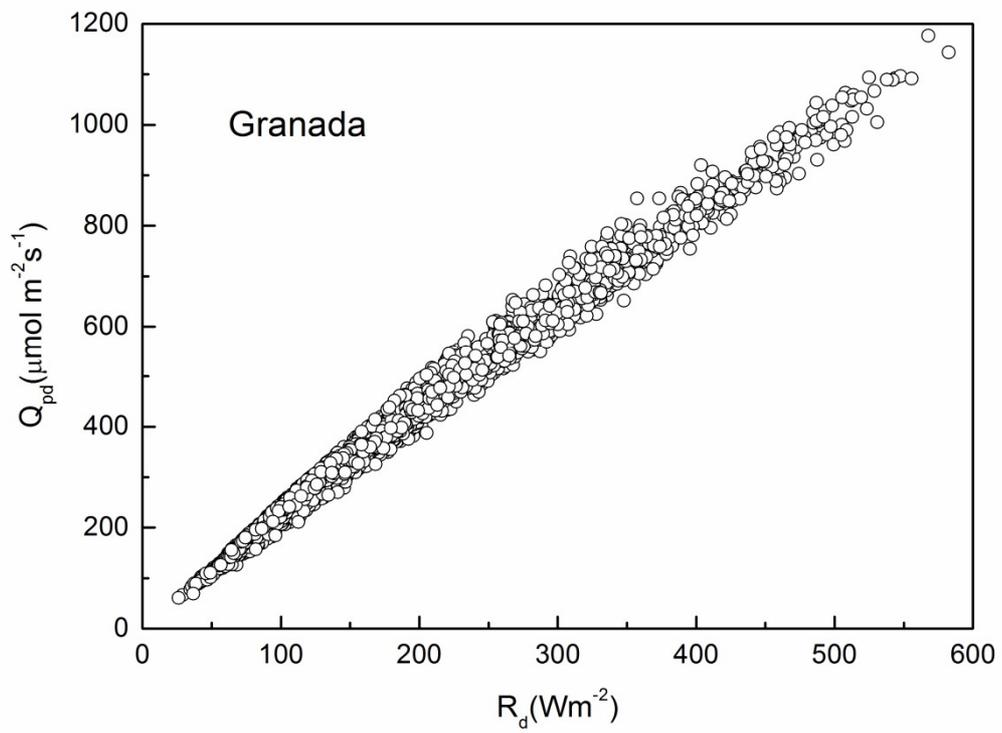
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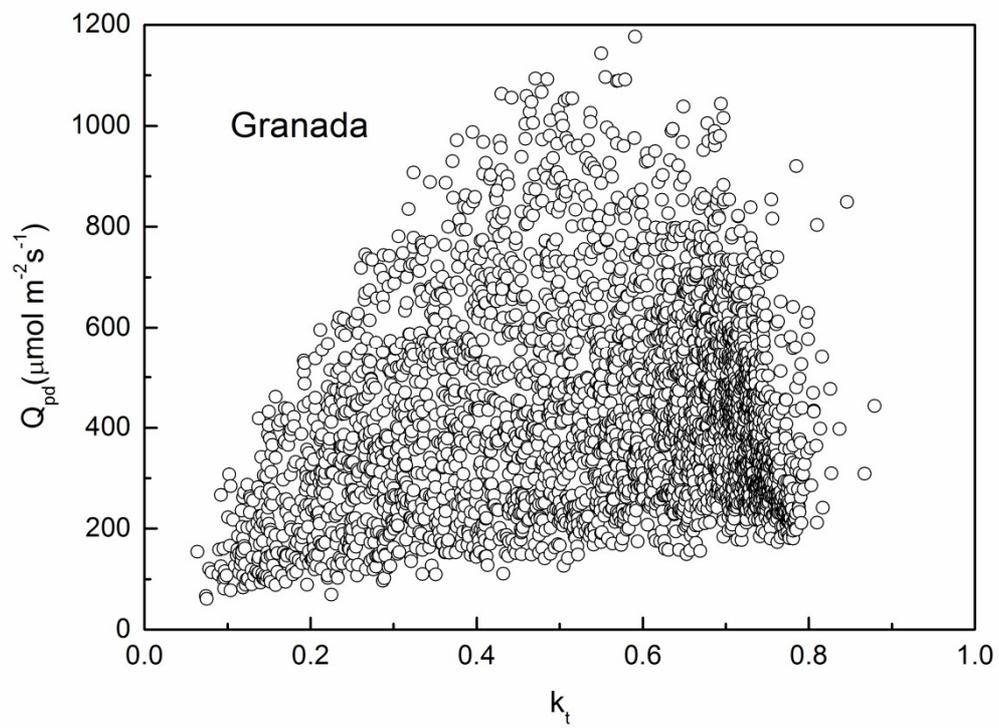
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Figure 11: Diffuse photosynthetic photon flux density (Q_{pd}) vs. R_d . Dots represent experimental data.

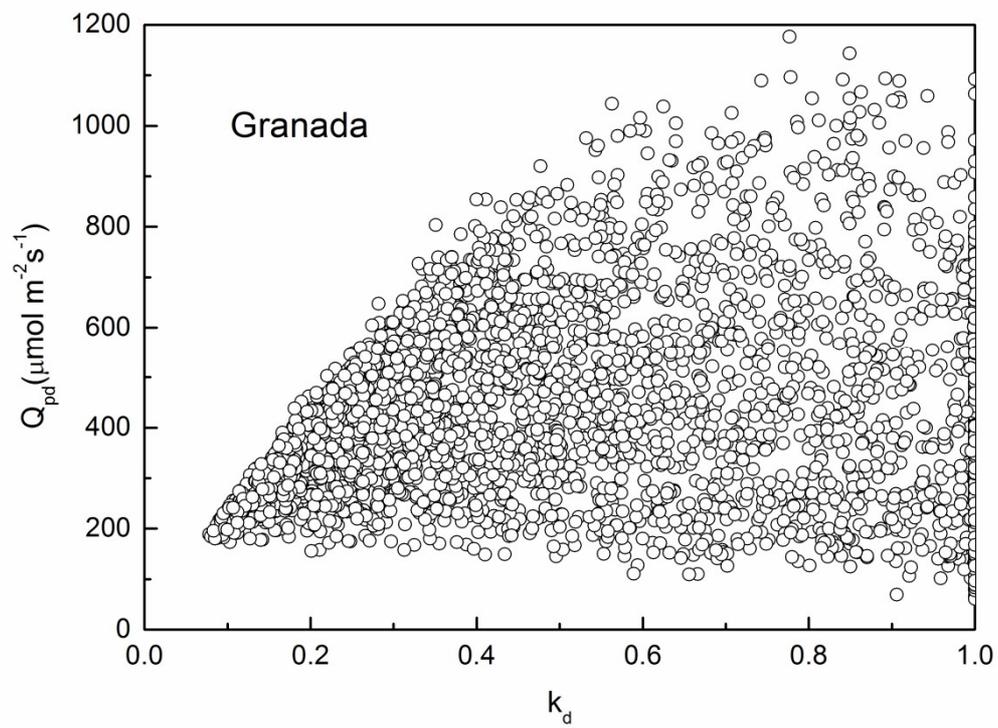
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Figure 12: Diffuse photosynthetic photon flux density (Q_{pd}) vs. k_t . Dots represent experimental data.

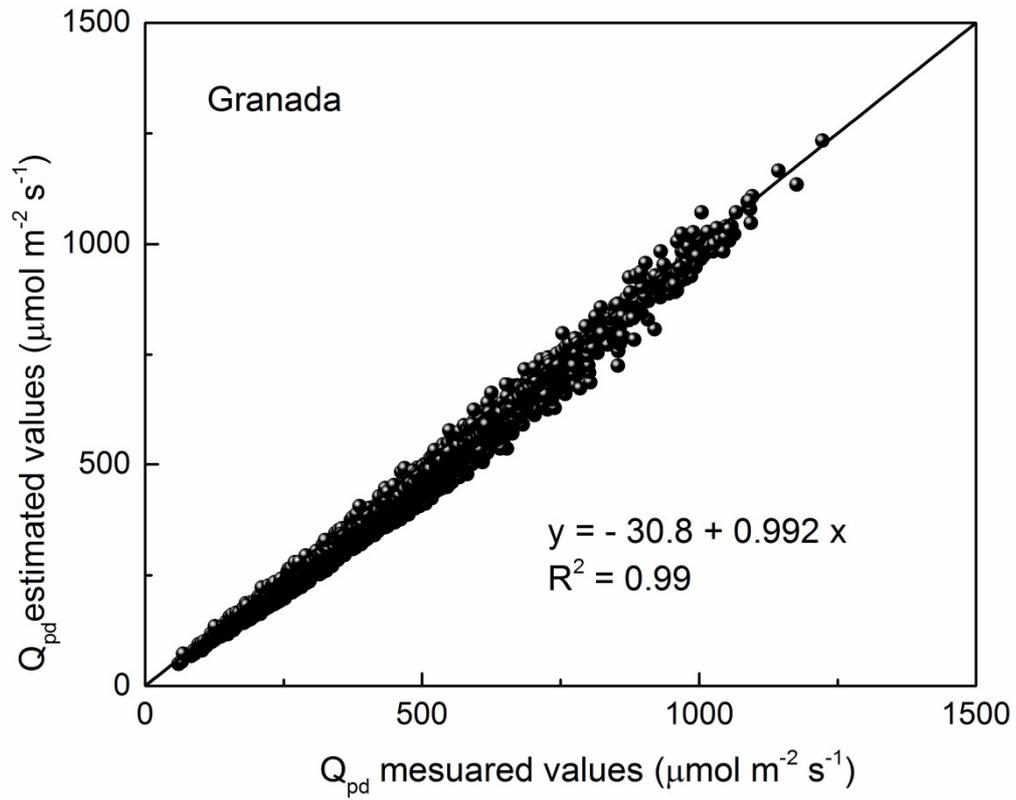
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Figure 13: Diffuse photosynthetic photon flux density (Q_{pd}) vs. k_d . Dots represent experimental data.

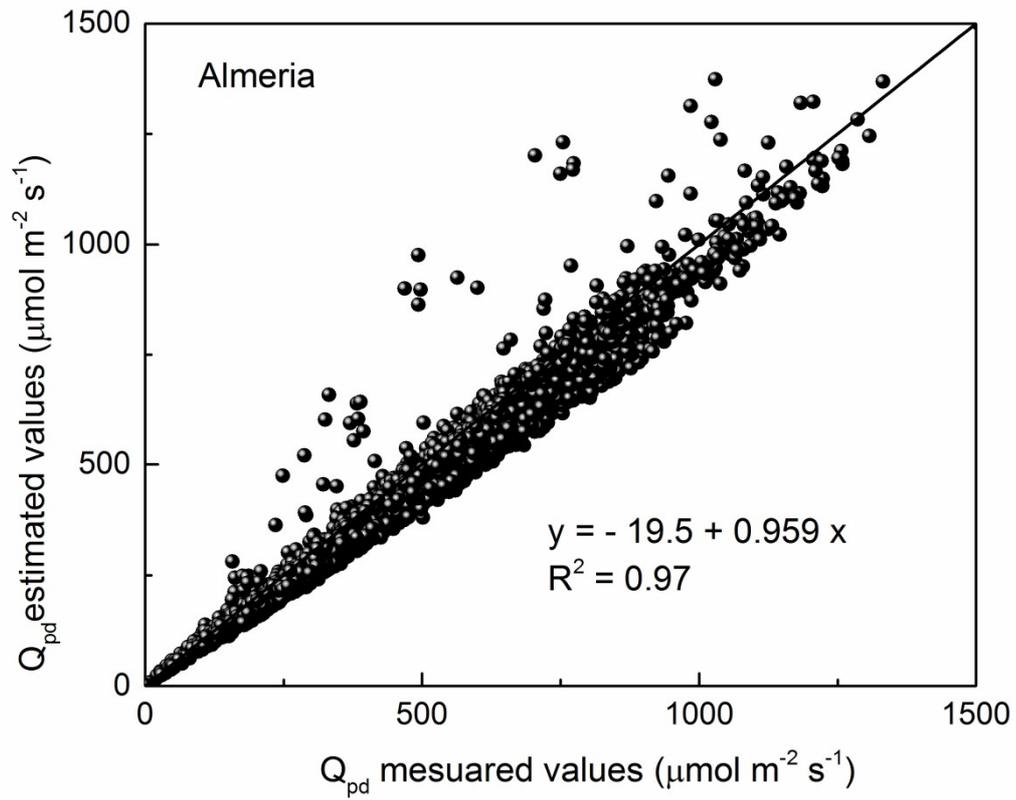
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691 Figure 14 a: Scatter plot of estimated vs. measured values of photosynthetic photon flux
692 density (Q_{pd}) at Granada site.

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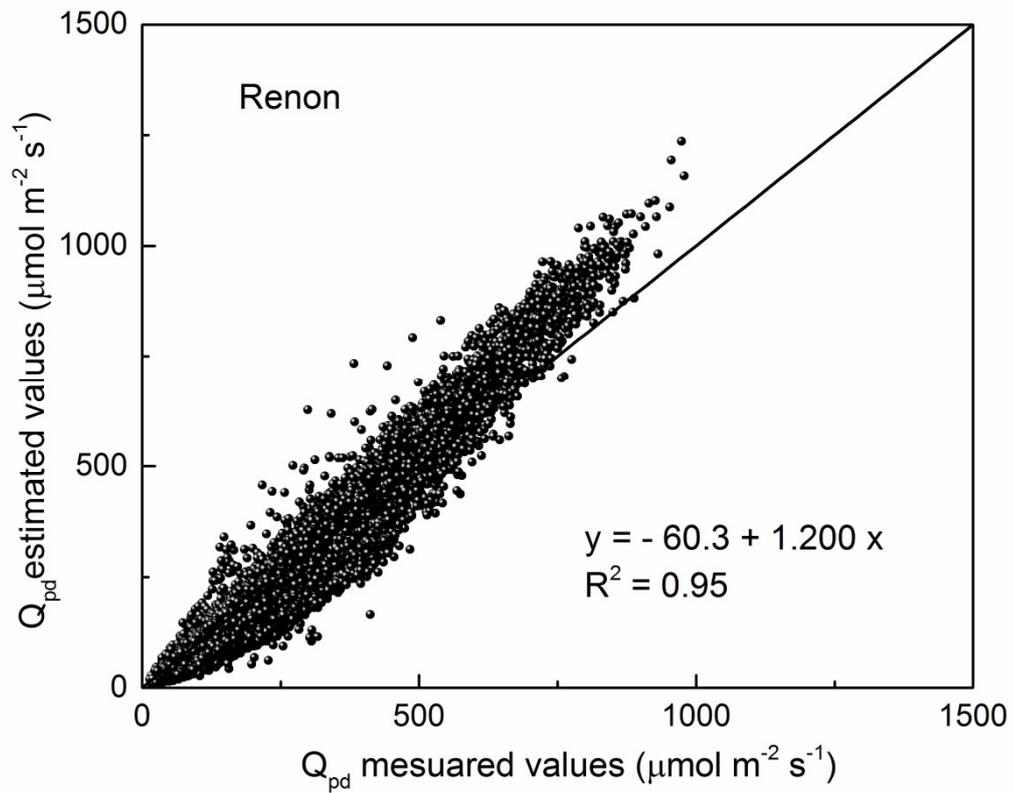
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Figure 14 b: Scatter plot of estimated vs. measured values of photosynthetic photon flux density (Q_{pd}) at Almeria site.

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728 Figure 14c: Scatter plot of estimated vs. measured values of photosynthetic photon flux
729 density (Q_{pd}) at Renon site.
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