

In Situ Observation of Biotite Dissolution at pH 1 Using Advanced Optical Microscopy

Chiara Cappelli,^{*,†} Alexander E. S. Van Driessche,[§] Jordi Cama,[‡] and F. Javier Huertas[†]

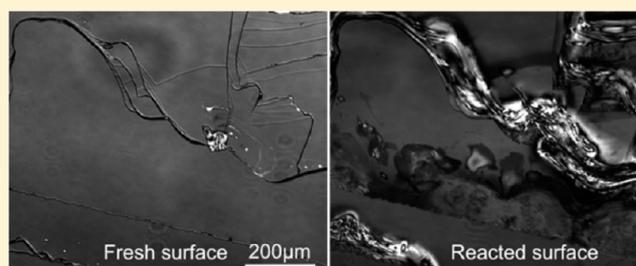
[†]Instituto Andaluz de Ciencias de la Tierra (IACT - CSIC-Universidad de Granada), Avda. de las Palmeras 4, 18100 Armilla, Granada, Spain

[‡]Institute of Environmental Assessment and Water Research (IDAEA), CSIC, Jordi Girona 18-26, 08034 Barcelona, Spain

[§]Laboratorio de Estudios Cristalograficos (IACT, CSIC-Universidad de Granada), Avd. de las Palmeras 4, 18100 Armilla, Granada, Spain

 Web-Enhanced Feature

ABSTRACT: Laser confocal microscopy with differential interference contrast microscopy (LCM-DIM) was used to study in situ the biotite (001) basal surface dissolution performing long-term flow-through experiments at pH 1 ($T = 11.5\text{--}70\text{ }^{\circ}\text{C}$). Time-lapse image sequences of large areas (up to 1 mm) of the evolving cleavage (001) surface showed that dissolution only occurs at surface edges. In addition, in contrast to an observed rapid dissolution at low steps (few layers), swelling and contraction of the edge layers occurred at high steps (many layers). An increase in temperature enhanced the surface edge dissolution from $7.5 \times 10^{-4}\text{ }\mu\text{m s}^{-1}$ at $11.5\text{ }^{\circ}\text{C}$ to $6.2 \times 10^{-2}\text{ }\mu\text{m s}^{-1}$ at $70\text{ }^{\circ}\text{C}$. The results obtained in this work demonstrate that LCM-DIM is a powerful technique to study in situ the dissolution mechanism and kinetics of phyllosilicates.



1. INTRODUCTION

Kinetics of phyllosilicate dissolution has been extensively studied over the last 20 years.¹ In particular, special attention has been paid to understand the effect that pH exerts on the dissolution mechanisms and rate. It is argued that at acidic pH clay mineral dissolution predominantly takes place at the clay mineral edge surface.² The interesting research on dissolution of kaolinite³ and biotite, using powder samples^{4,5} and single crystals (monolithes),⁶ showed that the contribution of the edge surface area to the dissolution rate is higher than the basal surface area, suggesting that the edge sites dissolve faster than sites on the basal surface and often control the overall dissolution rate under acidic conditions. After these initial studies, atomic force microscopy (AFM) has been the most frequently used technique to investigate, at the nanoscale, the reaction mechanisms of phyllosilicates dissolution. Rufe and Hochella⁷ assessed quantitatively the reactive surface area of dissolving phlogopite at acidic conditions and room temperature. Although the basal {001} surfaces showed certain reactivity, the dissolution rates normalized to the edge surface area were 2 orders of magnitude faster than rates normalized to the total surface area for this mineral. Bosbach et al.⁸ and Bickmore et al.⁹ showed that the basal surfaces of two smectite (hectorite and nontronite) crystallites appeared to be unreactive during the time scale of the experiments in acid solutions. At high temperature (100–140 °C), Aldushin et al.¹⁰ investigated the reactions of phlogopite and apophyllite with acidic solutions and observed that phlogopite dissolution occurred via nucleation of etch pits on the pristine surface,

implying that the basal surface plays an important role in the dissolution processes, at least at elevated temperatures. The apophyllite basal surface underwent a swelling process forming square-shaped hillocks. McMaster et al.¹¹ and Haward et al.¹² suggested that etch pit nucleation at the biotite basal plane controls the biotite dissolution in the presence of oxalic acid. In the absence of oxalate, however, formation of discrete etch-pits at room temperature was not observed. Pachana et al.¹³ based on SEM and AFM examinations observed stair-shaped etch pits on muscovite and biotite during hydrothermal (200 °C) alteration in acidic solutions (pH 1.1–5.7) and identified the role that nanogenic coating phases exert on the control of mineral alteration.

On the basis of these previous studies, one can assume that the likely mechanism controlling biotite dissolution at acidic pH (in the absence of oxalate) is the retreat of the sheet edge parallel to basal cleavage, resulting in an almost inert basal surface during dissolution. From the two types of biotite surface, the basal (001) and the edge (lateral) ($hk0$), the latter is dominantly involved. However, it appears that at high temperature the formation of etch pits contributes to the biotite dissolution at the basal surface,¹³ in contrast to the absence of etch pits formation at ambient temperature (10–35 °C).^{11,12}

Within this context, it still remains unclear whether the formation of etch pits at the biotite basal surface during biotite

Received: February 20, 2013

Revised: May 17, 2013

75 dissolution at acid pH could intervene in the overall reaction.
76 As formation of etch pits at the basal surface is apparently
77 temperature dependent, the study is conducted from 11.5 to 70
78 °C. In the previous studies, biotite (and other mica) dissolution
79 at acid pH was investigated from macroscopic to nanoscopic
80 perspectives based on the solution chemistry from flow-through
81 cells with powder samples and AFM observations on biotite
82 surfaces. Nonetheless, systematic, in situ, and real time
83 exploration of broad regions of the reacting biotite (001)
84 cleavage surface (in the order of tens of thousands square
85 micrometers) is still lacking.

86 To this end, the recently developed noninvasive laser
87 confocal microscopy with differential interference contrast
88 microscopy (LCM-DIM) technique¹⁴ allows in situ surface
89 exploration with vertical nanometric accuracy at different
90 temperatures.¹⁵ LCM-DIM has been successfully applied to
91 investigate in situ crystal-growth processes of gypsum crystals in
92 a broad temperature and supersaturation range.¹⁶ Although
93 unambiguous quantitative information on the vertical surface
94 topography is unavailable with LCM-DIM, horizontal quanti-
95 fication is possible, and measurements of step edge retreat with
96 time can be performed. This type of observation benefits from
97 the advantage that broad surface regions can be scanned for
98 long time periods and allows the visualization of the
99 mechanisms on surface areas with variable roughness and
100 topography. A new temperature controlled flow-through cell
101 was incorporated in the experimental microscope setup to
102 examine the alteration of the biotite (001) cleavage surface
103 during long-term experiments (up to 6 days) at pH 1 and
104 temperature range from 11.5 to 70 °C.

2. MATERIALS AND METHODS

105 **2.1. Experimental Setup.** Changes on biotite (001) cleavage
106 surface topography were monitored in situ employing LCM-DIM.¹⁴

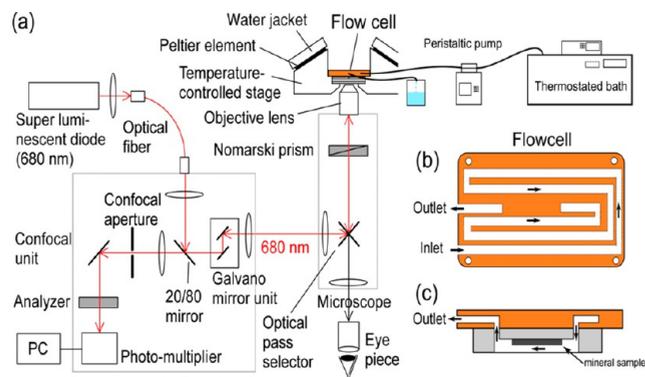


Figure 1. Schematic representation of the experimental setup: (a) the LCM-DIM equipment, (b) top view and (c) a cross-sectional view of the temperature-controlled observation flow cell.

107 This advanced optical system (Figure 1a) is a combination of two
108 microscopy techniques: a confocal system (FV300, Olympus) is
109 attached to an inverted optical microscope (IX70, Olympus), and a
110 Nomarski prism is introduced into the optical path. A partially
111 coherent super luminescent diode (Amonics Ltd., model ASLD68-
112 050-B-FA: 680 nm) is used to eliminate diffraction noise. This
113 combination provides high contrast and high resolution in the vertical
114 plane ($z \sim 1$ nm) of the crystal surface, and the x - y plane resolution is
115 ~ 1380 nm for a 10 \times objective lens (LUCplan FLN 10 \times , Olympus).¹⁵
116 The surface is scanned line-by-line deflecting the laser beam in the x - y
117 plane by galvanic mirrors. The light reflected at the crystal surface/
118 solution interface is detected by a photomultiplier tube (PMT).

A custom-made Teflon flow-through cell consisted of a rectangular
119 prism with a volume of 0.08 cm³ and a fissure on one side. The
120 observation cell was mounted inside a copper sample stage which
121 completely surrounds the cell, except for the observation area (7 mm
122 in diameter) at the bottom side (Figure 1b,c). To precisely control the
123 temperature of the sample stage, a curve-matched thermistor and two
124 Peltier elements were employed. Moreover a PR-59 (Supercool)
125 temperature controller PC-interfaced through a serial connection by
126 means of a Labview driver was used. Water jackets with circulating
127 water were placed on top of the Peltier hot sides to increase the
128 cooling capacity for set points below room temperature. Thermal
129 silicon grease was employed to minimize thermal contact resistance.
130 The precision of the temperature control is ± 0.5 °C.¹⁷

2.2. Dissolution Experiments. The biotite used in the present
132 study is from Bancroft (Ontario, Canada) and was purchased from
133 Ward's Natural Science Establishment. In principle, its composition
134 and structural formula should be similar to that reported by Turpault
135 and Trotignon (1994)⁶ who used a similar biotite sample.

In the flow-through experiments, biotite flakes with (001) cleavage
137 surfaces of $\sim 2 \times 8$ mm² and between 0.08 and 0.15 mm in thickness
138 reacted with 0.1 mol L⁻¹ HNO₃ (\sim pH 1) and 0.01 mol L⁻¹ NaNO₃
139 solutions prepared from ultrapure grade chemicals.

To observe the dissolving basal surface, the samples were fixed
141 parallel to the (001) surface on the bottom of the fissure of the reactor
142 cell by a silicone adhesive (Figure 1c). A small canal on each side of
143 the cell allowed circulation of the reacting solution at a constant flow
144 rate (0.01–0.07 mL min⁻¹), yielding a residence time from
145 approximately 1 to 8 min. The Teflon reactor was carefully sealed
146 with a cover glass glued with high vacuum grease (Dow Corning). The
147 experiments were conducted at pH ~ 1 over a wide temperature range
148 (11.5–70 °C). To increase the thermal stability in the experiments at
149 $T \geq 40$ °C, the reacting solutions were preheated in a thermostatted
150 bath (Figure 1a). Additionally, the input solutions were degasified with
151 a vacuum pump for ~ 2 h to reduce bubble formation during the high
152 temperature runs ($T \geq 40$ °C).

Experiments lasted from 2 h to 6 days according to the experimental
154 conditions (i.e., the lower the temperature, the longer the experiment
155 span due to the slower biotite dissolution kinetics). Images of the
156 (001) cleavage surface were taken every 20 s to 15 min with a time of
157 capture of 9.6 s. Replicas were performed for each experiment. Retreat
158 rate (R_r) at a given location of the surface was computed from the
159 surface horizontal retreat: $R_r = \Delta L / (t_2 - t_1)$, where ΔL is the change
160 in the surface horizontal length as measured over the period between
161 t_1 and t_2 at a given location.

3. RESULTS AND DISCUSSION

3.1. Biotite Dissolution Mechanism. The LCM-DIM
163 images show that the freshly cleaved biotite (001) surface is a
164 flat basal surface formed by terraces of variable extension and
165 height (Figure 2a). The limit between terraces is marked by
166 lines; high steps display darker outlines (A-A', B-B', C-C',
167 and D-D' in Figure 2a) than lower ones (e-e', f-f', and g-g'
168 in Figure 2a). The visible limits of the (001) cleavage planes do
169 not show preferential crystallographic directions. Surface
170 changes are readily visible as the biotite surface interacts with
171 the pH 1 solution. It is observed that the retreat of the higher
172 steps occurs rather irregularly along the edge as dissolution
173 starts. For instance, the evolution of a high step is depicted in
174 the images of Figure 2 (experiment B50-1-2, Table 1): the
175 terraces of the pristine surface are limited by high (AA', BB',
176 CC', and DD') and low (ee', ff', and gg') steps (Figure 2a).
177 After 53 min at 50 °C (Figure 2b), alteration of the high steps
178 is only visible at limited sections, resulting in a localized step
179 alteration and stationary step position. In the altered sections,
180 the upper layers appear "overexposed" (black arrows) due to
181 swelling and consecutive breaking that induce the process of
182 layer peeling. After ~ 6 h of reaction only the step edges AA'
183

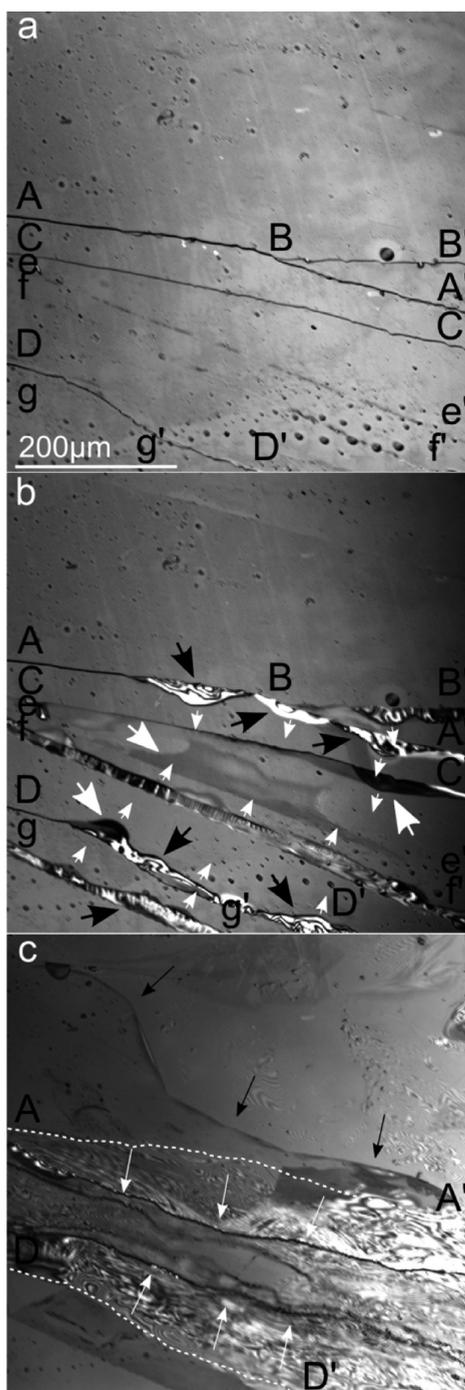


Figure 2. LCM-DIM images of biotite (001) cleavage surface: pristine surface (a), same surface reacted for 53 min (b) and ~6 h (c) at pH 1 and 50 °C. Capital letters and lower case letters indicate high and low step edges, respectively. Dissolution of higher steps at early stages of the reaction moves nonhomogeneously inward from the edges. The spherical structures in (a) are air bubbles adsorbed on the surface. Large white and black arrows in (b) indicate swelling and peeling layers, respectively, while small white arrows indicate dissolution directions. During the first hour of the reaction, the steps do not retreat, suggesting some dissolution of the upper layers. As the reaction continues, most of the steps disappear to leave altered areas. Two uniform reaction fronts remain after 6 h (AA' and DD' in c). Dotted lines in (c) indicate the AA' and DD' step positions after 53 min; white arrows show dissolution directions, and black arrows indicate dissolution of a new low step leaving fresh surface.

Table 1. Experimental Conditions

experiment	T (°C)	HNO_3 (M)	NaNO_3 (M)	flow rate (mL min^{-1})	duration (h)
B11.5-1-1	11.5	0.1	0.01	0.072	141.8
B11.5-1-2	11.5	0.1	0.01	0.072	89.8
B25-1-3	25	0.1	0.01	0.072	31.0
B25-1-4	25	0.1	0.01	0.072	18.3
B40-1-1	40	0.1	0.01	0.073	4.3
B40-1-1c	40	0.1	0.01	0.073	3.7
B50-1-1	50	0.1	0.01	0.01	97.6
B50-1-2	50	0.1	0.01	0.01	25.2
B70-1-3	70	0.1	0.01	0.068	4.8
B70-1-3b	70	0.1	0.01	0.068	2.1
B70-1-5b	70	0.1	0.01	0.072	2.0

and DD' are visible, remaining stable along the main part of its length (Figure 2c). Yet, the altered areas spread following uniform reaction fronts and originated a widely degraded region. The other high and lower steps disappear either after complete dissolution of a terrace or after coalescence into the main steps.

In general, the retreat of the step edges follows an approximately normal direction to the terrace limits, irrespective of the experimental temperature (Figure 3). Figure 3a shows the altered biotite cleavage surface after 5 h of dissolution at 70 °C (experiment B70-1-1, Table 1) with the presence of several steps of relatively low height. After 30 min of reaction, the steps have retreated and left a fresh new basal biotite surface (Figure 3b). The advance of the reaction front occurs approximately normal to the edge (indicated by white arrows). Nevertheless, inspection of the position of the reaction fronts (e.g., f_1 and f_2 in Figure 3) at discrete consequent intervals shows that they move at different velocities (see the change in the f_1 – f_2 distance with time), yielding a dissolution-front fluctuation likely caused by local surface defects (e.g., chemical inhomogeneities). In Figure 3 likewise, it is observed that upper layers move faster than lower ones and promote the formation of new surface edges (see black arrows in Figure 3c). This behavior mostly indicates step–step interaction.

Figure 4 shows a sequence of images obtained during an experiment performed at 11.5 °C for 6 days (experiment B11.5-1-1, Table 1). After 21 h dissolution takes mainly place at discrete sections along the steps (Figure 4a,b). Swelling and peeling also occur, and the retreat of lower steps leaves altered areas (Figure 4c). Interestingly, dissolution at the basal plane is visible (Figure 4c,d). Swelling of basal surface precedes the contraction and breaking of the layers (Figure 4c). The origin of this behavior is unclear and could depend on the presence of nonvisible defects in the basal plane. Figure 5 (experiment B25-1-4, Table 1) shows that the existence of small holes (indicated by black arrow), likely caused by the previous cleaving of the basal surface, could act as surface features that significantly promote dissolution by generating fronts spreading from these structures. Therefore, although this phenomenon can contribute to the overall dissolution rate, the dissolution of basal planes starting from these preexistent defects cannot be considered as etch pits nucleation and growth. This particular mechanism may be analogous to the defect-controlled mechanism observed in chlorite by Brandt et al. (2003).¹⁸

Mica dissolution is thought to be dominated by an edge attack mechanism that produces selective leaching of interlayer (K^+ , Na^+) and octahedral (Mg^{2+} , Fe^{2+} , Fe^{3+} , Al^{3+}) cations

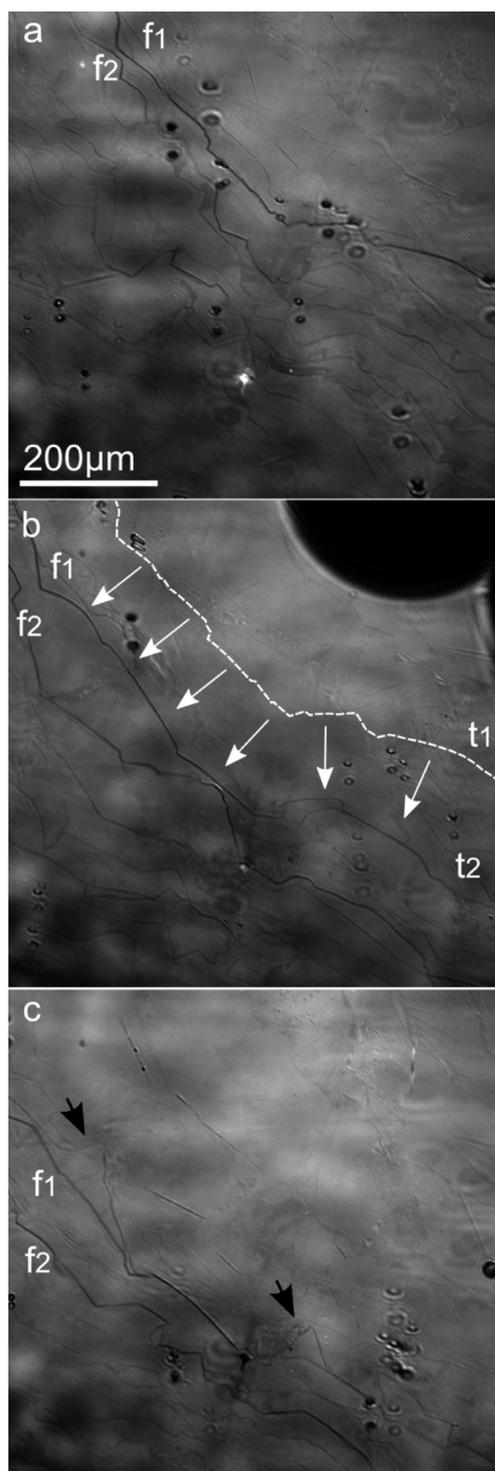


Figure 3. LCM-DIM images of reacted biotite (001) cleavage surface after ~5 h (t_1) (a), 5 h and 30 min (t_2) (b), and 5 h and 46 min (c) at pH 1 and 70 °C. In (b) the dotted line indicates the previous position of the main edge in (a). White arrows indicate retreat direction. Black arrows in (c) indicate areas where upper layers dissolved faster than the lower ones. The different distance between consecutive dissolution fronts for different time reactions (f_1 and f_2 in a, b, and c) is indicative of retreat fluctuation. The small dark spherical structures and the large dark area in the upper right corner in (b) are air bubbles; the pale elliptical structures are dust particles derived from the objective lens. A time-lapse movie of this process in AVI format is available in the HTML version.

through the $\{hk0\}$ surfaces and forms an altered silica-enriched rim.^{4,6,7,10,19} Fast exchange of interlayer potassium by protons and loss of octahedral cations lead to biotite layers curling and peeling off.⁶ Atomic force microscopy (AFM) revealed a change in the phlogopite layer volume induced by alteration in HCl solutions at pH 2 that was explained by depolymerization and repolymerization reactions occurring during leaching layer formation.⁷ After the initial leaching of octahedral cations, the hydrolysis of Si–O–Si and Si–O–Al in the tetrahedral sheet allows the layer to expand. The consecutive cross-linking of neighboring Si–OH groups to reform Si–O–Si (repolymerization) might cause layer contraction by the expulsion of water. Aldushin et al.¹⁰ observed that the dissolution of phlogopite basal surface at 100–135 °C was accompanied by an alteration of the topmost layers. Initially the layer expands and then corrugates and decomposes. They interpreted the mechanism as Rufe and Hochella⁷ but did not observe contraction of the depolymerized layers likely due to the higher alteration temperature.

The current observation is substantially consistent with these previous studies, and the topographical changes in biotite can be explained by initial leaching of interlayered and octahedral cations, followed by depolymerization and polymerization of tetrahedral sheet. These reaction stages lead to swelling and peeling of the uppermost layers in higher steps and formation of a degraded, low reflective surface, associated with the residual silica-rich layer and biotite incongruent dissolution. The retreat of low steps that leave a fresh basal surface can be interpreted as congruent dissolution.

The above results show that the dissolution mechanism of the biotite (001) cleavage surface is the same for the entire temperature range of this study and that formation of etch pits does not occur. This is consistent with the study by Rufe and Hochella⁷ who etched the phlogopite surface with HF to produce an edge surface and did not observe the nucleation of etch pits on the $\{001\}$ surfaces. On the other hand, McMaster et al.¹¹ show that dissolution of the biotite (001) surface in oxalic acid is dominated by the formation of etch pits, but in the absence of oxalate the formation of discrete etch pits is not observed.

3.2. Biotite Dissolution Rate. Steps of different thickness retreat at different rates. As expected, higher steps retreat more slowly than the lower ones. For example, in the experiment conducted at 50 °C and pH 1 (B50-1-2, Figure 2), the retreat rate of the higher steps (e.g., AA' and DD') is $5.4 \times 10^{-3} \mu\text{m s}^{-1}$, whereas that of the lower steps, that dissolves leaving fresh surface (see black arrows) is $2.6 \times 10^{-2} \mu\text{m s}^{-1}$. On the basis of the LCM-DIM measurements of the horizontal low steps retreat variation with time, dissolution rates of the biotite ($hk0$) surface were obtained at different temperatures.

In fact, the low-steps variation with time was considered to be optimal to quantify the retreat rates because they actually correspond to the dissolution of a few layers.¹⁰ Hence, the horizontal retreat rates (R_r in $\mu\text{m s}^{-1}$) were obtained with an estimated error of 25% (standard deviation calculated from the step retreat values of replicated experiments) (Table 2). The Arrhenius plot derived from the rates with temperature yields an apparent activation energy (E_{app}) of 61.7 kJ mol^{-1} (Figure 6). This value suggests that biotite dissolution at pH 1 is surface controlled.²⁰

It is not easy to compare the retreat rates obtained in this study with the rates reported in two previous studies that dealt with the dissolution of the biotite basal surface at acid pH.^{12,13}

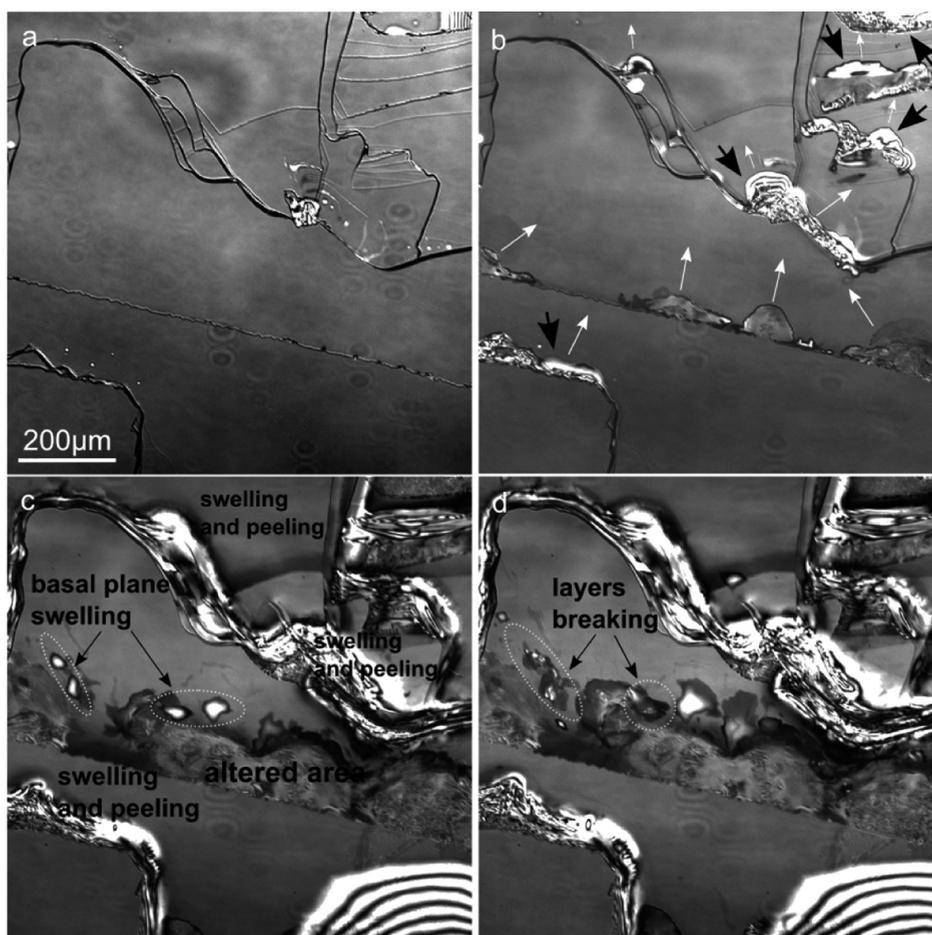


Figure 4. LCM-DIM images of biotite (001) cleavage surface: pristine surface (a) and after reaction at pH 1 and 11.5 °C for 21 h (b), ~5 days (c), and ~6 days (d). Note the presence of high step edges (dark outlines). Dissolution started irregularly along the edges (white arrows in (b) indicate dissolution direction). Large black arrows indicate overlit areas (regions with PMT saturation) where roughening and peeling of layers occur. A time-lapse movie of the last 42 h of the experiment is available.

294 The main difficulty arises from the fact that the basal surface
 295 dissolution was controlled by the etch pit formation in the
 296 presence of oxalate and at high temperature (200 °C).
 297 However, the E_{app} value inferred from the step-retreat rates
 298 falls in the E_{app} range associated with etch pit growth into the
 299 $\{hk0\}$ surfaces.

4. CONCLUSIONS

300 This study demonstrates the following:

301 (1) LCM-DIM is a powerful noninvasive optical technique to
 302 investigate in situ the topographic changes on dissolving
 303 mineral surface areas over a broad temperature range. Hence,
 304 this technique is useful to unravel the major dissolution
 305 mechanisms of phyllosilicates.

306 (2) On the basis of the time-lapse high-resolution image
 307 sequences, biotite dissolution is mainly controlled by surface-
 308 edge retreat. Dissolution from existing defects on the basal
 309 plane is not a proper biotite dissolution mechanism. Formation
 310 of basal etch pits did not occur during the reaction time of the
 311 experiments (up to 6 days) in the temperature range of this
 312 study (11.5–70 °C).

313 (3) The apparent activation energy value of 61.5 kJ mol⁻¹
 314 suggests that biotite dissolution at pH 1 is surface controlled.

■ ASSOCIATED CONTENT

315

W Web-Enhanced Feature

316

Time-lapse movies of Figures 3 and 4 in AVI format.

317

■ AUTHOR INFORMATION

318

Corresponding Author

319

*Phone: +34 958 23 00 00 (ext. 190123). Fax: +34 958 55 26
 20. E-mail: chiaracappelli@ugr.es.

320

321

Notes

322

The authors declare no competing financial interest.

323

■ ACKNOWLEDGMENTS

324

Financial support was obtained from Projects CGL2008-01652
 325 and CGL2011-22567 (Plan Nacional I+D and EU). C.C. 326
 benefitted from a FPU grant (MEC), and A.E.S.V.D. is grateful
 327 for the support by the Consolider-Ingenio 2010 project
 328 “Factoría Española de Crystalización”. Luis David Patiño
 329 López is acknowledged for his assistance during the
 330 experimental work. The comments and suggestions of Dr.
 331 Felix Brandt and two anonymous reviewers have largely
 332 improved the quality of the manuscript and are gratefully
 333 acknowledged.
 334

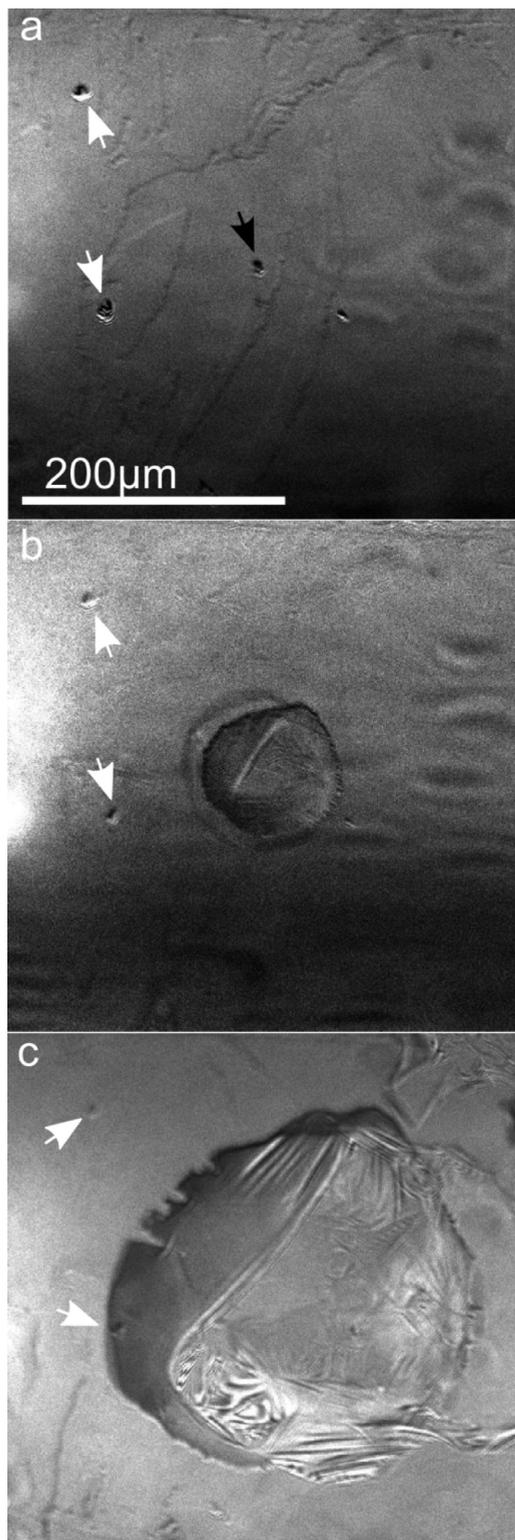


Figure 5. LCM-DIM images of biotite (001) cleavage surface: pristine surface with a visible defect indicated by the black arrow (a); same surfaces after ~7.5 h (b) and after 23 h (c) of reaction at pH 1 and 25 °C. Dissolution proceeded from the edge surface of the defect (hole). White arrows indicate holes that did not react, most likely because they were below the upper layer that interacted with solution.

Table 2. Horizontal-Retreat Rates (R_r) Calculated from the Horizontal Surface Retreat of the Lower Steps at the Respective Temperature

T (°C)	R_r ($\mu\text{m s}^{-1}$)
11.5	7.5×10^{-4}
25	3.5×10^{-3}
40	1.9×10^{-2}
50	2.6×10^{-2}
70	6.2×10^{-2}

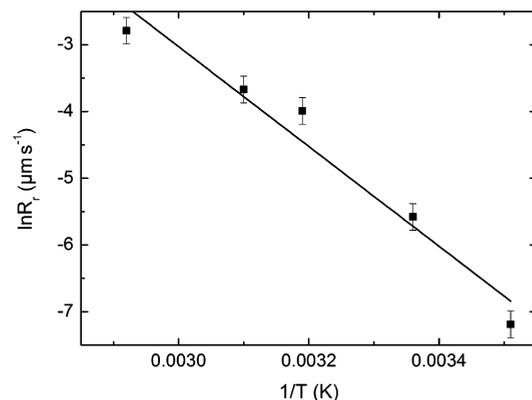


Figure 6. Arrhenius plot calculated from measurement of horizontal retreat rates at pH 1 and 11.5 to 70 °C ($R^2 = 0.96$).

REFERENCES

- Brantley, S. L. In *Kinetics of Water-Rock Interaction*; Brantley, S. L., Kubicki, J. D., White, A. F., Eds.; Springer: New York, 2008; pp 151–210. 336
- Brantley, S. L.; Conrad, C. F. In *Kinetics of Water-Rock Interaction*; Brantley, S. L., Kubicki, J. D., White, A. F., Eds.; Springer: New York, 2008; pp 1–38. 339
- Ganor, J.; Mogollón, J. L.; Lasaga, A. C. *Geochim. Cosmochim. Acta* **1995**, *59*, 1037–1052. 341
- Acker, J. G.; Bricker, O. P. *Geochim. Cosmochim. Acta* **1992**, *56*, 3073–3092. 342
- Kalinowski, B. E.; Schweda, P. *Geochim. Cosmochim. Acta* **1996**, *60*, 367–385. 343
- Turpault, M. P.; Trotignon, L. *Geochim. Cosmochim. Acta* **1994**, *58*, 2761–2775. 344
- Rufe, E.; Hochella, M. F., Jr. *Science* **1999**, *285*, 874–876. 345
- Bosbach, D.; Charlet, L.; Bickmore, B.; Hochella, M. F., Jr. *Am. Mineral.* **2000**, *85*, 1209–1216. 346
- Bickmore, B. R.; Bosbach, D.; Hochella, M. F., Jr.; Charlet, L.; Rufe, E. *Am. Mineral.* **2001**, *86*, 411–423. 347
- Aldushin, K.; Jordan, G.; Schmahl, W. W. *Geochim. Cosmochim. Acta* **2006**, *70*, 4380–4391. 348
- McMaster, T. J.; Smits, M. M.; Haward, S. J.; Leake, J. R.; Banwart, S.; Ragnarsdóttir, K. V. *Mineral. Mag.* **2008**, *72*, 115–120. 349
- Haward, S. J.; Smits, M. M.; Ragnarsdóttir, K. V.; Leake, J. R.; Banwart, S. A.; McMaster, T. J. *Geochim. Cosmochim. Acta* **2011**, *75*, 6870–6881. 350
- Pachana, K.; Zuddas, P.; Censi, P. *Appl. Geochem.* **2012**, *27*, 1738–1744. 351
- Sazaki, G.; Matsui, T.; Tsukamoto, K.; Usami, N.; Ujihara, T.; Fujiwara, K.; Nakajima, K. *J. Cryst. Growth* **2004**, *262*, 536–542. 352
- Van Driessche, A. E. S.; Otálora, F.; Sazaki, G.; Sleutel, M.; Tsukamoto, K.; Gavira, J. A. *Cryst. Growth Des.* **2008**, *8*, 4316–4323. 353
- Van Driessche, A. E. S.; García-Ruiz, J. M.; Delgado-Lpez, J. M.; Sazaki, G. *Cryst. Growth Des.* **2010**, *10*, 3909–3916. 354
- Van Driessche, A. E. S.; Gavira, J. A.; Patiño Lopez, L. D.; Otalora, F. *J. Cryst. Growth* **2009**, *311*, 3479–3484. 355

- 372 (18) Brandt, F.; Bosbach, D.; Krawczyk-Bärsch, E.; Arnold, T.;
373 Bernhard, G. *Geochim. Cosmochim. Acta* **2003**, *67*, 1451–1461.
- 374 (19) Kaviratna, H.; Pinnavaia, T. J. *Clays Clay Miner.* **1994**, *42*, 717–
375 723.
- 376 (20) Lasaga, A. C. *Kinetic Theory in the Earth Sciences*; Princeton
377 University Press: New Jersey, 1998.