



Synergistic effects of UVR and simulated stratification on commensalistic phytoplankton–bacteria relationship in two optically contrasting oligotrophic Mediterranean lakes

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Abstract. An indirect effect of global warming is a reduction in the depth of the upper mixed layer (UML) causing organisms to be exposed to higher levels of ultraviolet (UVR, 280–400 nm) and photosynthetically active radiation (PAR, 400–700 nm). This can affect primary and bacterial production as well as the commensalistic phytoplankton–bacteria relationship. The combined effects of UVR and reduction in the depth of the UML were assessed on variables related to the metabolism of phytoplankton and bacteria, during in situ experiments performed with natural pico- and nanoplankton communities from two oligotrophic lakes with contrasting UVR transparency (high-UVR versus low-UVR waters) of southern Spain. The negative UVR effects on epilimnetic primary production (PP) and on heterotrophic bacterial production (HBP), intensified under increased stratification, were higher in the low-UVR than in the high-UVR lake, and stronger on the phytoplanktonic than on the heterotrophic bacterial communities. Under UVR and increased stratification, the commensalistic phytoplankton–bacteria relationship was strengthened in the high-UVR lake where excretion of organic carbon (EOC) rates exceeded the bacterial carbon demand (BCD; i.e., BCD : EOC(%) ratio < 100). This did not occur in the low-UVR lake (i.e., BCD : EOC(%) ratio > 100). The greater UVR damage to phytoplankton and bacteria and the weakening of their commensalistic interaction found in the low-UVR lake indicates that these ecosystems would be especially vulnerable to UVR and increased stratification as stressors related to global climate change. Thus, our findings

may have important implications for the carbon cycle in oligotrophic lakes of the Mediterranean region.

1 Introduction

Rising levels of greenhouse gases (mainly CO₂), attributed to human activities, have led to an increase of 0.56 °C in the Earth's surface temperature over the past 150 years (IPCC, 2013). Model predictions indicate greater temperature increases, ranging from 1.5 to 6.4 °C by the end of the century. Major changes in precipitation have accompanied these temperature variations and are expected to become more pronounced (IPCC, 2013). These climate changes affect aquatic ecosystems by increasing water temperature, altering mixing regimes, shortening the thaw time and the duration of ice cover, and/or strengthening water-column stratification (de Senerpont Domis et al., 2013). These alterations in physical conditions have different effects on primary and bacterial production, plankton growth, nutrient supply, and trophic interactions, among other ecological processes (de Senerpont Domis et al., 2013). In addition, variations in stratification patterns are known to strongly affect biogeochemical cycles (van de Waal et al., 2009).

Higher temperatures in the upper layers of freshwater bodies increase density differences between the upper mixed layer (UML) or epilimnion and deeper waters, augmenting the vertical temperature gradient, and thus the stratification.

This process has contrasting effects on nutrient and light availability for organisms' growth. On the one hand, stratification reduces the flow of nutrients from deep and nutrient-rich areas into the UML, limiting nutrient availability for growth (Huisman et al., 2006). On the other hand, stratification traps phytoplankton populations in surface layers, increasing the light available for growth, but also exposing them to higher levels of ultraviolet radiation (UVR, 280–400 nm). In this regard, it has been widely reported that greater exposure to UVR exerts an inhibitory effect on autotrophic and heterotrophic organisms (Häder et al., 2011), and that UV-B (280–315 nm), in particular, harms primary and bacterial production (Carrillo et al., 2002), enzymatic activity (Korbee et al., 2012), and cell viability (Helbling et al., 1995), among other effects. However, it has been also reported (Aas et al., 1996; Medina-Sánchez et al., 2002; Gao et al., 2007) that UVR does not produce negative effects and it can even stimulate bacterial production and photosynthetic activity. These opposite effects may be attributable to the high acclimation capacity of organisms in high-UVR ecosystems (Medina-Sánchez et al., 2002; Ruiz-González et al., 2013) or to differences in physical–chemical factors (e.g., temperature or nutrient content) among ecosystems (Harrison and Smith, 2009).

With respect to physical factors, it has been experimentally demonstrated (Helbling et al., 1994) that vertical mixing can alter UVR-induced effects on planktonic organisms by generating a regime of fluctuating irradiance, with high values near the surface and low values at the bottom of the UML. The depth of the UML also influences the mean UVR and PAR irradiance received by organisms and the duration of their residence in the photoactive zone (Neale et al., 2003). Studies on the interactive effects of UVR and vertical mixing on phytoplankton (Helbling et al., 1995; Neale et al., 2003) and bacteria (Bertoni et al., 2011) have shown that these organisms can recover from UVR-induced damage when UVR exposure is subsequently reduced or avoided. The outcome of damage vs. repair depends not only on the amount of damaging UVR received but also on photorepair wavelengths (UV-A, PAR) to which organisms are subsequently exposed during the fluctuating radiation regime. Moreover, the effects of different mixing depths, and thus of different mean irradiances, can act synergistically or antagonistically with UVR, depending on the composition, structure, and size of the species, as well as on the environmental conditions (Villafañe et al., 2007). For instance, Barbieri et al. (2002) found that the impact of UVR in Patagonian coastal waters was negative or positive depending on the fraction of the euphotic zone (Z_{eu}) that was mixed; thus, UVR was used for photosynthesis when vertical mixing reached $\sim 90\%$ of the Z_{eu} , but carbon fixation was reduced by UVR when the UML was shallow ($\sim 60\%$ of the Z_{eu}).

Besides increased stratification of the water column, more extreme rainfall events and storms are predicted in many parts of the Earth in the global-change scenario (IPCC,

2013). This would increase the amount of allochthonous dissolved organic matter (DOM) reaching inland and coastal aquatic ecosystems, reducing the penetration of incident UVR (Rose et al., 2009). The UVR filtering characteristics of colored DOM (CDOM) result in a more effective attenuation of shorter (UV-B) than longer (UV-A, 315–400 nm) wavelengths, as also observed for stratospheric ozone. Concomitantly, the photochemical reactions mediated by UVR lead to (i) the photodegradation of DOM, altering the composition and absorbance of CDOM, and (ii) the photooxidation of DOM, producing oxygen free-radicals (Kitidis et al., 2014). These changes would modulate the response of aquatic organisms to UVR (Williamson and Rose, 2010), making it more complex to predict the interactive effects of UVR and stratification on the planktonic community.

Recent experiments carried out by our group have demonstrated that fluctuating irradiance increases the harmful UVR effects on primary producers in oligotrophic mountain lakes with high DOM, whereas the opposite effects were detected in those with low DOM content (Helbling et al., 2013). Several authors have highlighted the importance of the quality of the radiation, which can interact with DOM and either increase or decrease the availability of organic carbon for bacteria (Pérez and Sommaruga, 2007). However, despite the key role of phytoplankton and heterotrophic bacteria production as a link between the microbial and grazing food webs, no studies on the interactive effects of radiation quality and increased stratification on the commensalistic phytoplankton–bacteria relationship have been done in ecosystems with high- and low-CDOM contents.

A growing body of literature supports the strong dependence of planktonic heterotrophic prokaryotes on organic matter released in situ by phytoplankton in the upper layers of aquatic ecosystems (Baines and Pace, 1991; Norrman et al., 1995; Morán and Alonso-Sáez, 2011). It has also been demonstrated that UVR exposure in the upper layers of the water column can increase the proportion of photosynthate released as exudates (Carrillo et al., 2008; Korbee et al., 2012), which would stimulate the growth of UVR-resistant bacteria (Xenopoulos and Schindler, 2003) and give rise to a coupled phytoplankton–bacteria relationship in clear oligotrophic lakes (Carrillo et al., 2002). Coupling between phytoplankton and bacterioplankton has been defined as the capacity of the carbon (C) released by phytoplankton to support the bacterial carbon requirement (Morán et al., 2002) and will therefore differ depending on (i) the availability of alternative (allochthonous or autochthonous) carbon sources (Gasol et al., 2009) and (ii) the supply of inorganic nutrients (Medina-Sánchez et al., 2010; 2013; López-Sandoval et al., 2011). Although the bacterial dependence on C released by phytoplankton is a well-established paradigm in aquatic microbiology (Cole et al., 1988), it is currently under renewed debate. Thus, Fouilland and Mostajir (2010, 2011) proposed that C dependency of bacteria on phytoplankton is uncertain because C sources other than those from algal origin

might support the bacterial growth more significantly. However, Morán and Alonso-Sáez (2011) rebutted this idea due to uncertainty found in the application of different conversion factors to raw data and modeled rates in the Fouilland and Mostajir's calculations.

With this background, the aim of the present study was to improve our understanding about the combined effects of UVR exposure and increased stratification on (i) phytoplanktonic and heterotrophic bacterial production and (ii) the commensalistic relationship between them in lakes with different transparency to UVR. We hypothesized that the interactive effects of UVR and increased stratification will accentuate the harmful UVR effects on primary production (PP) and heterotrophic bacterial production (HBP), thus resulting in a greater C release by phytoplankton, which will strengthen the commensalistic phytoplankton–bacteria relationship. These effects will be more acute in low-UVR than in high-UVR lakes, because UVR-resistant populations are likely not selected in low-UVR lakes.

To test our hypothesis, we carried out in situ experiments to assess the combined impact of solar radiation (i.e., quality) and simulated stratification on metabolism of phytoplankton and bacteria, and their commensalistic relationship, in two oligotrophic lakes with contrasting transparency to UVR in the Mediterranean region.

2 Methods

2.1 Model ecosystems

The study was performed during September 2011 in two Spanish oligotrophic lakes: La Caldera in Sierra Nevada National Park (37°03' N; 3°19' W; 3050 m a.s.l.) (Granada) and La Conceja in Ruidera Natural Park (38°55' N; 2°47' W; 850 m a.s.l.) (Ciudad Real). La Caldera is a mixed oligotrophic (total phosphorus [TP] < 0.3 µM and chlorophyll *a* [Chl *a*] < 5 µg L⁻¹) high-mountain lake above the treeline on a siliceous bedrock in a glacial cirque (Carrillo et al., 2006). This lake has a surface area of 2 ha and a mean depth of 4.3 m, with a maximum depth interannually varying from 2 to 14 m. UVR of considerable intensity penetrates deeply in the lake due to the high transparency of the water and low values of dissolved organic carbon ([DOC] < 0.08 mM) as reported in Carrillo et al. (2008) and Helbling et al. (2013). Therefore, this lake is hereafter called the “high-UVR” lake. The pelagic community is relatively simple (Carrillo et al., 2006) and it is characterized by the scarcity of ciliates and absence of heterotrophic nanoflagellates and autotrophic picoplankton, and no size overlaps exist between phytoplankton and heterotrophic bacteria (Medina-Sánchez et al., 2002). La Conceja is a stratified oligotrophic lake (TP < 0.03 µM and Chl *a* < 5 µg L⁻¹), although it has an elevated nitrate concentration which can exceed 800 µM due to agricultural use of the land. This lake has a surface area of 29 ha and

maximum depth of 14 m. The DOC content ranges from 0.15 to 0.25 mM. Therefore, this lake is hereafter called the “low-UVR” lake. The autotrophic community is composed of pico- and nanoplankton (Rojo et al., 2012).

2.2 Experimental setup

To assess the interactive effects of solar radiation quality (“UVR” factor) and stratification conditions (“STRAT” factor) on PP, HBP, TPR (total planktonic respiration < 45 µm fraction), and BR (bacterial respiration < 1 µm fraction in the high-UVR lake alone), samples were collected from the surface (0–0.5 m) epilimnetic water. An acid-cleaned 6 L horizontal Van Dorn sampler was used to collect the water that was pre-screened through a 45 µm mesh to remove large zooplankton prior to the experiments. Samples for PP were placed in 50 mL quartz flasks, and those for HBP, TPR, and BR in 25 mL quartz flasks. In the low-UVR lake, samples for PP, HBP, and TPR analyses were also gathered from the hypolimnetic water below the thermocline at 6 m depth, where UV-B did not reach the cells. The idea behind sampling these two communities in the low-UVR lake was to compare the responses of phytoplankton and bacterial communities that had different light histories and degree of acclimation to solar radiation when exposed to similar light quality treatments and irradiance conditions. Since this sharp contrast did not occur in the high-UVR lake, only samples from the 0–0.5 m were used in these experiments.

The experimental design consisted of three (for TPR and BR), four (for PP, HBP), or two (for TPR in the low-UVR lake) “UVR” treatments combined with the two stratification conditions. (1) The UVR treatments (triplicates for each condition) were as follows: (i) PAB – full solar radiation, uncovered quartz flasks; (ii) PA – exclusion of UV-B (280–320 nm), flasks wrapped with Folex 320 film (Folex, Germany); (iii) P, control – exclusion of UVR (280–400 nm), flasks wrapped with Ultraphan UV Opak395 film (Digrefa, Germany); and (iv) dark – flasks wrapped with black tape. The optical properties of the filters used for the radiation treatments have been published elsewhere (Villafañe et al., 2003); the filters were replaced before each experiment and tested using a double-beam spectrophotometer (Perkin-Elmer Lambda 40). (2) The stratification treatments were the following: (i) subsurface, samples incubated at 0.5 m depth; (ii) mixed, samples subjected to vertical mixing from 0 to 5 m depth. To simulate these reductions in the depth of the UML (i.e., from 5 m to near the surface), two round trays containing the samples were exposed in situ to solar radiation. One tray was placed at 0.5 m depth (subsurface) subjected to irradiance oscillations associated with waves at the surface. This treatment represents the worst-case scenario in terms of solar radiation (i.e., high summer irradiance conditions), in combination with a sharp increase of thermal stratification (i.e., simulating the formation of near-surface thermoclines) during the usually warm Mediterranean summer. Transient

thermoclines trapping phytoplankton very close to the surface have previously been detected in aquatic environments (Neale et al., 2003). The second tray was vertically moved between the surface and 5 m depth to simulate the irradiance changes in the upper 5 m of the water column (mixed). The speed of movement was 1 m every 2 min, achieved by a custom-made mixing simulator, using a frequency-controlled DC motor (Maxon motor, Switzerland) to impose a linear transport rate on the vessels from the surface to the mixing depth and back. The tray was placed on a boat anchored in a deep area of each lake in such a manner as to avoid shadows or any type of interference from the shoreline or boat. All incubations lasted for 3.5 h centered on local noon, and a total of 10 cycles (from the surface to 5 m depth to the surface again) were completed for the mixed condition.

Unfortunately, space restrictions within the trays prevented the performance of all experimental treatments in the low-UVR lake for TPR, which was measured only in samples exposed to PAB and P in the subsurface and mixed treatments. The overlapping between autotrophic and heterotrophic picoplankton precluded the measurements of BR in the low-UVR lake.

2.3 Physical measurements

Incident solar radiation was continuously monitored by means of a BIC radiometer (deck unit, Biospherical Instruments Inc., CA, USA) that has three channels in the UVR region of the spectrum (305, 320, and 380 nm) and one broadband channel for PAR (400–700 nm). Vertical profiles of solar radiation in the water column were performed at noon using a BIC radiometer (underwater unit) with temperature and depth sensors, in addition to the aforementioned channels. Vertical profiles of temperature and pH in the water column were measured using a multiparameter probe (Turo Water Quality Analysis T-611 Sandy Bay, Tasmania, Australia). These profiles were done daily at noon, and the temperature data were used to estimate the strength and depth of the epilimnion.

2.4 Chemical analyses

Samples for chemical and biological variables were collected with a 6 L Van Dorn sampler at the deepest central station at four depths in the high-UVR lake (surface, 5, 8, and 10 m) and six in the low-UVR lake (surface, 2, 4, 6, 8, and 10 m). Water samples were taken to determine the bacterial abundance (BA, 20 mL), phytoplankton species composition and abundance (250 mL), and Chl *a* (1 L). Samples were also collected for the chemical determination of total nitrogen (TN), TP, total dissolved nitrogen (TDN), total dissolved phosphorus (TDP), nitrate (NO_3^-), and soluble reactive phosphorus (SRP). The samples for TDN, TDP, NO_3^- , and SRP analyses were filtered through GF/F Whatman filters (47 mm in diameter) before the determinations. Samples for TP and

TDP were persulfate-digested at 120 °C for 30 min and determined (as for SRP) using 10 cm quartz cuvettes (following the acid molybdate technique; APHA, 1992). TN and TDN samples were also persulfate-digested and measured as NO_3^- by means of the ultraviolet spectrophotometric screening method (APHA, 1992). Blanks and standards were run in all procedures. DOC values were determined by filtering the samples through pre-combusted (2 h at 500 °C) glass fiber filters (Whatman GF/F) and acidifying them with HCl. Samples were then measured in a total organic carbon analyzer (TOC-VCSH/CSN Shimadzu).

2.5 Analysis biological variables

Chl a concentration: For measurements of the Chl *a* concentration, water samples collected from different depths in the water column were filtered onto Whatman GF/F filters (25 mm in diameter), which were frozen at –20 °C until analyses. For Chl *a* analysis, samples were thawed and placed in centrifuge tubes (15 mL) with 5 mL of acetone (90 %) for 24 h in the dark at 4 °C. Next, the samples were centrifuged, and the fluorescence of the supernatant was measured with a fluorometer (LS 55 Perkin Elmer, USA) (APHA, 1992).

Identification and cell counting: Samples for the identification and counting of phytoplankton were placed in 250 mL brown glass bottles and fixed with Lugol's reagent (approx. 1 % *v/v*). Sub-samples (100 mL) were settled for 48 h in Utermöhl chambers (Hydro-Bios GmbH), and species were then identified and counted using an inverted microscope (Leitz Fluovert FS, Leica, Wetzlar, Germany). Phytoplankton biovolumes were estimated from measurements of 20–30 cells of each species using image analysis (inverted microscope Axio Observer A1, Zeiss; high-resolution microscopy camera AxioCam HRC, Zeiss). Cell volume was calculated according to Carrillo et al. (1995), and converted to phytoplankton carbon using the conversion factors reported by Rocha and Duncan (1985). Bacterial abundance (BA) was determined by the 4',6-diamidino-2-phenylindole (DAPI) direct-count method described by Porter and Feig (1980). Water samples were fixed with neutralized formaldehyde (2 %), stained with DAPI to a final concentration of 2.5 $\mu\text{g mL}^{-1}$, and then filtered onto 0.2 μm pore-size polycarbonate black filter. At least 400 cells per sample were counted by means of epifluorescence microscopy (Carl Zeiss AX10). Bacterial biomass (BB) was estimated from bacterial biovolume, measured from bacterial images obtained by transmission electron microscopy (TEM) as described by Medina-Sánchez et al. (1999).

2.6 Analysis of biotic functional variables

Primary production and excreted organic carbon: For PP measurements, samples of phytoplankton communities were placed in 50 mL round quartz flasks (three clear and one dark per radiation treatment), inoculated with 0.37 MBq of

$\text{NaH}^{14}\text{CO}_3$ (specific activity: $310.8 \text{ MBq mmol}^{-1}$; DHI Water and Environment, Germany), and exposed to solar radiation in situ, as described above. The total organic carbon (TOC) produced was measured on 4 mL aliquots before filtration. The samples for PP were filtered onto $0.2 \mu\text{m}$ filters (25 mm diameter, Nuclepore, Whatman) under low vacuum ($< 100 \text{ mm Hg}$) to minimize cell breakage. Excretion of organic carbon (EOC) was measured on 4 mL aliquots from the filtrates ($< 0.2 \mu\text{m}$). Both filters and filtrates were placed in 20 mL scintillation vials and acidified with $100 \mu\text{L}$ of 1 N HCl for 24 h (no bubbling) to remove inorganic ^{14}C before the addition of liquid scintillation cocktail (Ecoscint A) to the vials. The amount of organic carbon produced was obtained by counting disintegrations per minute (dpm), using an auto-calibrated scintillation counter (Beckman LS 6000 TA). The total CO_2 in the lake water was calculated from alkalinity and pH measurements (APHA, 1992). In all calculations, dark values were subtracted from the corresponding light values (more details in Carrillo et al., 2002). The %EOC was estimated as

$$\% \text{EOC} = 100 \times (\text{EOC}/\text{TOC}). \quad (1)$$

Heterotrophic bacterial production: Samples for HBP measurements were placed in 25 mL quartz flasks and exposed in situ for 3.5 h under the radiation and stratification conditions as described above. Then, the HBP was determined as incorporation of ^3H -thymidine (specific activity = 52 Ci mmol^{-1} , Amersham Pharmacia) into the bacterial DNA, in darkness. Briefly, ^3H -thymidine was added to independent sets of five (three replicates plus two blanks per treatment) sterile microcentrifuge tubes filled with 1.5 mL of the pre-exposed samples to a final (saturating) concentration of 15.2 nM. The vials were then incubated at in situ temperature in darkness for 1 h. After incubation, the incorporation of ^3H -thymidine was stopped by adding trichloroacetic acid (TCA, 6 % final concentration). Likewise, blanks were TCA-killed before the radiotracer was added. After the cold TCA extraction, the precipitate was collected by centrifugation at 14 000 rpm for 10 min. The conversion factor $1.5 \times 10^{18} \text{ cell mol}^{-1}$ was used to estimate the number of bacteria produced per mole of incorporated ^3H -thymidine (Bell, 1993). The factor $20 \text{ fg C cell}^{-1}$ was applied to convert bacterial production into C (Lee and Fuhrman, 1987).

Respiration rates: Samples for TPR ($< 45 \mu\text{m}$ fraction) and BR ($< 1 \mu\text{m}$ fraction) measurements were placed in 25 mL quartz flasks and exposed in situ for 3.5 h under the radiation and stratification conditions described above. TPR and BR rates were measured in darkness using optode sensor spots (SP-PSt3-NAU-D5-YOP; PreSens GmbH, Germany) and an optic-fiber oxygen transmitter (Fibox 3; PreSens GmbH, Germany) connected to a computer. Data were recorded using OxyView 3.51 software (PreSens GmbH). The system was calibrated by a two-point calibration, together with data of atmospheric pressure and temperature,

before each experiment, following the manufacturer's recommendations. Measurements were made at the initial time (t_0) and then every hour over 8 h. Every oxygen measurement was done over 30 s with a frequency of 1 datum s^{-1} ; only the last 10 data points of each measurement were used in our analysis to ensure the stability of the data. Oxygen data were then adjusted to a linear model via least-squares regression. Slope of the regressions provided the oxygen consumption rates ($\mu\text{M O}_2 \text{ h}^{-1}$) (Warkentin et al., 2007). Oxygen was converted into carbon units using a respiratory quotient of 1 (del Giorgio and Cole, 1998).

The bacterial carbon demand (BCD) is the HBP plus BR. The bacterial growth efficiency (BGE) is the proportion of C entering the bacterial pool that is incorporated into the biomass, and was calculated as $\text{BGE} = \text{HBP}/\text{BCD}$. The absence of size overlapping between phytoplankton and bacteria in the high-UVR lake (Medina-Sánchez et al., 2002) allowed for a direct measurement of BR. This, however, was not possible in the low-UVR lake, since autotrophic picoplankton and bacteria co-existed in the $< 3 \mu\text{m}$ fraction. Therefore, BCD in this lake was estimated by assuming that BR values lie within two limits: (i) a conservative value of 75 % of TPR, which is an average value based on data reported for oligotrophic waters (Lemeé et al., 2002), and (ii) a potential minimum value of 50 % of TPR (Robinson, 2008), comparable with direct measurements made in this study on the TPR vs. BR in La Caldera (Herrera et al., unpublished data).

2.7 Data calculation and statistical analysis

The effect size of the UVR was quantified as

$$\text{Effect size of UV-B (\%)} = 100 \times \left[\frac{((C_P - C_{\text{PAB}})/C_P) - ((C_P - C_{\text{PA}})/C_P)}{C_P} \right], \quad (2)$$

$$\text{Effect size of UV-A (\%)} = 100 \times \left[(C_P - C_{\text{PA}})/C_P \right], \quad (3)$$

where C_P , C_{PA} , and C_{PAB} represent the carbon production by phytoplankton or bacteria in samples under the P, PA, and PAB treatments, respectively. We used propagation errors to calculate the variance of the effect size (as percentage) due to UV-B and UV-A. The change (Δ) in the effect size of UV-B and UV-A, between the subsurface and mixed treatments, was calculated as the difference of the effect size for each radiation band.

The effects of solar radiation quality ("UVR" factor) and stratification ("STRAT" factor) on the response variables were tested using two-way ANOVA. When the interactive effects were significant, a post hoc Bonferroni test was used to determine significant differences among treatments. The normality (Kolmogorov–Smirnov test) and homoscedasticity (Cochran–Hartley–Bartlett test) were checked for each data group before ANOVA application. HBP data from the hypolimnetic community in the low-UVR lake were log-transformed to meet ANOVA assumptions. Significance of

Table 1. Mean values of the main physical, chemical, and biological variables measured in the water column in La Caldera (high-UVR lake) and La Conceja (low-UVR lake). Values of vertical attenuation coefficients (kd , m^{-1}) in the UVR (305, 320, 380 nm) and photosynthetically active radiation (PAR, 400–700 nm) regions are shown. Values are mean (\pm SD) of concentrations for four (La Caldera) or six (La Conceja) depths of inorganic, total, and dissolved nitrogen (N) and phosphorus (P), chlorophyll a , and phytoplankton and bacterial abundances. TN: total nitrogen; TDN: total dissolved nitrogen; NO_3^- : nitrate; TP: total phosphorus; TDP: total dissolved phosphorus; SRP: soluble reactive phosphorus; Chl a : chlorophyll a concentration; PA: phytoplankton abundance; PB: phytoplankton biomass; BA: bacterial abundance; BB: bacterial biomass.

Variable	High-UVR lake	Low-UVR lake
kd_{305}	0.61	4.84
kd_{320}	0.52	2.53
kd_{380}	0.34	0.93
kd_{PAR}	0.25	0.28
TN (μ M)	21.50 ± 1.54	787.1 ± 10.7
TDN (μ M)	20.71 ± 1.46	786.4 ± 12.9
NO_3^- (μ M)	14.28 ± 1.02	702.1 ± 6.7
TP (μ M)	0.10 ± 0.003	0.06 ± 0.012
TDP (μ M)	0.051 ± 0.002	0.038 ± 0.012
SRP (μ M)	0.02 ± 0.001	0.018 ± 0.012
Chl a (μ g L^{-1})	2.02 ± 0.42	2.66 ± 0.46
PA ($cell\ mL^{-1}$) $\times 10^3$	7.03 ± 1.65	4.03 ± 0.72
PB (μ g CL^{-1})	15.10 ± 4.31	95 ± 5.72
BA ($cell\ mL^{-1}$) $\times 10^6$	1.94 ± 0.17	1.28 ± 0.21
BB (μ g CL^{-1})	8.66 ± 1.32	0.98 ± 0.03

the effect size of UV-B and UV-A on PP and HBP between subsurface and mixed conditions was evaluated by means of a t test. Regression analyses were done to assess the dependence of the BGE on the EOC rates for the experimental data in each lake. Statistica 7.1 software for Windows was used for the statistical analyses.

3 Results

3.1 Physical, chemical, and biological variables in the water column

The lakes greatly differed in their transparency to UVR, but not to PAR (Table 1). Thus, in the high-UVR lake, 1% of the surface energy at 305 nm reached the bottom of the lake, whereas in the low-UVR lake most of the UVR energy was attenuated in the upper layers (1% of the surface energy at 305 nm reached only ca. 1 m depth). This differential penetration of solar UVR resulted in two contrasting environments, with organisms being exposed to UV-B throughout the water column in the high-UVR lake but only in the upper 1–2 m of the water column in the low-UVR lake. This was related to the different DOC concentrations in the lakes,

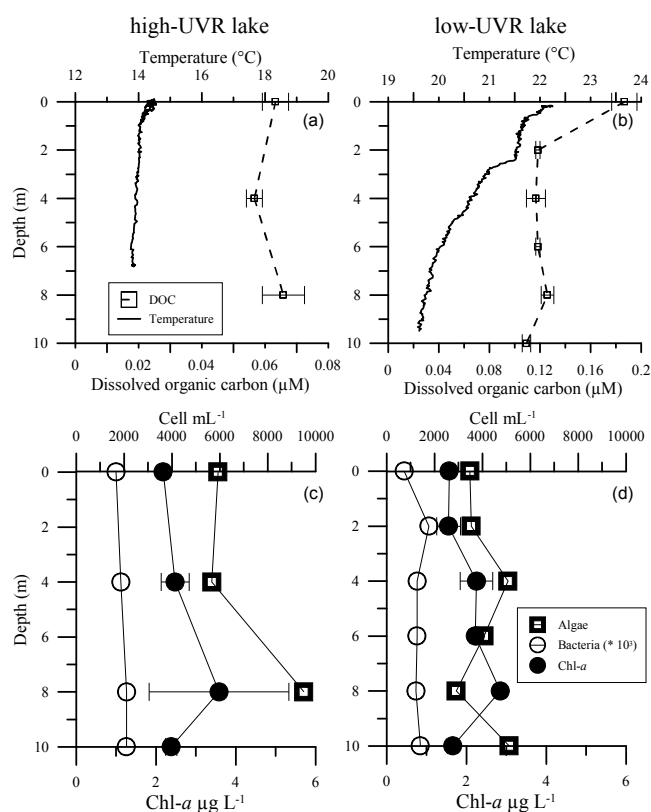


Figure 1. Water column characteristics of the high-UVR lake (a, c) and low-UVR lake (b, d). Depth profiles of temperature ($^{\circ}$ C) and dissolved organic carbon (DOC in μ M) (a, b), as well as phytoplankton and bacterial abundances ($cell\ mL^{-1}$) and chlorophyll a concentration (μ g L^{-1}) (c, d). Each symbol represents the mean of triplicate samples, while the horizontal error lines are the standard deviation.

reaching values of 0.07 and 0.18 mM in the high-UVR and low-UVR lakes, respectively (Fig. 1a, b). Vertical temperature profiles also differed between the lakes: the temperature was 14 $^{\circ}$ C, ranging by only 0.4 $^{\circ}$ C between the surface and bottom in the high-UVR lake, whereas a weak thermal stratification between 2 and 3 m was detected in the low-UVR lake, where the temperature ranged from 22 to 19.5 $^{\circ}$ C between the surface and bottom layers (Fig. 1a, b).

The concentrations of total dissolved and inorganic forms of N and P were homogeneous in the water column in both lakes; therefore, only mean values are reported in Table 1. TN values were higher in the low-UVR than in the high-UVR lake by up to 1 order of magnitude, and NO_3^- constituted most of the TN (90% in the low-UVR and 68% in the high-UVR lake). By contrast, TP values were $< 0.16\ \mu$ M and mostly in organic form in both lakes. The NO_3^- : TP ratio was > 100 in the high-UVR lake and $> 10\ 000$ in the low-UVR lake, indicating a strong P limitation (Table 1).

Chl a concentrations had small variations with depth in both lakes (Fig. 1c, d). However, the vertical distribution of

Table 2. Mean irradiances during the incubations for 305 nm, 320 nm and 380 nm within the UVR wavelengths ($\mu\text{W cm}^{-2} \text{nm}^{-1}$) and for PAR ($\mu\text{mol photons m}^{-2} \text{s}^{-1}$). The ratio of the mean irradiances of 380 and 305 nm is also presented.

Wavelength		305 nm	320 nm	380 nm	PAR	UV-A ₃₈₀ : UV-B ₃₀₅
High-UVR lake	Subsurface	3.90	23.40	60.10	1480	15.41
	Mixed	1.40	9.50	31.50	900	22.50
Low-UVR lake	Subsurface	1.44	12.90	47.90	1428	33.26
	Mixed	0.16	1.80	12.80	824	80.00

Table 3. Results of the two-way ANOVA of the interactive effect of “UVR” (PAB, PA, P, dark) and “STRAT” factors on carbon incorporation by phytoplankton (PP, $\mu\text{g CL}^{-1} \text{h}^{-1}$), and excreted organic carbon (EOC, $\mu\text{g CL}^{-1} \text{h}^{-1}$), heterotrophic bacterial production (HBP, $\mu\text{g CL}^{-1} \text{h}^{-1}$), bacterial respiration (BR, $\mu\text{g CL}^{-1} \text{h}^{-1}$; directly measured in the high-UVR lake or calculated as 50 % of total planktonic respiration (TPR) in the low-UVR lake), bacterial growth efficiency (BGE), and ratio of bacterial carbon demand to excretion of organic carbon (BCD : EOC) as a percentage. Numbers in bold indicate $p < 0.05$. df1, df2, and df3, df4, are the degrees of freedom.

		df ₁	df ₂	PP		EOC		df ₃	df ₄	HBP		BR		BGE		BCD : EOC (%)				
				$F_{df1,df2}$	p	$F_{df1,df2}$	p			$F_{df3,df4}$	p	df ₁	df ₂	$F_{df1,df2}$	p	$F_{df1,df2}$	p			
High-UVR lake	Epilimnetic	STRAT	1	12	42.29	<0.001	44.00	<0.001	1	16	6.41	0.022	1	12	1.07	0.321	0.26	0.619	6.15	0.029
		UVR	2	12	124.12	<0.001	6.33	0.013	3	16	8.65	0.001	2	12	12.38	0.001	7.22	0.009	35.47	<0.001
		UVR × STRAT	2	12	20.90	<0.001	0.11	0.895	3	16	5.46	0.009	2	12	3.71	0.056	4.80	0.029	14.59	0.001
Low-UVR lake	Epilimnetic	STRAT	1	12	0.61	0.450	2.46	0.143	1	16	7.37	0.015	1	8	5.28	0.05	1.45	0.263	18.76	0.002
		UVR	2	12	6.78	0.011	9.78	0.003	3	16	27.96	<0.001	1	8	0.14	0.72	46.13	<0.001	14.42	0.005
		UVR × STRAT	2	12	16.71	<0.001	16.51	<0.001	3	16	6.38	0.005	2	8	0.63	0.45	0.06	0.810	44.86	<0.001
Hypolimnetic	UVR	STRAT	2	12	0.33	0.574	4.33	0.060	1	16	32.98	<0.001	1	8	0.29	0.604	6.01	0.040	4.65	0.063
		UVR	2	12	41.58	<0.001	52.75	<0.001	3	16	12.05	<0.001	1	8	8.39	0.020	0.15	0.711	0.81	0.394
		UVR × STRAT	2	12	0.39	0.688	3.21	0.076	3	16	7.98	0.002	2	8	0.90	0.372	5.24	0.061	1.99	0.196

phytoplankton and bacteria differed between the lakes: in the high-UVR lake (Fig. 1c) bacterial abundance was rather homogeneous, but phytoplankton abundance was higher at the deepest depth; however, in the low-UVR lake (Fig. 1d) the abundances of bacteria and phytoplankton were rather uniform with depth. Mean phytoplankton and bacterial abundance values were greater in the high-UVR than in the low-UVR lake (Table 1). In terms of taxonomic composition, the Chlorophyceae *Monoraphidium* sp. represented > 90 % and ~ 80 % of the total abundance and biomass, respectively, in the high-UVR lake, whereas the Bacillariophyceae *Cyclotella ocellata* was the dominant species in the low-UVR lake (> 75 % abundance and 95 % biomass).

3.2 Variations in solar mean irradiance during experiments

The mean irradiance for the three wavelengths within the UVR and PAR region received by the samples under the experimental conditions are shown in Table 2. The mean irradiance at 305, 320, and 380 nm in the high-UVR lake were 2.8-, 2.5-, and 1.9-fold higher, respectively, in the subsurface than in the mixed conditions. The ratios between subsurface and mixed treatments in the low-UVR lake were 8.7-, 7.1-, and 3.7-fold higher for the 305, 320, and 380 nm wavelengths, respectively. The energy ratio at 380 and 305 nm (i.e., UV-A₃₈₀ : UV-B₃₀₅ ratio) had higher values in the low-UVR lake as compared to the high-UVR lake, reflecting the lower penetration of UV-B in the former.

3.3 Joint effects of UVR and stratification on phytoplanktonic and bacterial metabolism in the high-UVR lake

The PP values did not show significant differences between subsurface and mixed conditions in the PAB treatment, while samples under the PA and P treatments had significantly higher PP values at the subsurface than under mixed conditions (Fig. 2a). A significant UVR × STRAT effect was found for PP (Table 3), and according to our hypothesis the subsurface incubations resulted in higher UV-B (11.5 %) and UV-A (18.3 %) inhibition as compared to the mixed incubations (Table 4). UVR at the subsurface also significantly increased the rates of EOC, with significantly higher values in samples under the PAB and PA treatments (Fig. 2b). Likewise, the %EOC was significantly affected by UV-B, increasing to 22 and 21 % in subsurface and in mixed treatments, respectively (graph (1) in Appendix A). Like PP, HBP did not differ between PAB-subsurface and PAB-mixed treatments. However, HBP was significantly lower under PA-subsurface than under PA-mixed treatments (Fig. 2c), resulting in a significant UVR × STRAT effect (Table 3). By contrast, only the “UVR” factor significantly affected BR (Fig. 2d, Table 3), with the lowest BR value determined in the PAB-subsurface treatment (Fig. 2d). BGE had higher values in the PAB-subsurface treatment as compared to the other radiation treatments under subsurface conditions; other comparisons between paired treatments did not result in

Table 4. Effect size of UV-B and UV-A on primary production (PP, $\mu\text{gCL}^{-1}\text{h}^{-1}$) and heterotrophic bacterial production (HBP, $\mu\text{gCL}^{-1}\text{h}^{-1}$) in the effect size of UV-B and UV-A, between the subsurface and mixed treatments, was calculated as the difference of the effect size for each radiation band. Numbers in bold indicate $p < 0.05$.

		PP				HBP			
		%UV-B	$\Delta\%$ UV-B	%UV-A	$\Delta\%$ UV-A	%UV-B	$\Delta\%$ UV-B	%UV-A	$\Delta\%$ UV-A
High-UVR lake									
Epilimnetic	Subsurface	37.3 ± 2.4	11.55	25.6 ± 7.6	18.32	2.7 ± 18.3	-20.3	51.9 ± 26.7	110.2
	Mixed	25.7 ± 5.0		7.3 ± 7.1		23.0 ± 1.5		-58.3 ± 0.2	
Low-UVR lake									
Epilimnetic	Subsurface	33.7 ± 4.2	40.00	17.4 ± 13.9	27.41	42.9 ± 6.2	-4.2	30.0 ± 8.7	1.2
	Mixed	-6.3 ± 10.9		-10.0 ± 23.5		47.1 ± 2.0		28.2 ± 6.7	
Hypolimnetic	Subsurface	27.2 ± 22.5	0.09	20.8 ± 28.9	-5.98	52.1 ± 5.8	45.6	12.0 ± 24.4	-11.5
	Mixed	27.1 ± 5.6		26.8 ± 12.8		6.5 ± 12.2		23.6 ± 2.6	

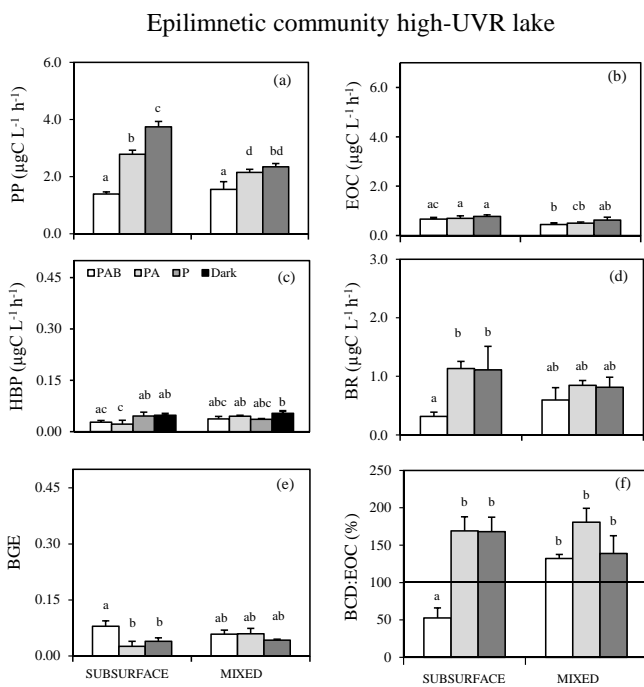


Figure 2. Metabolic variables of epilimnetic community under different radiation treatments (PAB, PA, P, dark) and different stratification conditions (subsurface, mixed) in the high-UVR lake. (a) Primary production (PP, in $\mu\text{gCL}^{-1}\text{h}^{-1}$), (b) rates of excretion of organic carbon (EOC, in $\mu\text{gCL}^{-1}\text{h}^{-1}$), (c) heterotrophic bacterial production (HBP, in $\mu\text{gCL}^{-1}\text{h}^{-1}$), (d) bacterial respiration (BR, in $\mu\text{gCL}^{-1}\text{h}^{-1}$), (e) bacterial growth efficiency (BGE), and (f) ratio of bacterial carbon demand to excretion of organic carbon (BCD:EOC) as a percentage. The horizontal line in (f) represents a balanced commensalistic phytoplankton–bacteria relationship. The vertical error lines on top of the bars are the standard deviation, whereas the letters indicate differences among treatments.

significant differences of BGE (Fig. 2e). There was, nevertheless, a significant UVR \times STRAT interaction on BGE

(Table 3). No relationship was found between EOC rate and BGE ($R^2 = 0.149$ $p > 0.05$). Finally, to quantify the capacity of C released by phytoplankton to support the bacterial C demand (BCD) in each treatment, the BCD:EOC ratio (as a percentage) was calculated (Fig. 2f). Carbon released by phytoplankton resulted in excess to meet BCD (i.e., BCD:EOC values $< 100\%$) only in the PAB-subsurface treatment (Fig. 2f).

3.4 Joint effects of UVR and stratification on phytoplanktonic and bacterial metabolism in the low-UVR lake

UVR exerted negative effects on both epilimnetic (Fig. 3) and hypolimnetic (Fig. 4) communities. For the epilimnetic community, PP was significantly lower in the PAB than in PA and P treatments under subsurface conditions, while UVR did not affect PP under mixed conditions (Fig. 3a). A significant UVR \times STRAT effect on PP (Table 3) was found, with the lowest PP values for the PAB-subsurface treatment. The highest values of UV-B (37%) and UV-A (25%) inhibition were found at the subsurface (Table 4). As for PP, EOC was significantly lower in the PAB than in the PA and P treatments at the subsurface, but non-significant differences among radiation treatments under mixed conditions were found (Fig. 3b). %EOC did not show differences among radiation treatments in both stratification conditions (graph (2) in Appendix A). HBP showed significantly higher values in dark treatments at the subsurface than under mixed conditions (Fig. 3c), generating a significant interactive effect of UVR \times STRAT on HBP (Table 3). Noticeably, a strong inhibition of HBP by UV-B and UV-A in subsurface and in mixed conditions was found (Table 4). By contrast, the estimated BR was not significantly affected by any factor (Table 3; Fig. 3d shows $\text{BR}_{50\%}$). UVR was the only factor that significantly reduced BGE values in both mixed and subsurface conditions (Fig. 3e). No relationship between EOC rate and BGE was found ($R^2 = 0.055$ $p > 0.05$). The BCD:EOC (%)

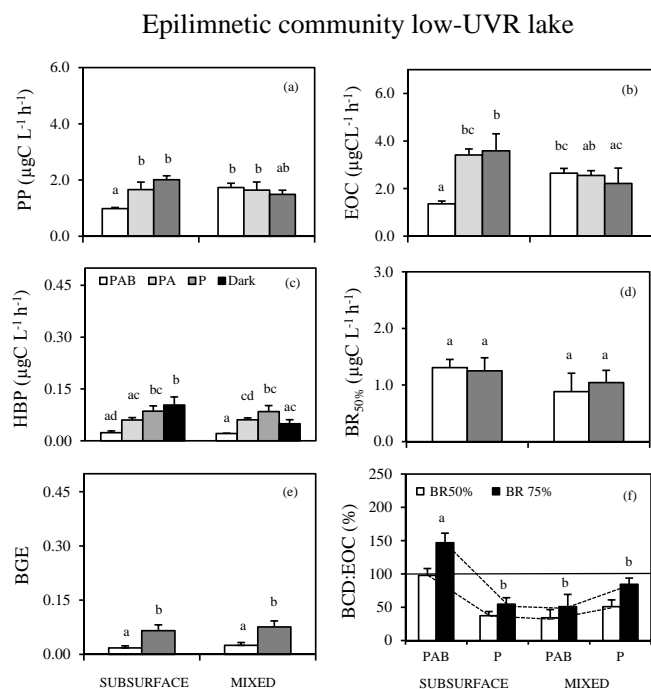


Figure 3. Metabolic variables of epilimnetic community under different radiation treatments (PAB, PA, P, dark) and different stratification conditions (subsurface, mixed) in the low-UVR lake. (a) Primary production (PP, in $\mu\text{gC L}^{-1} \text{h}^{-1}$); (b) rate of excretion of organic carbon (EOC, in $\mu\text{gC L}^{-1} \text{h}^{-1}$); (c) heterotrophic bacterial production (HBP, in $\mu\text{gC L}^{-1} \text{h}^{-1}$); (d) bacterial respiration ($\text{BR}_{50\%}$ in $\mu\text{gC L}^{-1} \text{h}^{-1}$), calculated as 50% of total planktonic respiration (TPR); (e) bacterial growth efficiency (BGE); and (f) ratio of bacterial carbon demand to excretion of organic carbon (BCD:EOC) as a percentage. The broken lines indicate the min-max range of the BCD:EOC(%) ratio, with BCD calculated assuming BR as either 50 or 75% of TPR. The horizontal line in (f) represents a balanced commensalistic phytoplankton–bacteria relationship. The vertical error lines on top of the bars are the standard deviation, whereas the letters indicate differences among treatments.

was $< 100\%$ for every experimental condition except for that under PAB-subsurface treatment, where the BCD:EOC (%) reached values from $\sim 100\%$ (assuming $\text{BR} = 50\%$ of TPR) to 145% (assuming $\text{BR} = 75\%$ of TPR) (Fig. 3f). Thus, in this latter case (PAB-subsurface), EOC was not enough to meet BCD.

For the hypolimnetic community, UVR was the only factor that significantly inhibited PP (Fig. 4a). Samples under the PAB and PA treatments had significantly lower PP values than those under the P treatment in both subsurface and mixed conditions (Fig. 4a). The EOC rates were significantly lower in the PAB and PA treatments than in the P treatment at the subsurface (Fig. 4b). No significant differences among both stratification treatments were determined when comparing each radiation treatment (Fig. 4b). HBP was only significantly inhibited by UV-B (Fig. 4c), whereas it was stimu-

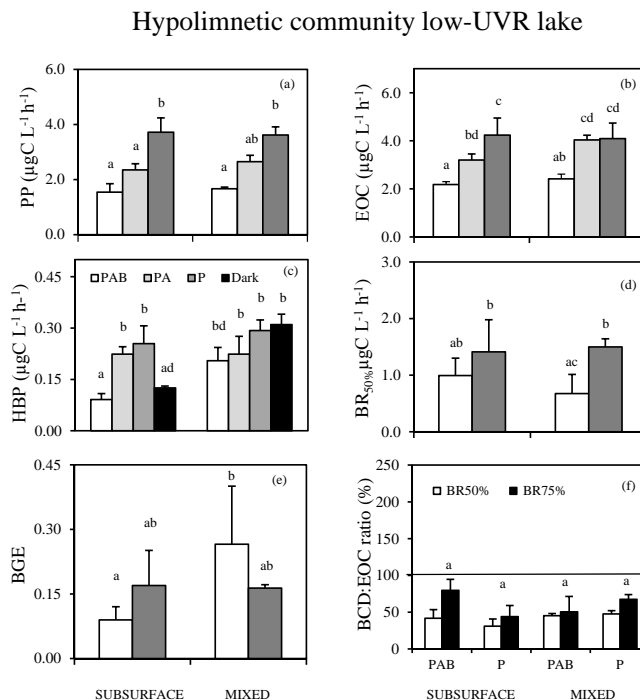


Figure 4. Metabolic variables of hypolimnetic community under different radiation treatments (PAB, PA, P, dark) and different stratification conditions (subsurface, mixed) in the low-UVR lake. (a) Primary production (PP, in $\mu\text{gC L}^{-1} \text{h}^{-1}$); (b) rates of excretion of organic carbon (EOC, in $\mu\text{gC L}^{-1} \text{h}^{-1}$); (c) heterotrophic bacterial production (HBP, in $\mu\text{gC L}^{-1} \text{h}^{-1}$); (d) bacterial respiration ($\text{BR}_{50\%}$ in $\mu\text{gC L}^{-1} \text{h}^{-1}$), calculated as 50% of total planktonic respiration (TPR); (e) bacterial growth efficiency (BGE); and (f) ratio of bacterial carbon demand to excretion of organic carbon (BCD:EOC) as percentage, calculated assuming BR as either 50 or 75% of TPR. The horizontal line in (f) represents a balanced commensalistic phytoplankton–bacteria relationship. The vertical error lines on top of the bars are the standard deviation, whereas the letters indicate differences among treatments.

lated by PA and P under subsurface conditions (Fig. 4c). Under mixing, however, UVR did not affect HBP. Therefore, subsurface exposure triggered the inhibition due to UV-B by 45.6% (Table 4). Only UVR, as a single factor, significantly affected BR (Table 3), with the lowest values under the PAB-mixed treatment (Fig. 4d), whereas only the STRAT factor affected BGE, with the lowest BGE values in the PAB-subsurface treatment (Fig. 4e). The BCD:EOC(%) was $< 100\%$ under all conditions (assuming $\text{BR} = 50$ or 75% of TPR), indicating the EOC was always capable of supporting BCD (Fig. 4f).

To summarize, and taking into account the changes (Δ) in the inhibitory UVR effect (UV-B and UV-A) on PP and HBP with increased stratification (Table 4), our results reveal greater UV-B sensitivity of (i) epilimnetic phytoplankton and heterotrophic bacteria communities in the low-UVR lake than in the high-UVR lake, (ii) epilimnetic phytoplankton

than heterotrophic bacteria in both lakes, and (iii) hypolimnetic heterotrophic bacterial than phytoplankton community in the low-UVR lake. In addition, significant interactive UVR \times STRAT effects were observed on the BCD : EOC (%) only in the epilimnetic communities (Table 3). Thus, partially supporting our hypothesis, the BCD : EOC (%) significantly decreased under PAB-subsurface treatment in the high-UVR lake but increased in the low-UVR lake.

4 Discussion

The main outcome of our work is that the increased stratification of the water column altered the commensalistic phytoplankton–bacteria relationship in oligotrophic lakes. The present study is the first so far to directly assess the interactive effects of UVR and stratification changes on phytoplankton and bacteria and their commensalistic relationship in freshwater ecosystems. Furthermore, in our complex experimental approach, we simulated reductions in the depth of the UML due to the stratification of the water column (one of the potential consequences of global warming; Gao et al., 2012; de Senerpont-Domis et al., 2013). Under these conditions, we measured the extracellular carbon release by phytoplankton and directly determined the BR, as these are the key variables implied in the ratio of bacterial carbon demand to C supply. Moreover, since a strong feedback between physical processes (e.g., mixing, stratification) and changes in DOC concentration in small lakes has previously been reported (Read and Rose, 2013), we further achieved an advance in our knowledge by investigating two oligotrophic ecosystems that differed in their UVR penetration in the water column due to their DOC content, selected as model lakes representing two ends of an optical gradient of transparency to UVR in Mediterranean inland waters. This provides a framework for disentangling the complex processes that underlie biological interactions under changing physical (stratification, UVR) and chemical (DOC) conditions, which can then modify the C flux in aquatic ecosystems.

4.1 Sensitivity of phytoplankton and bacteria to UVR and stratification

Despite the physical and ecological differences between the two lakes, PP and HBP responses to the joint effect of UVR and stratification were quite similar in that the latter augmented the effect size of UV-B, mainly on the epilimnetic communities in both ecosystems. This effect reached a higher magnitude in the low-UVR lake (Table 4), which coincided with a greater relative exposure to UV-B (9-fold) and an more accentuated decrease in the UV-A : UV-B ratio (58 %) at the shallower layer in the low-UVR than in the high-UVR lake. This result agrees with the findings of higher UVR damage to primary producers in low-UVR lakes than in high-UVR lakes as reported by Helbling et al. (2013), although in their study

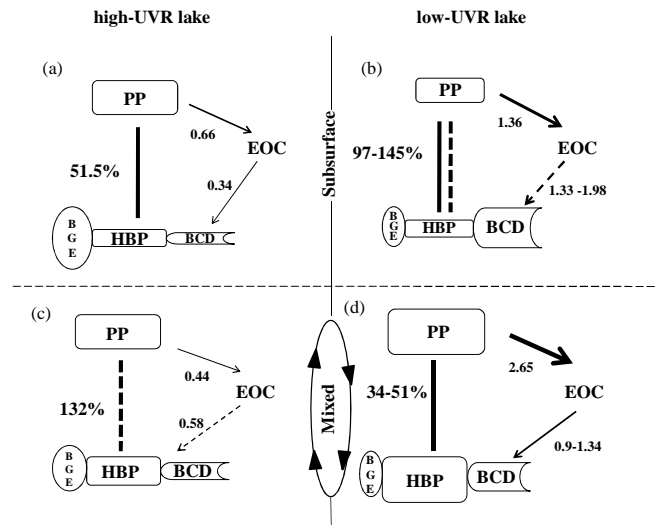


Figure 5. Epilimnetic phytoplankton–bacteria relationship under PAB-subsurface and PAB-mixed conditions in the high-UVR lake (a, c) and low-UVR lake (b, d). The sizes of the boxes are proportional to the magnitude of the rates (in $\mu\text{g C L}^{-1} \text{h}^{-1}$). The absolute numbers are values of excretion of organic carbon (EOC) rates and bacterial carbon demand (BCD), and the percentage numbers are values of the BCD : EOC (%) ratio. The thicknesses of the arrows indicate the relative magnitude of a particular carbon flux. The broken-line arrows indicate that EOC is not enough to satisfy BCD. Thick black lines represent the BCD : EOC (%) ratio, indicating either coupled (solid lines) or uncoupled (broken lines) phytoplankton–bacteria relationship. PP: primary production; HBP: heterotrophic bacterial production; BGE: bacterial growth efficiency. Numbers are rates of C flux (in $\mu\text{g C L}^{-1} \text{h}^{-1}$).

this response was found only under fluctuating irradiances. The results presented here indicate increased susceptibility to UVR of bacteria and phytoplankton communities relatively less exposed to UV-B during their life cycles (Pakulski et al., 2007; Harrison and Smith, 2011a). Interestingly, the UVR effect on %EOC was only significant in the high-UVR lake; the release of C has been described as a protective mechanism to prevent photosystem damage from reducing power excess under high irradiance of PAR (Wood and Van Valen, 1990) as well as UVR (Carrillo et al., 2002, 2008). The lack of this “escape valve”, which helps to prevent over-excitation of PSII, might be the final cause of the higher sensitivity of phytoplankton communities in the low-UVR lakes. In addition, a higher sensitivity to UVR was found for epilimnetic phytoplankton than for bacteria mainly under subsurface conditions, suggesting that photosynthetic processes are more sensitive under extreme conditions that mimic the global-warming scenario. This result contrasts with previous reports of greater UVR damage to bacterioplankton than to phytoplankton in oligotrophic waters of the Mediterranean Sea (Bertoni et al., 2011), the northern South China Sea (Yuan et al., 2011), high-mountain lakes (Sommaruga et al., 1997), and boreal lakes (Xenopoulos and Schindler, 2003).

Taken all together, our results show that increased stratification, by trapping the cells in a shallower epilimnion, with increased UVR exposure triggered or exacerbated the inhibitory effect of UVR on phytoplanktonic and bacterial metabolism measured under mixed conditions. Because this negative effect was greater in high-DOC waters, we propose that the “ideal” photoprotective DOM may become harmful on planktonic communities in a scenario of increased stratification. Our proposal is based on the indirect harmful UV-B effects due to the free radicals (O_2^- , H_2O_2 , OH^-) generated by photooxidation of the DOC (Banaszak, 2003; Pullin et al., 2004), which can exacerbate the negative UVR effect in low-UVR lakes. In addition, DOC would become bleached and therefore the lake would be more UVR transparent (Reche et al., 2001), thus increasing the negative effect of UVR on organisms. However, cell acclimation to UVR or a shift in the taxonomic composition towards UVR-resistant species could counteract the net negative UVR effect on a long-term scale.

As expected, UVR was the main factor that affected the non-acclimated hypolimnetic community, since PP and HBP underwent negative UV-B and UV-A effects in both subsurface and mixed conditions (Table 4). These responses reflect the higher sensitivity of the hypolimnetic than the epilimnetic community to UVR, because only the hypolimnetic community was negatively affected by UVR under mixed conditions. These results agree with previous reports of higher photosynthetic impairment under UVR exposure of phytoplankton from deep chlorophyll maxima (Harrison and Smith, 2011b) or from the bottom of the mixed layer (Xenopoulos and Schindler, 2003).

Nevertheless, HBP of the hypolimnetic community was stimulated by UV-A and PAR when exposed to shallower conditions. These results suggest that the hypolimnetic bacteria possessed photorepair mechanisms, via UV-A and PAR-promoted photolyase activity (DNA repair), which may be activated after 4 h of UVR and PAR exposure (Jeffrey et al., 1996; Bertoni et al., 2012). This photorepair mechanism has a low energy cost and may be an important adaptive mechanism to attenuate the gross negative effect of UVR when a non-UVR-acclimated bacterioplankton community is exposed to high PAR and UV-A intensity and harmful UV-B levels in ecosystems with low nutrient availability (Medina-Sánchez et al., 2002). Nevertheless, and in agreement with our hypothesis, photorepair mechanisms were insufficient to completely counteract UVR-induced damage, this being concordant with a sharp decrease in the UV-A : UV-B ratio (58 %) in the upper layers (subsurface conditions). Moreover, the increased HBP found after exposure of samples to higher PAR intensity in the upper layers is consistent with the previously reported stimulatory effect of PAR on HBP (Morán et al., 2001; Medina-Sánchez et al., 2002; Pakulski et al., 2007). Furthermore, a potential presence of aerobic anoxygenic phototrophic bacteria (Bertoni et al., 2011; Mašín et al., 2012; Ferrara et al., 2011) should not be ruled out in account-

ing for the increased HBP under high PAR in the low-UVR lake.

4.2 UVR and increased stratification effect on the commensalistic phytoplankton–bacteria relationship

As noted above, UVR and stratification exerted an interactive effect on PP and HBP in the epilimnetic layer in both lakes. However, this interactive effect was only exerted on EOC in the low-UVR lake, where the EOC rates values were 3-fold higher (except under PAB-subsurface treatment) than in the high-UVR lake. The carbon released by phytoplankton is composed mainly of low-molecular-weight compounds that are readily assimilable by bacteria (Amon et al., 2001). This source of carbon is preferred by bacteria, even in lakes with considerable input of terrestrial carbon to subsidize their growth (Kritzberg et al., 2005, 2006), because the non-readily assimilable organic matter, mostly composed of high-molecular-weight compounds, must be hydrolyzed by bacterial ectoenzymes before the assimilation.

Quantification of the dependence of heterotrophic bacteria on organic substrate released by phytoplankton requires an accurate assessment of the BCD (Morán et al., 2002). Our study offers a quite precise estimate of the BCD, because both HBP and BR were directly measured in the high-UVR lake, due to the absence of size overlap between auto- and heterotrophic organisms. In the low-UVR lake, where segregation between both biological fractions was not feasible, BR was estimated from direct measurements of TPR and the reported percentages of the latter variable accounted for BR (i.e., 50 and 75 %; Lemeé et al., 2002; Robinson, 2008). This procedure brought about a min–max range where the actual BR should safely fall. In addition, its reliability is supported in that our estimated mean BGE and BR values fell within the range reported for oligotrophic ecosystems (Biddanda et al., 2001; Amado et al., 2013).

In the high-UVR lake, BGE was increased under full-sunlight and subsurface conditions, reflecting greater changes in bacterial respiration than in production. The reduction in BR and, as a consequence, the increase in bacterial growth efficiency could be interpreted as a tolerance-related mechanism under full-sunlight exposure in accordance with the non-inhibitory effect of UV-B on HBP found under shallower conditions. By contrast, in the low-UVR lake, BGE values were lower under full-sunlight and subsurface (stratified) conditions. The lack of the inhibitory effect of full-sunlight (PAB vs. P) on TPR (and hence BR) concomitantly with a strong inhibitory effect of UV-B on HBP determined a reduction in bacterial growth efficiency according to the high sensitivity of the bacterial community. The differences in the bacterial responses between the lakes could be the outcome of specific bacterial composition inhabiting each lake. These results agree with previous laboratory findings of a negative UV-B effect on BGE or BR in some bacterial strains isolated

from alpine lakes, but a positive effect on others, suggesting a strain-specific response (Hörtnagl et al., 2010). Nevertheless, changes in BGE are frequently observed when bacterial growth is limited by substrate availability (del Giorgio and Cole, 1998; López-Urrutia and Morán, 2007). Although our experiments were not specifically designed to test the role of organic substrates on BGE, we did not find a significant direct relationship between EOC rate and BGE in each lake. Thus, our data support the view that BGE can be altered by direct solar UVR impact.

Regarding the commensalistic phytoplankton–bacteria relationship, it was noticeable that EOC rates in the high-UVR ecosystem increased with full-sunlight under subsurface conditions, reaching values that exceeded the C demand of a bacterial community which seemed to have undergone an inactivation or dormancy under PAB, reflected by lower respiration. This slowing of the bacterial metabolism, concomitant with an increase in the availability of C released by phytoplankton, was the mechanism that determined the “coupling” phytoplankton–bacteria relationship. However, the fate of the C released by phytoplankton could be a transitory accumulation in lake water until its consumption by enhanced bacterial metabolic processes (growth and respiration) after an improvement in the light conditions, or could be definitively incorporated into the dissolved-C pool of the lake water.

In the low-UVR ecosystem, particularly the epilimnetic community, the strong inhibitory effect of UV-B at the subsurface on PP (i.e., decreasing C incorporation) was also reflected in a lesser C release by phytoplankton under these conditions. These decreased EOC rates did imply a change in their capability to meet the BCD, which ranged from barely sufficiency (if a 50 % loss of TPR is assumed) to non-sufficiency (if a 75 % loss of TPR is assumed). Therefore, the estimated min–max interval for each experimental condition shows an unexpected trend to a weakening of the bacterial dependence on phytoplankton C under full-sunlight and subsurface conditions in the low-UVR lake, which may be induced by global warming. These results partially support our hypothesis because the interaction between UVR and stratification strengthened the commensalistic phytoplankton–bacteria relationship (decreasing the BCD : EOC(%) ratio to < 100) in the high-UVR lake, but weakened (increasing the BCD : EOC(%) ratio to > 100) this relationship in the low-UVR lake (Figs. 2f and 3f). Moreover, they underline the capability of UVR in altering the efficiency of phytoplankton C excretion to support bacterial demands in optically contrasting ecosystems. Since the interaction of UVR and simulated stratification on this crucial biotic interaction in high-UVR and low-UVR lakes has not been previously examined, more data are needed in order to generalize these responses by microbial organisms, not only on a short-term (as considered in this study) but also a long-term basis.

To summarize our findings, we propose a conceptual functioning model that embraces both contrasting model ecosystems (Fig. 5). According to the global-warming scenario, it is expected that (i) the vertical stratification of aquatic ecosystems will intensify (de Senerpont Domis et al., 2013), (ii) the depth of the mixed layer will be altered as a consequence of micro-stratification in shallow lakes (van de Waal et al., 2009), and (iii) microbial communities and DOC will be confined within a highly irradiated layer. Based on our results, the synergistic effect of UVR and increased stratification on the microbiota might strengthen the C flux through the microbial loop in the high-UVR lake (or increasing the DOC pool in the lake) but might weaken it in the low-UVR lake. Therefore, our results showing greater UVR damage in the low-UVR lake imply that these types of ecosystems might be especially vulnerable to these factors related to global change.

Appendix A

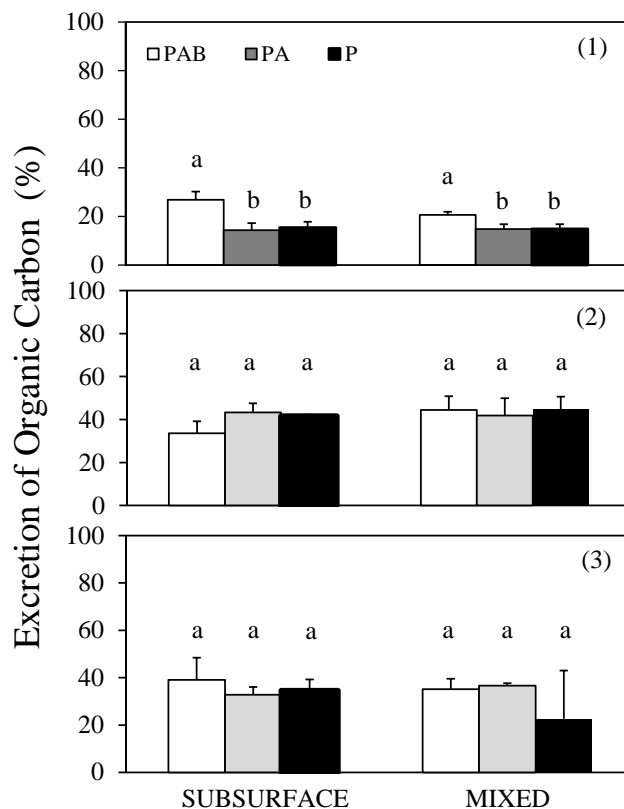


Figure A1. Percentage of excretion of organic carbon rates (%EOC) under different radiation (PAB, PA, P) and stratification conditions (subsurface, mixed) in (1) an epilimnetic community in the high-UVR lake, (2) an epilimnetic community in the low-UVR lake, and (3) a hypolimnetic community in the low-UVR lake. The vertical error lines on top of the bars are the standard deviation, whereas the letters indicate differences among treatments.

Author contributions. P. Carrillo, E. W. Helbling, and V. E. Villafañe conceived and designed the experiments. P. Carrillo, J. M. Medina-Sánchez, C. Durán, G. Herrera, E. W. Helbling, and V. E. Villafañe performed the experiments. P. Carrillo, J. M. Medina-Sánchez, and E. W. Helbling analyzed the data. P. Carrillo contributed reagents/materials/analysis tools. P. Carrillo, J. M. Medina-Sánchez, E. W. Helbling, and V. E. Villafañe wrote the paper.

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