
A new conventional regression model to estimate hourly photosynthetic photon flux density under all sky conditions

I. Foyo-Moreno,^{a*} I. Alados^b and L. Alados-Arboledas^a

^a *Departamento de Física Aplicada, Universidad de Granada, Spain*

^b *Departamento de Física Aplicada II, Universidad de Málaga, Spain*

ABSTRACT: In this work, we propose a new and simple empirical model to estimate photosynthetic photon flux density under all sky conditions, developed using experimental measurements carried out at Granada, an urban site in Southeastern Spain during 2 recent years (2014–2015). The model uses the solar zenith angle and clearness index as input parameters, and thus needs only global irradiance measurements usually registered in most radiometric networks. Five stations located in the Northern and Southern Hemispheres with different climatological characteristics at Europe, Asia and America (Spain, Japan and Argentina) were used to validate the model. The model provides satisfactory results, with low mean bias error (MBE) for all stations, particularly MBE, being less than 1% in absolute values in three stations and root mean square error below 6% for all stations except one with 6.1%. These results show better accuracy in comparison with other earlier empirical models and suggest the effectiveness of the model by its general applicability.

KEY WORDS photosynthetic photon flux density; global irradiance; solar zenith angle; clearness index

Received 15 October 2016; Revised 17 February 2017; Accepted 21 February 2017

2
3
4
5
6
7
8
9
10
11
12
13
14
15
16
17
18
19
20
21
22
23
24
25
26

A new conventional regression model to estimate hourly photosynthetic photon flux density under all sky conditions

I. Foyo-Moreno, (1), I. Alados (2) and L. Alados-Arboledas (1)
(1) Dpto de Física Aplicada, Universidad de Granada, Granada, Spain.
(2) Dpto de Física Aplicada II, Universidad de Málaga, Málaga, Spain.

Corresponding author:

I. Foyo-Moreno

Departamento de Física Aplicada

Facultad de Ciencias

Universidad de Granada

18071, Granada

Spain.

Phone: 34 58 240022

FAX: 34 58 243214

E-mail: ifoyo@ugr.es

27 **ABSTRACT**

28

29 In this work we propose a new and simple empirical model to estimate photosynthetic photon
30 flux density under all sky conditions, developed using experimental measurements carried out
31 at Granada, an urban site in Southeastern Spain during two recent years (2014-2015). The
32 model uses the solar zenith angle and clearness index as input parameters, and thus needs only
33 global irradiance measurements usually registered in most radiometric networks. Five stations
34 located in the northern and southern hemisphere with different climatological characteristics
35 at Europe, Asia and America (Spain, Japan and Argentina) were used to validate the model.
36 The model provides satisfactory results, giving low mean bias error for all stations,
37 particularly Mean Bias Error, MBE, being less than 1% in absolute values in three stations
38 and Root Mean Square Error, RMSE, below 6% for all stations except one with 6.1%. These
39 results show better accuracy in comparison to other earlier empirical models and suggest the
40 effectiveness of the model by its general applicability.

41

42 **KEYWORDS.**

43 Photosynthetic photon flux density, global irradiance, solar zenith angle, clearness index.

44

45 **1. INTRODUCTION.**

46

47 Photosynthetically active radiation is defined as the visible portion of global irradiance
48 (R_s) that is utilized by plant biochemical processes in photosynthesis to convert light energy
49 into biomass (Udo and Aro, 1999; Jacovides et al., 2004; Tang et al., 2013). This radiation
50 lies between 400 and 700 nm and covers both photon and energy terms. Consequently it is

51 expressed in energy units (Wm^{-2}) or photon units ($\mu\text{mol m}^{-2} \text{s}^{-1}$) ($1 \mu\text{mol photons m}^{-2} \text{s}^{-1}$
52 $=6.022 \times 10^{17}$ photons $\text{m}^{-2} \text{s}^{-1}$). Thus, photosynthetic photon flux density, Q_p , is defined as the
53 photon flux density, that is, the number of photons in the 400-700 nm waveband incident per
54 unit time on a unit surface.

55 Q_p is very important in comprehensive studies of radiation climate, remote sensing of
56 vegetation, radiation regimes of plant canopy and photosynthesis, playing important roles in
57 agriculture, atmospheric physics, forestry, ecology, energy management and photon science
58 (Cao et al., 2005). However, reliable measurements of Q_p are seldom measured on a routinely
59 basis around the world (Tsubo and Walker, 2005; Ge et al., 2011, Tang et al., 2013). In fact,
60 these measurements are only taken over a few experimental networks, such as FLUXNET
61 (Baldocchi et al., 2001), Baseline Surface Radiation Network (BSRN) (Ohmura et al., 1998),
62 National Oceanic and Atmospheric Administration Surface Radiation Budget Network
63 (SURFRAD) (Augustine et al., 2000), or Chinese Ecosystem Research Network (CERN) (Yu
64 et al., 2006).

65 Q_p is strongly affected by the presence of clouds, which are responsible for scattering
66 processes that affect more markedly the shorter wavelengths in the solar spectrum, which
67 include the photosynthetically active spectral range. However, the use of a model including
68 direct information about cloudiness presents the problem that cloudiness observations are
69 usually instantaneous and most irradiance measurements are averaged over intervals between
70 5 min and 1 hour. Thus, if sky conditions change considerably during the averaging time, the
71 matching of averaged radiation measurements with instantaneous cloud observations may
72 produce confusing results (Calbó et al., 2005). Q_p is also affected by atmospheric aerosols,
73 especially under cloudless conditions (Alados et al., 2000), hence it is necessary to model the
74 influence produced by both factors: clouds and aerosols.

75 Different methodologies have arisen for the analysis of clouds effects on this
76 radiometric flux, either radiative transfer models or empirical approaches. The first method
77 implies the application of Mie theory and requires adequate input information regarding cloud
78 optical thickness and drop size distributions at high temporal and spatial resolution; such
79 information is limited to specific sites and campaigns. Furthermore, these spectral models
80 require information concerning aerosol optical properties or an appropriate aerosol model,
81 which is not easily accessible in most radiometric stations. The second method estimates Q_p
82 from parameters usually measured at most radiometric stations, such as meteorological
83 information and measurements of global irradiance (R_s).

84 For these reasons, one widely adopted method is to model Q_p from R_s assuming the
85 ratio between both variables to be constant. However, this ratio presents different values
86 depending on the study area and season (Alados et al., 1996), local time and weather
87 conditions (Akitsu et al., 2015). Different empirical models that use appropriate input
88 parameters for the description of sky conditions can be found in the bibliography. These
89 parameters are indices such as the sky clearness ($\varepsilon = (R_d + R_b)/R_d$), the sky brightness
90 ($\Delta = R_d/R_{s0} \cos\theta$), or the clearness index ($k_t = R_s/R_{s0}$), where R_d is diffuse irradiance, R_b is
91 direct normal irradiance and R_{s0} is the extraterrestrial global irradiance (Pérez et al., 1990;
92 Alados et al., 1996). For example, Tsubo and Walker (2005) proposed a model to estimate Q_p
93 with k_t . These empirical models include several meteorological variables such as the
94 dewpoint temperature, or/and water vapour pressure to consider the absorption of radiation by
95 water vapour present in the atmosphere (Alados et al., 1996; Ge et al., 2011, Aguiar et al.,
96 2012). But, in practice these approaches include inherent systematic errors that propagate
97 from their reference (calibration) data and therefore these models are not well established
98 (Akitsu et al., 2015). Moreover, these studies have been mostly conducted at sites in the

99 northern hemisphere and few direct measurements have been carried out in the southern
100 hemisphere (Aguilar et al., 2012).

101 In order to find a simple empirical model to estimate Q_p from available measurements
102 in most radiometric stations, we have analysed the dependences of the ratio Q_p/R_s and Q_p on
103 different parameters. These parameters are: solar position through the solar zenith angle (θ),
104 global irradiance (R_s), and sky condition. Sky condition includes the effect of clouds and
105 aerosols through the clearness index (k_t). To consider the effects of water vapour absorption
106 in this spectral range, we also have analysed the dependence on meteorological parameters
107 such as water vapour pressure (e) or dewpoint temperature (T_d).

108 We propose a simple model which only needs global irradiance measurements as input
109 variable. This model presents the advantage that both variables (Q_p and R_s) can be obtained
110 with the same time interval and R_s is available in most radiometric stations. The model has
111 been evaluated against a set of independent data at various sites with different climatological
112 characteristic located in the northern and southern hemisphere.

113

114 **2. EXPERIMENTAL SITE AND MEASUREMENTS**

115

116 Ground-based data from 2014 to 2015 at 1 min intervals were recorded at the
117 radiometric station located on the rooftop of the Andalusian Institute for Earth System
118 Research in Granada (IISTA-CEAMA; 37.17° N, 3.61° W, 680 m a.s.l.). The year 2014 has
119 been used to develop the proposed model in this work, and 2015 to validate it along with
120 other stations mentioned later. From this data base, hourly values have been generated
121 covering the two-year period, guaranteeing the inclusion of a wide range of seasonal
122 conditions and solar zenith angles. The photosynthetic photon flux density (Q_p) was measured
123 using LICOR model 190 SA quantum sensors (Lincoln, Nebraska, USA). Global irradiance,

124 R_s , was measured using a Kipp and Zonen model CM-11 radiometer (Delft, Netherlands). The
125 quantum sensor has a relative error less than 5% estimated by the manufacturer and
126 measurements of global irradiance have an estimated experimental error of about 2-3%. The
127 calibration constants of the instruments were checked periodically. To avoid problems
128 associated to the instrument deviations from the ideal cosine law, we limited our study to
129 solar zenith angles less than 85° (Alados et al., 1996, Alados-Arboledas et al., 2000).

130 Granada is a non-industrialized medium-sized located in south-eastern Spain and
131 situated in a natural basin surrounded by mountains with elevations between 1000 and 3500 m
132 a.s.l.. Near continental conditions prevailing at this site are responsible for large seasonal
133 temperature differences, providing cool winters and hot summers. Most rainfall occurs during
134 winter and spring.

135 In order to evaluate the proposed model we have used data from different stations not
136 used in the model development, two in Spain (Granada and Jaén), two in Japan (Sapporo and
137 Fujiyoshida) and another one in Argentina (Luján). The quantum sensors used in Jaén (Spain)
138 and Japan are photodiodes (LI-190; Li-Cor, Lincoln, NE, USA) and that in Argentina was an
139 instrument designed by Argentina researchers with similar characteristics to photodiode LI-
140 190. Detailed information about instrumentation and sites characteristics can be found in
141 Chamizo et al. (2016) for Jaén (Spain), in Mizoguchi et al. (2014) for Japan and in Denegri
142 (2014) for Argentina. Earlier studies with Q_p measurements were conducted in Europe and
143 the Americas, where key climate conditions such as precipitation are different from those in
144 Asia. The stations selected here furthermore cover the northern and southern hemispheres.
145 Thus, model evaluation can be carried out at sites with very different climatic characteristics
146 and altitudes. Table 1 presents some climatic data for stations used in this study. Especially
147 relevant are the differences in yearly precipitation.

148

149 **3. RESULTS AND DISCUSSION.**

150

151 **3.1. ANALYSIS OF RATIO Q_p/R_s**

152

153 Before doing an in-depth, direct study of the variable of interest (Q_p), we have
154 analysed the ratio between Q_p and R_s (Q_p/R_s), as it is a common practice to express Q_p as a
155 fraction of R_s , in order to explore the systematic relationships between both variables. The
156 most important factor influencing the levels of solar radiation reaching the Earth's surface is
157 solar position. Some authors showed that Q_p/R_s increases with solar zenith angle (θ) (e.g.
158 Meek et al., 1984; Udo and Aro 1999) whereas others indicated the opposite (e.g., Ge et al.,
159 2011; González and Calbó, 2002). On the other hand, many authors reported dependences on
160 site or season (Alados et al., 1996; Aguiar et al., 2012; Li et al., 2010, Jacovides et al., 2003),
161 with higher values in summer and lower values during winter. In any case, the reasons for any
162 dependence of this ratio on solar position remain unclear. Akitsu et al. (2015) found a slight
163 negative correlation ($R = -0.323$), but considered it a false correlation caused by the lower
164 solar elevation in winter. Thus, previous reported dependencies on θ may be attributed to this
165 false correlation or to artifacts (such as cosine and spectral errors) in the quantum sensor and
166 the pyranometers (Akitsu et al., 2015).

167 Figure 1 shows the dependence of Q_p/R_s on θ through the cosine function. We can
168 observe that this ratio varies between 1.52 and 2.39 $\mu\text{mol J}^{-1}$ with a mean value of 1.95 ± 0.12
169 ($\mu\text{mol J}^{-1}$), and with slightly more scatter for longer path lengths. Thus, we find no clear
170 dependence of the ratio Q_p/R_s on solar position, even though the mean values tend to higher
171 values for solar position close to the zenith. Alados et al. (1996) also found no dependence on
172 θ although with more dispersion around the mean value for longer path lengths.

173 If we assume this ratio being constant it implies the possibility of a linear model to
174 estimate Q_p from global irradiance measurements, whose intercept would be zero. This is the
175 simplest model to estimate Q_p from R_s with acceptable estimation errors, but the relationship
176 between Q_p and R_s changes with cloud condition and season (Mizoguchi et al., 2014). Our
177 mean value is similar to that proposed by Aguiar et al. (2012) in their model ($1.94 \mu\text{mol J}^{-1}$)
178 although with a non-zero intercept ($2.20 \mu\text{mol J}^{-1}$). Bai (2012) found similar mean values for
179 this ratio in two stations in North China: ($1.95 \pm 0.12 (\mu\text{mol J}^{-1})$) at Yucheng and 1.94 ± 0.14
180 ($\mu\text{mol J}^{-1}$) at Luancheng). Hu et al. (2007) found values between 1.75 ± 0.12 to 2.30 ± 0.15
181 ($\mu\text{mol J}^{-1}$) in many places of China.

182 The ratio Q_p/R_s shows scatter values that increase with decreasing R_s (Figure 2). Thus,
183 for high values of R_s measurements the ratio tends to a constant value close to the mean value
184 of the ratio ($1.95 \pm 0.12 (\mu\text{mol J}^{-1})$), whereas for a given low value of R_s , the ratio can reach
185 values between 1.52 and $2.39 \mu\text{mol J}^{-1}$. The ratio Q_p/R_s shows no dependence on clearness
186 index (Figure 3). This parameter (k_t) is defined as the ratio of the global irradiance (R_s) to the
187 extraterrestrial global irradiance (R_{so}), both on a horizontal surface:

$$188 \quad k_t = \frac{R_s}{R_{so}} \quad (1)$$

189 The solar constant value used to calculate R_{so} is 1367 Wm^{-2} (Iqbal, 1983). The
190 parameter k_t characterizes the sky condition including the attenuation effects of the most
191 dominant factors controlling solar radiation, such as clouds and aerosols. Previous studies
192 have used similar parameters such as the sky clearness and the sky brightness (Pérez et al.,
193 1990) to characterize the sky condition (Alados et al., 1996). Despite the lack of dependence,
194 higher values are detected for lower values of k_t similarly to Figure 2, but with the difference

195 of the same high dispersion for the whole range of k_t values. For $k_t < 0.2$, this ratio can attain
196 values close to $2.5 \mu\text{mol J}^{-1}$. In this sense, Tsubo and Walker (2005) found a simple function
197 between both variables in which the ratio increases with decreasing k_t . Also a slight negative
198 dependence ($R = -0.380$) was found by Akitsu et al. (2015), a result that can be attributed to
199 the presence of clouds in skies, with high solar absorption in the infrared region. Clouds
200 absorb NIR more strongly than Q_p , hence the transmittance of Q_p through clouds is larger than
201 NIR. In fact, clouds attenuate more markedly total solar spectral range than shorter
202 wavelengths such as ultraviolet range (Foyo-Moreno et al., 2001, 2003).

203 Now, to explicitly account for absorption by water vapour on the solar spectrum, we
204 have studied the dependence on meteorological parameters such as the dewpoint temperature
205 (T_d) – relevant due to its correlation with the amount of precipitable water (Reitan, 1963) –
206 and the partial vapour pressure (e), shown in Figure 4. Our results shows no evidence about
207 the dependence with these parameters, although the mean values tend to higher values with
208 increasing T_d and e . Akitsu et al. (2015) found positive correlation with e ($R=0.847$), with the
209 ratio Q_p/R_s increasing with e regardless of sky condition, from 1.9 to $2.2 \mu\text{mol J}^{-1}$ as e
210 increased from 2 to 30 hPa. This increase can be attributed to the absorption of near-infrared
211 radiation by water vapour. The dependence of Q_p/R_s on e has been roughly described in
212 seasonal variations such as higher values in summer and lower in winter (Rao, 1984;
213 Papaioannou et al., 1996). Alados et al. (1996) found a positive correlation with T_d , with the
214 ratio Q_p/R_s increasing as T_d increases, because of the enhancement of the extinction process in
215 the infrared region of the solar spectrum. There is a seasonal dependence with higher values
216 in summer and lower in winter, which can be explained by the seasonal pattern of the water
217 vapour content, with higher values in summer and lower values in winter.

218

3.2. ANALYSIS OF Q_p .

219

220

221

222

223

224

225

226

227

228

229

230

231

232

233

234

235

236

In this section, we directly study the dependence of Q_p on the same parameters considered in the above section. First, we show that the solar zenith angle alone does not allow for a correct determination of Q_p (Figure 5). While Q_p increases with decreasing θ , due to high dispersion it is possible to parameterise only the envelope of the data through a simple linear dependence on the cosine of θ . The envelope of the data corresponds to the maximum values that can be assumed to values corresponding to clear skies. This relationship is similar to other spectral solar ranges, such as UV or erythemal ultraviolet irradiance (Foyo-Moreno et al., 1998; 2007). So we can propose a linear function to estimate the maximum values of Q_p using $\cos \theta$, with the intercept selected as zero. Once the maximum values are determined for a given solar position, the real value will be modulated or attenuated depending on sky conditions (i.e., k_t). Consequently, Q_p measurements can be estimated from a linear expression using the product of two parameters, k_t and $\cos \theta$, as input data following the procedure used by Foyo-Moreno et al. (1998, 2007) with the proposal of a new empirical model to estimate UV and later adaptation to estimate erythemal ultraviolet irradiance extending to other spectral wavelengths ranges. In order to establish a simple model to estimate Q_p , in this work we present a version slightly different.

237

238

239

240

241

Following the same scheme until now, we show the high positive correlation between Q_p and R_s (Figure 6; $R^2=0.998$) as it was anticipated, with a slope of $1.989 \pm 0.001 \mu\text{mol J}^{-1}$, close to the mean value found for the ratio Q_p / R_s ($1.95 \pm 0.12 (\mu\text{mol J}^{-1})$). Similar values for the slope have been found for three sites in Europe by means of data from the EUROFLUX database (Ceulemans et al., 2003).

242 We found no dependence of Q_p on e nor T_d (not shown here). Regarding dependence
243 on k_t , we were able to parameterize two envelopes for the maximum and minimum possible
244 values of Q_p at a given value of k_t (Figure 7). Wang et al. (2013, 2014, 2015) carried out a
245 similar analysis and proposed a model including two variables: k_t and θ . The equation was:

$$246 \quad Q_p = (a + b k_t + c k_t^2 + d k_t^3) \cos \theta^e \quad (2)$$

247 where a , b , c , d and e are the fitting coefficients.

248 The above analyses demonstrate that it is more appropriate to model Q_p rather than
249 Q_p/R_s , and that the determinants parameters are θ , R_s and k_t . So we have eliminated the
250 meteorological parameters e and T_d from our analysis, although some models do consider
251 them (e.g. Aguiar et al., 2012).

252 Prior to proposing a simple model, we have used Artificial Neural Network (ANN)
253 including the combination of various input data with these parameters to finally select one or
254 more parameters. The combination of input variables taken into consideration in this work
255 were the product k_t and $\cos \theta$, since Figure 5 showed a linear dependence of Q_p on $\cos \theta$ for
256 the maximum values of Q_p , and for a given solar position the real value is modulated by k_t
257 representing broadband transmittance depending on both, cloudiness and aerosol load. This
258 can be viewed as a multivariable interpolation problem requiring estimation of the function
259 relating the input to the output (Alados et al., 2004, 2007). For six model versions, ANN used
260 as input variables: (model 1) R_s , (model 2) $k_t \cdot \cos \theta$; (model 3) R_s and $\cos \theta$, (model 4) R_s and
261 k_t , (model 5) k_t and $\cos \theta$; and (model 6) R_s , k_t and $\cos \theta$. The ANN used in our study is a
262 multi-layer perceptron (MLP) with three layers (input layer, hidden layer and out layer). The
263 optimal number of neurons in the hidden layer was selected following an empirical procedure.
264 The output layer has one neuron, the estimated Q_p . The design of an ANN requires the use of
265 training and testing data sets. Data from different sites are used in training and selecting the

266 best ANN. The validation data set was carried out with data from different stations shown in
267 Table 1. The performance of the models was evaluated using the RMSE (Root Mean Square
268 Error) and the MBE (Mean Bias Error), calculated as percentage of the mean experimental
269 values, given by the following expressions:

$$270 \quad MBE = \frac{100}{M_{ave}} \frac{1}{N} \sum_{i=1}^N (E_i - M_i) \quad (3)$$

$$271 \quad RMSE = \frac{100}{M_{ave}} \left(\frac{1}{N} \sum_{i=1}^N (E_i - M_i)^2 \right)^{0.5} \quad (4)$$

272 where E_i is the estimated value (i th number), M_i is the measured value, M_{ave} is the average of
273 the measured values and N is the number of data. These statistics allow detection of both the
274 differences between model estimates and experimental data, and any systematic data over- or
275 underestimation tendencies. In general, all models present good results with R^2 above 0.99
276 with MBE lower than 5% and RMSE lower than 10% (Table 2), but at all sites model 2
277 presents lower MBE and model 5 lower RMSE; both of these superior models used θ and k_t
278 as input variables. Thus, the most adequate variables to model Q_p are θ and k_t . Zempila et al.
279 (2016) developed and assessed the performance of linear regression, multiple linear
280 regression and nonlinear neural networks to calculate Q_p from R_s measurements using also
281 information about θ , the columnar perceptible water vapour and the aerosol optical depth.
282 Jacovides et al. (2015) also used Artificial Neural Network models for estimating daily solar
283 global UV, Q_p and broadband radiant fluxes in an eastern Mediterranean site.

284

285 **3.3. PERFORMANCE OF MODEL.**

286

287 Taking into account the already shown results we propose a new simple empirical
288 model to estimate Q_p through the expression:

289
$$Q_p = a k_t \cos \theta \quad (5)$$

290 with $a = 2681 \pm 2 \mu\text{mol m}^{-2}\text{s}^{-1}$ and a determination coefficient (R^2) of 0.999. The model has
291 been evaluated with experimental data from different stations which were not used in its
292 development, two in Spain (Granada, 2015; and Jaén), two in Japan, and another in Argentina
293 (Table 1). These stations can be considered to represent a wide range of global climatic
294 conditions. Other empirical models include these parameters and others but involving a more
295 complicated formulation. Moreover, small improvements of nonlinear models over linear
296 models have been detected (Zempila et al., 2016).

297 Table 3 shows the results obtained for each station including the coefficient of
298 determination R^2 - the fraction of experimental data variance explained by the model -, the
299 slope b , and the intercept a of the linear regression of estimated Q_p versus measured values.
300 Table 3 also shows the Mean Bias Error (MBE) and Root Mean Square Error (RMSE), both
301 as percentage of the mean experimental values.

302 Figure 8 shows Q_p estimated via eq. (5) at all localities analysed in this work. The model
303 evaluation is highly satisfactory for all radiometric stations, the variance explained for the
304 models is better than 99%, and the values for MBE and RMSE are low. In fact, for three
305 localities the MBE values are less than 1%, with larger values at Sapporo (-3.1%) and Conde
306 (-1.3%). The RMSE values do not exceed 6.1%. The model underestimates in all locations. In
307 any case, both the slope and correlation coefficient of the linear regression between measured
308 and estimated values reveals the goodness of the model estimations. Wang et al. (2016)
309 developed and evaluated Q_p estimating models at different types of ecosystems in China, their
310 results showed large differences in model accuracy for each model at each ecosystem. RMSE
311 ranges between 6.45 and 13.08 % and MBE ranges between 4.02 and 8.89 % for a semi-
312 empirical all-sky model using as input variables k_t and $\cos \theta$. Alados et al. (2000), found MBE

313 values about 4% and RMSE lower than 16% at two stations located at Granada and Almería
314 using a cloudless parametric model combined with the cloud transmittance, but their scheme
315 requires information about cloud cover.

316

317 **3. CONCLUSIONS**

318

319 In order to formulate a simple empirical model to estimate the photosynthetic photon
320 flux density (Q_p) under all sky conditions, with measurements widely available at most
321 radiometric stations such as global irradiance (R_s), this work presents the results showing the
322 dependences of the ratio Q_p/R_s and Q_p with different parameters using the data acquired in
323 Granada, an urban site in Southeast Spain, during two recent years (2014-2015).

324 The solar position and cloudiness and aerosols conditions are found to be the most
325 important factors to be considered for all sky conditions in order to estimate R_s , and
326 consequently to estimate Q_p . Our results show no dependence of the ratio Q_p/R_s on the
327 variables selected, but a certain dependence of Q_p on them. Based on an analysis using
328 Artificial Neural Network (ANN), including various input data with combinations of these
329 parameters, we propose a simple model using one only input variable, defined as product of
330 the cosine of solar zenith angle ($\cos \theta$) and the clearness index (k_t), which depends both on
331 cloud cover and aerosol load. This model has the advantage of requiring only measurements
332 of R_s , available at most radiometric stations, and Q_p and R_s that can be obtained with the same
333 time interval. The model has been evaluated at five stations in Spain, Japan and Argentina,
334 with different climatic characteristics located in the northern and southern hemispheres. The
335 model provides satisfactory results, with low mean bias error for all stations; at three stations
336 Mean Bias Error (MBE) is less than 1% and the Root Mean Square Error (RMSE) below

337 6.1% for all stations. These results suggest that this model can be widely applied with better
338 accuracy than other empirical models.

339

340 **ACKNOWLEDGEMENTS:** This work was supported by the Andalusia Regional
341 Government through projects P11-RNM-7186 and P12-RNM-2409, by the Spanish Ministry
342 of Economy and Competitiveness through projects CGL2013-45410-R and CGL2014-52838-
343 C2-1-R and by the European Union's Horizon 2020 research and innovation programme
344 through project ACTRIS-2 (grant agreement No 654109). The authors would like to thank the
345 Forestry and Forest Products Research Institute FluxNet Database, FFPRI
346 <http://www2.ffpri.affrc.go.jp/labs/flux/>), also to Dr. E. Sánchez Cañete by the data from
347 Conde and Dra. M. J. Denegri for the data from Luján.

348

349 **REFERENCES**

350

- 351 Aguiar LJG, Fischer GR, Ladle RJ, Malhado ACM, Justino FB, Aguiar RG, Costa JMN.
352 2012. Modeling the photosynthetically active radiation in South West Amazonia under all
353 sky conditions. *Theor. Appl. Climatol.* 108, 631–640.
- 354 Augustine JA, DeLuisi JJ, Long CN. 2000. SURFRAD-A national surface radiation budget
355 network for atmospheric reserach, *Bull. Am. Meteorol. Soc.*, 81, 2341-2358.
- 356 Alados I, Foyo-Moreno I, Alados-Arboledas L. 1996. Photosynthetically active radiation:
357 measurements and modeling. *Agric. For. Meteorol.* 78, 121-131.
- 358 Alados-Arboledas L, Olmo FJ, Alados I, Pérez M. 2000. Parametric models to estimate
359 photosynthetically active radiation in Spain. *Agric. For. Meteorol.* 101, 187-201.

360 Alados I, Olmo FJ, Foyo-Moreno I, Alados-Arboledas L. 2000. Estimation of
361 photosynthetically active radiation under cloudy conditions. *Agric. For. Meteorol.* 102, 39-
362 50.

363 Alados I, Mellado JA, Ramos F, Alados-Arboledas L. 2004. Estimating UV erythemat
364 irradiance by means of neural networks. *J. Photochem. Photobiol.*, 80(2), 351–358.

365 Alados I, Mellado JA, Foyo-Moreno I, Alados-Arboledas L. 2007. Neural network for
366 estimation of UV erythemat irradiance using solar broadband irradiance. *Int. J. Climatol.*,
367 27, 1791–1799.

368 Akitsu T, Kume A, Hirose Y, Ijima O, Nasahara KN. 2015. On the stability of radiometric
369 ratios of photosynthetically active radiation to global solar radiation in Tsukuba, Japan.
370 *Agric. For. Meteorol.* 290-2010, 59-68.

371 Baldocchi D, Falge E, Gu L, Olson R, Hollinger D, Running S, Anthoni P, Bernhofer Ch,
372 Davis K, Evans R, Fuentes J, Goldstein A, Katul G, Law B, Lee X, Malhi Y, Meyers T,
373 Munger W, Oechel W, Paw KT, Pilegaard K, Schmid HP, Valentini R, Verma S, Vesala
374 T, Wilson K, Wofsy S.. 2001, FLUXNET : A new tool to study the temporal and spatial
375 variability of ecosystem-scale carbon dioxide, water vapor, and energy flux densities, *Bull.*
376 *Am. Meteorol. Soc.*, 82, 2415-2434.

377 Bai J. 2012. Observations and estimations of PAR and solar visible radiation in North China.
378 *J. Atmos. Chem.*, 69, 231-252.

379 Calbó J, Pagès D, González J. 2005. Empirical studies of cloud effects on UV radiation: A
380 review, *Rev. Geophys.*, 43, RG2002, doi:10.1029/2004RG000155.

381 Cao MK, Prince SD, Tao B, Li KR. 2005. Regional pattern and interannual variations in
382 global terrestrial carbon uptake in response to changes in climate and atmospheric CO₂.
383 *Tellus B*, 57, 210-217.

384 Ceulemans R, Kowalski AS, Berbigier P, Dolman AJ, Grelle A, Janssens IA, Lindroth A,
385 Moors E, Rannik Ü, Vesala T. 2003. Coniferous Forests (Scots and Maritime Pine):
386 Carbon and water fluxes, balances, ecological and ecophysiological determinants,
387 Ecological Studies, 163, 71 – 97.

388 Chamizo S, Serrano-Ortiz S, López-Ballesteros A, Sánchez-Cañete EP, Vicente-Vicente JL,
389 Kowalski AS. 2016. Net ecosystem CO₂ exchange in an irrigated olive orchard of SE
390 Spain: influence of weed cover. Agriculture, Ecosystems and Environment (submitted).

391 Denegri MJ. 2014. Medición de la Radiación Fotosintéticamente Activa (PAR) en planos con
392 diferentes orientaciones, mediante el desarrollo y ensayo de un dispositivo ad-hoc, y
393 verificación de modelos. Doctoral Thesis, pp. 316.

394 Foyo-Moreno I, Vida J, Alados-Arboledas L. 1998. A simple all weather model to estimate
395 ultraviolet solar radiation (290-385nm). J. Apl. Metereol., 38, 1020–1026.

396 Foyo-Moreno I, Alados I, Olmo FJ, Vida J, Alados-Arboledas L. 2001. On the use of a cloud
397 modification factor for solar UV (290-385 nm) spectral range. Theor. Appl. Climatol., 68,
398 41-50.

399 Foyo-Moreno I, Alados I, Olmo FJ, Alados-Arboledas L. 2003. The influence of cloudiness
400 on UV global irradiance (295-385 nm). Agric. For. Meteorol. 120, 101-111.

401 Foyo-Moreno I, Alados I, Alados-Arboledas L. 2007. Adaptation of an empirical model for
402 erythemal ultraviolet irradiance. Ann. Geophys., 25, 1–10.

403 Ge S, Smith RG, Jacovides CP, Kramer MG, Carruthers RI. 2011. Dynamics of
404 photosynthetic photon flux density (PPFD) and estimates in coastal northern California.
405 Theor. Appl. Climatol. 105, 107-118.

406 González J, Calbó J. 2002. Modelled and measured ratio of PAR to global radiation under
407 cloudless skies. Agric. For. Meteorol. 110, 319–325.

408 Hu B, Wang Y, Liu G. 2007. Spatiotemporal characteristics of photosynthetically active
409 radiation in China. *J. Geophys. Res.*, 112, doi: 10.1029/2006JD007965.

410 Iqbal M. 1983. *An Introduction to Solar Radiation*. Academic Press: London.

411 Jacovides CP, Tymvios FS, Asimakopoulos DN, Theofilou KM, Pashiardes S. 2003. Global
412 photosynthetically active radiation and its relationship with global solar radiation in the
413 Eastern Mediterranean basin. *Theor. Appl. Climatol.* 74, 227–233.

414 Jacovides CP, Tymvios FS, Papaioannou G, Asimakopoulos DN, Theofilou KM. 2004. Ratio
415 of PAR to broadband solar radiation measured in Cyprus, *Agric. For. Meteor.*, 121(3-4),
416 135-140.

417 Jacovides CP, Tymvios FS, Boland J, Tsitouri M. 2015. Artificial Neural Network models for
418 estimating daily solar global UV, PAR and broadband radiant fluxes in an eastern
419 Mediterranean site, *Atmos. Res.*, 152, 138-145.

420 Li R, Zhao L, Ding Y, Wang S, Ji G, Xiao Y, Liu G, Sun L. 2010. Monthly ratios of PAR to
421 global solar radiation measured at northern Tibetan Plateau, China. *Sol. Energy* 84, 964–
422 973.

423 Meek DW, Hatfield JL, Howell TA, Idso SB, Reginato RJ. 1984. A generalized relationship
424 between photosynthetically active radiation and solar radiation. *Agron. J.* 76, 939–945.

425 Mizoguchi Y, Yasuda Y, Ohtani Y, Watanabe T, Kominami Y, Yamanoi K. 2014. A practical
426 model to estimate photosynthetically active radiation using meteorological elements in a
427 temperate humid area and comparison among models. *Theor. Appl. Climatol.* 115, 583-
428 589.

429 Ohmura A, Dutton E, Forgan B, Fröhlich C, Gilgen H, Hegner H, Heimo A, König-Langlo G,
430 Mcarthur B, Müller G, Philipona R, Pinker R, Whitlock CH, Wild M. 1998. Baseline

431 Surface Radiation Network (BSRN/WCRP): New precision radiometry for climate change
432 research, *Bull. Am. Meteorol. Soc.*, 79, 2115-2136.

433 Papaioannou G, Nikolidakis G, Asimakopoulus DN, Redalis D. 1996. Photosynthetically
434 active radiation in Athens. *Agric. For. Meteorol.* 81, 287–298.

435 Perez R, Ineichen P, Seals R, Michalsky JJ, Stewart R. 1990. Modelling daylight availability
436 and irradiance components from direct and global irradiance. *Solar Energy* 44, 271-289.

437 Rao CR. 1984. Photosynthetically active components of global solar radiation: measurements
438 and model computations. *Arch. Met. Geophys. Bioclim. Ser. B* 34, 353–364.

439 Reitan CH. 1963. Surface dewpoint and water vapour aloft. *J. Appl. Meteorol.*, 2: 776-779.

440 Tang W, Qin J, Yang K, Niu X, Zhang X. 2013. Reconstruction of daily photosynthetically
441 active radiation and its trend over China. *J. Geophys. Res. Atm.*, 118, 13292-13302.

442 Tsubo M, Walker S. 2005. Relationships between photosynthetically active radiation and
443 clearness index at Bloemfontein, South Africa. *Theor. Appl. Climatol.* 80:17–25.

444 Udo S, Aro T. 1999. Global PAR related to global solar radiation for central Nigeria.
445 *Agric. For. Meteorol.* 97, 21–31.

446 Wang L, Gong W, Lin A, Hu B, Ma Y. 2013. Measurement and estimation of
447 photosynthetically active radiation from 1961 to 2011 in Central China. *Appl. Energy* 111,
448 1010-1017.

449 Wang L, Gong W, Hu B, Zhu Z. 2014. Analysis of photosynthetically active radiation in
450 Northwest China from observations and estimation. *Int. J. Biometeorol.*, doi:
451 10.1007/s00484-014-0835-3.

452 Wang L, Gong W, Feng L, Lin A, Hu B, Zhou M. 2015. Estimation of hourly and daily
453 photosynthetically active radiation in Inner Mongolia, China, from 1990 to 2012. *Int. J.*
454 *Climatol.*, 35, 3120-3131.

455 Wang L, Kisi O, Zounemat-Kermani M, Hu B, Gong W. 2016. Modeling and comparison of
456 hourly photosynthetically active radiation in different ecosystems. *Renew. & Sustain.*
457 *Energy Rev.*, 56: 436-453.

458 Yu G, Wen X, Sun X, Tanner B, Lee X, Chen J. 2006. Overview of China FLUX and
459 evaluation of its eddy covariance measurements, *Agric. For. Meteor.*, 137, 125-137.

460 Zempila MM, Taylor M, Bais A, Kazadzis S. 2016. Modeling the relationship between
461 photosynthetically active radiation and global horizontal irradiance using singular
462 spectrum analysis. *J. of Q. Spec. & Rad. Transfer*, 182, 240-263.

463